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Physics of the ocean boundary layer below ice shelves: the relationship between modeling and observations

Xylar Asay-Davis

OASIIS Workshop, 15 June 2017



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More observations like Tim Stanton's, please!

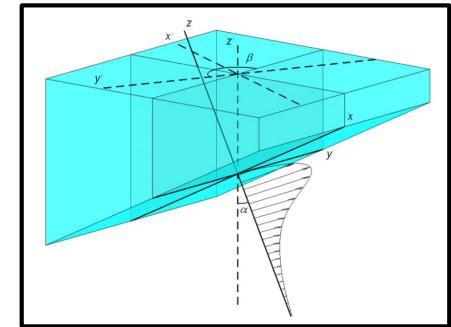
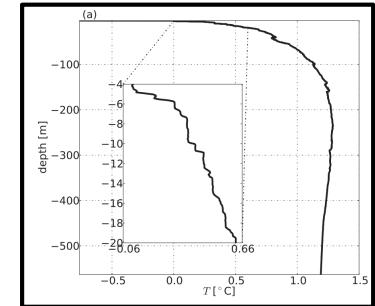
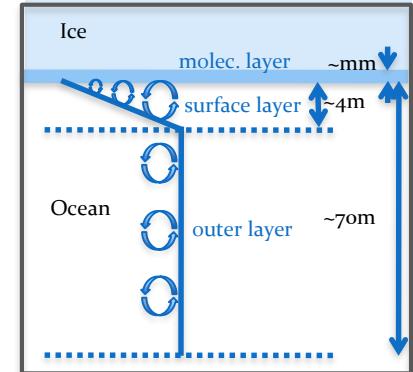
**Physics of the ocean boundary layer below ice shelves:
the relationship between modeling and observations**

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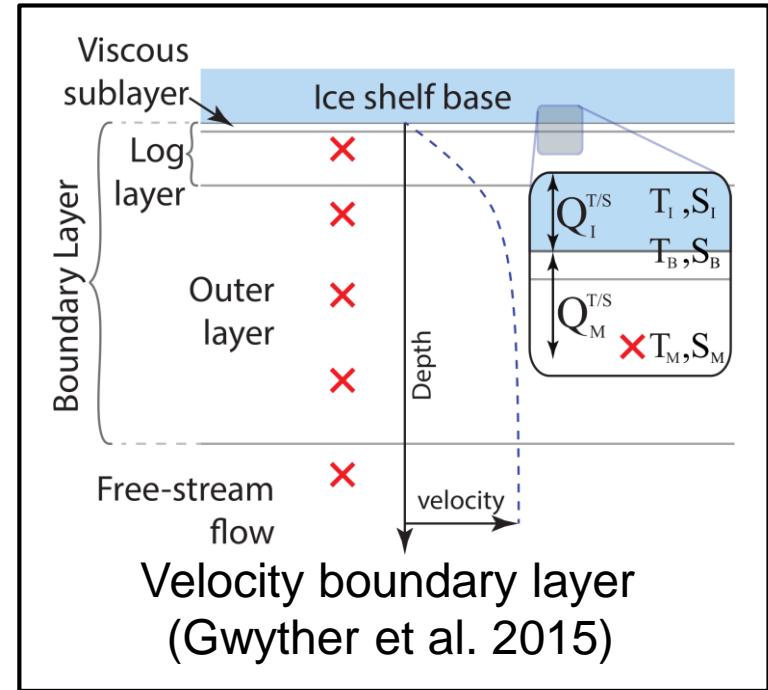
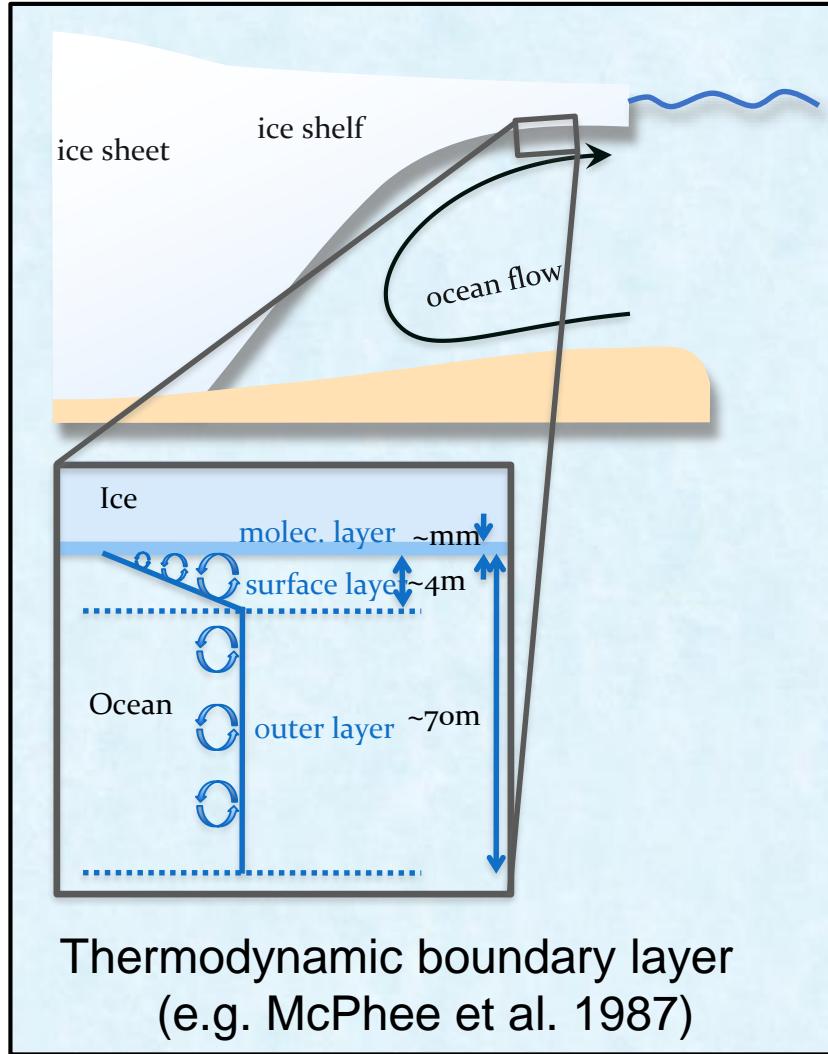
OASIIS Workshop, 15 June 2017

Outline

- Observations and theory behind boundary-layer parameterizations used in ocean models
- Observations, theory and modeling suggesting need for new parameterizations
- Components of an improved boundary-layer parameterization
- New observations needed to better constrain parameterizations



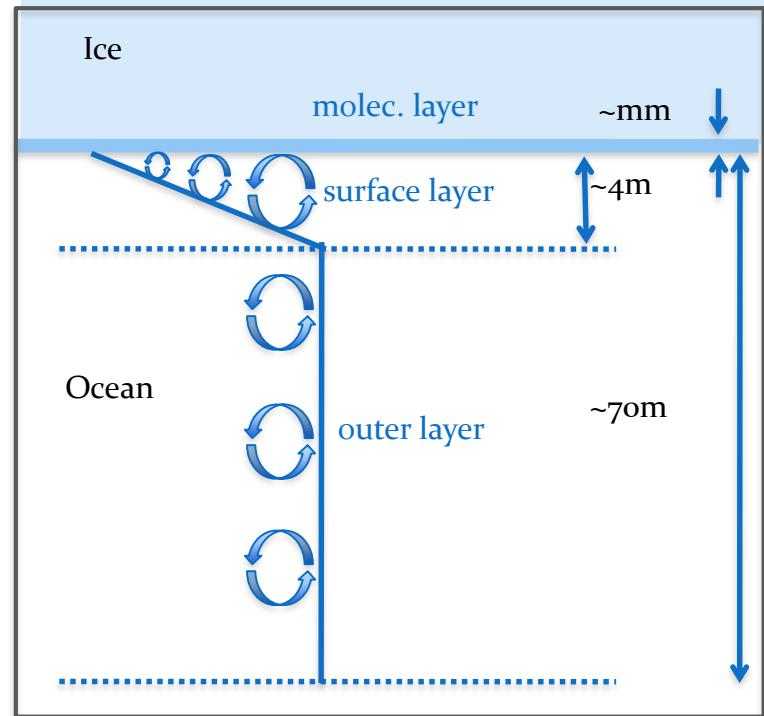
Theory and modeling of the boundary layer (BL)



- Surface roughness, molecular diffusivities dominant in surface layer
- Turbulence (therefore eddy viscosity/diffusivities) dominant in outer layer

Theory: McPhee et al. 1987

- Applied to sea ice (not ice shelves)
- 1D vertical model

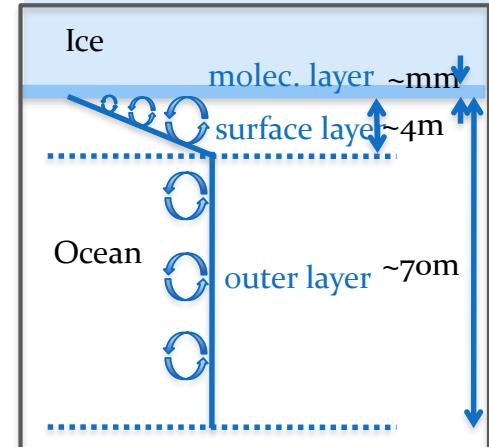


McPhee, M. G., Maykut, G. A., & Morison, J. H. (1987). Dynamics and Thermodynamics of the Ice/Upper Ocean System in the Marginal Ice Zone of the Greenland Sea. Journal of Geophysical Research, 92(C7), 7017–7031.



Theory: McPhee et al. 1987

- Constant heat, salt flux through BL



$$\langle w'T' \rangle_0 = wQ_L = -K_h \frac{\partial T}{\partial z} \quad (1)$$

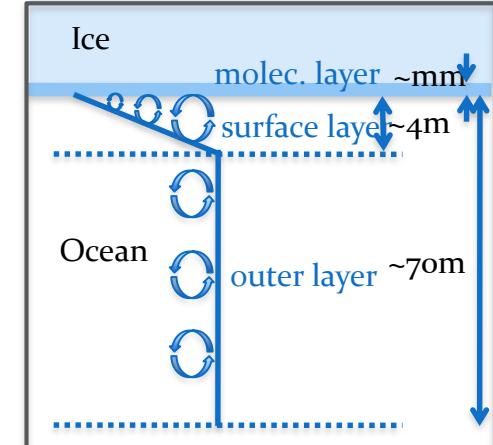
$$\langle w'S' \rangle_0 = w(S_w - S_i) = -K_s \frac{\partial S}{\partial z} \quad (2)$$

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Journal of Geophysical Research, 92(C7), 7017–7031.



Theory: McPhee et al. 1987

- Constant heat, salt flux through BL
- Diffusivities $K_h(z)$ and $K_s(z)$ vary with space:
 - Weak, molec. diffusion in molec. Layer
 - Transition to turbulence in surface layer
 - Fully turbulent mixing in outer layer



$$\langle w'T' \rangle_0 = wQ_L = -K_h \frac{\partial T}{\partial z} \quad (1)$$

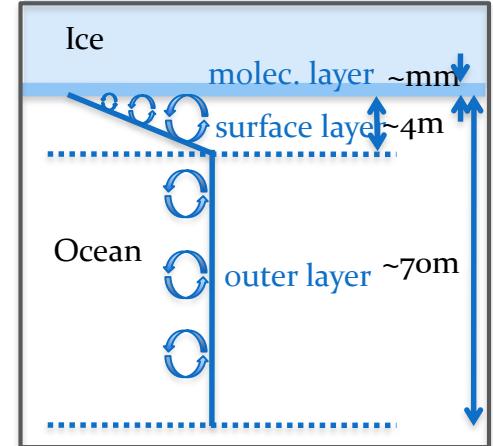
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Theory: McPhee et al. 1987

- Vertically integrate to get 2 of the 3 boundary conditions
- 3rd boundary condition is freezing point



$$\rho_i a_b L_i = \rho_i c_i \kappa_i \left. \frac{\partial T_i}{\partial z} \right|_b - \rho_w c_w u_* \Gamma_T [T_f(S_b, P_b) - T_w],$$

$$\rho_i a_b (S_b - S_i) = -\rho_w u_* \Gamma_S (S_b - S_w),$$

$$T_f = \lambda_1 S + \lambda_2 + \lambda_3 P.$$

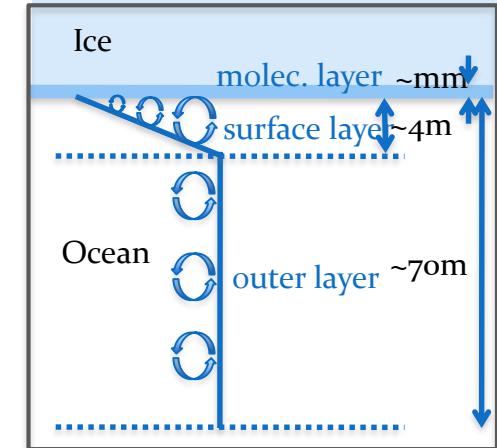
“Three equations” using notation from Jenkins et al. (2010)

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Theory: McPhee et al. 1987

- Vertically integrate to get 2 of the 3 boundary conditions
- 3rd boundary condition is freezing point
- Heat and salt transfer coeffs. $\Gamma_T(z_w)$ and $\Gamma_S(z_w)$
 - Parameterize transfer across BL
 - Functions of $K_h(z)$ and $K_s(z)$
 - Asymptote to constants as $z_w \rightarrow \infty$
- Similarly for drag coeff. $C_d(z_w)$



$$\rho_i a_b L_i = \rho_i c_i \kappa_i \left. \frac{\partial T_i}{\partial z} \right|_b - \rho_w c_w u_* \Gamma_T [T_f(S_b, P_b) - T_w],$$

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“Three equations” using notation from Jenkins et al. (2010)

$$u_*^2 = C_d U^2,$$

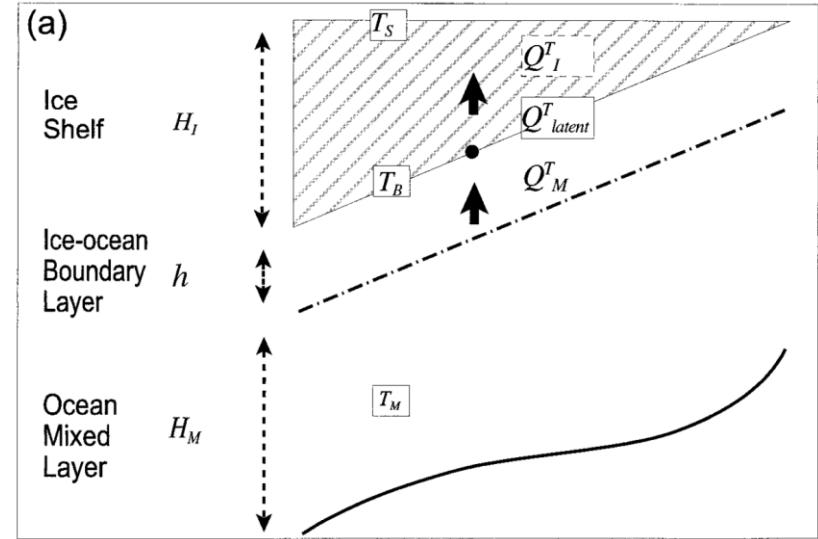
Quadratic drag law using notation from Jenkins et al. (2010)

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Modeling: Holland and Jenkins (1999)

- Applied McPhee et al. (1987) to ice shelves
- Currently most common BL parameterization in ocean models with ice-shelf cavities

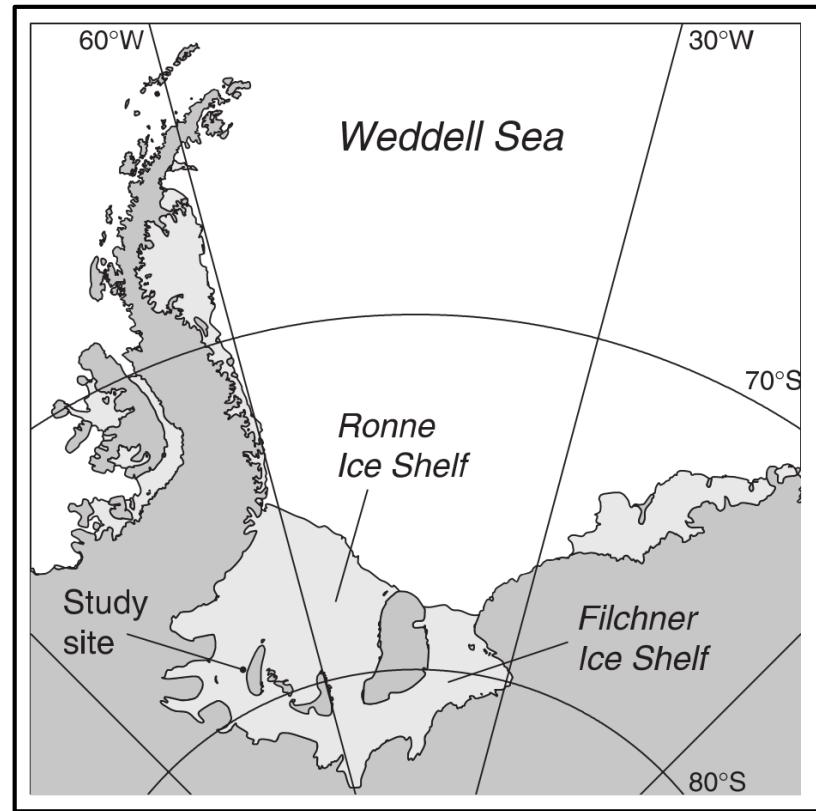


Holland, D. M., & Jenkins, A. (1999). Modeling Thermodynamic Ice–Ocean Interactions at the Base of an Ice Shelf. *Journal of Physical Oceanography*, 29(8), 1787–1800.



Observations: Nicholls et al. (1997); Corr et al. (2002)

- “Site 3” on Ronne Ice Shelf



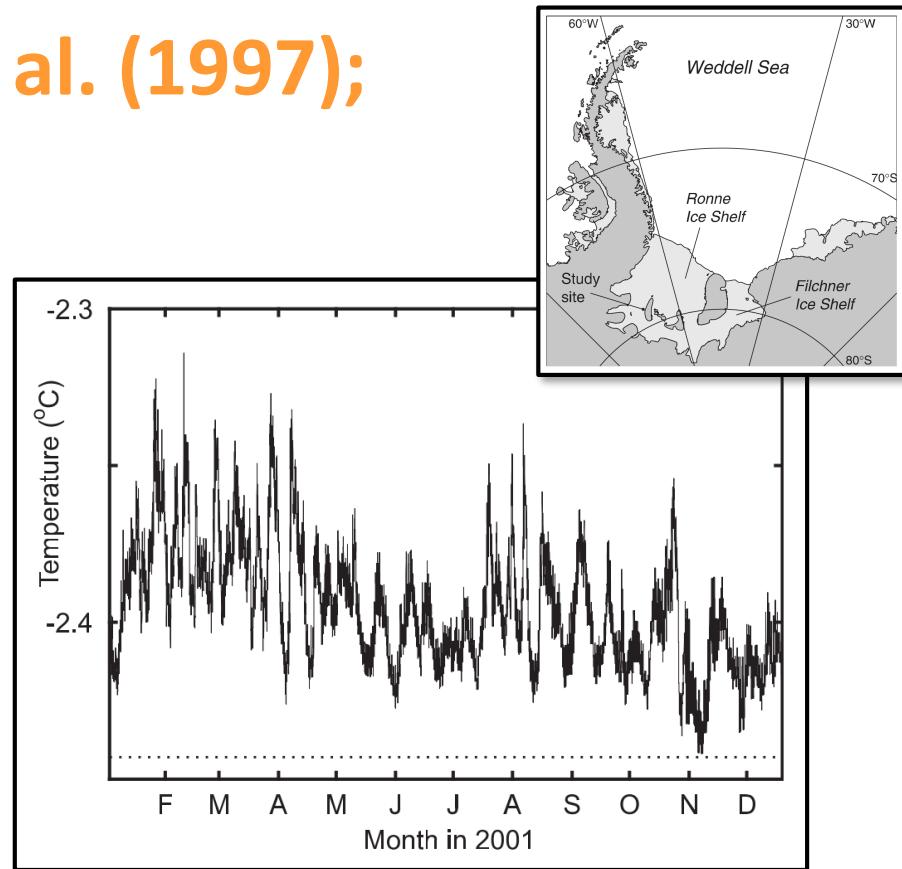
Nicholls et al. (1997). New oceanographic data from beneath Ronne Ice Shelf, Antarctica. GRL, 24(2), 167.
Corr et al. (2002). Precise measurement of changes in ice-shelf thickness by phase-sensitive radar to determine basal melt rates. GRL, 29(8), 73-1-74-4.

Figures from Jenkins et al. (2010)



Observations: Nicholls et al. (1997); Corr et al. (2002)

- “Site 3” on Ronne Ice Shelf
- Long CTD time series of T, S and $|u|$ at various depths



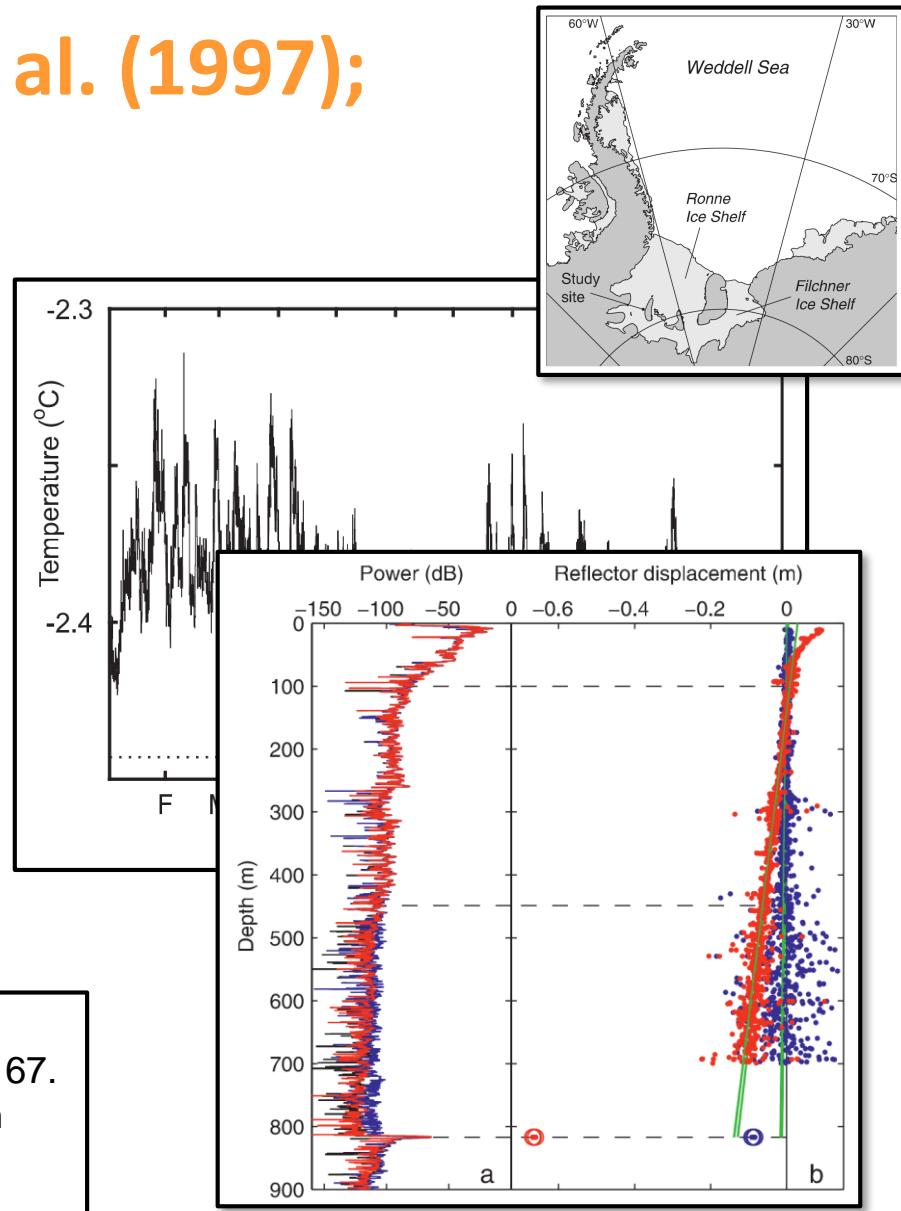
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Observations: Nicholls et al. (1997); Corr et al. (2002)

- “Site 3” on Ronne Ice Shelf
- Long CTD time series of T, S and $|u|$ at various depths
- Concurrent radar measurements of ice layers
 - Inferred mean melt rates

Nicholls et al. (1997). New oceanographic data from beneath Ronne Ice Shelf, Antarctica. GRL, 24(2), 167.
Corr et al. (2002). Precise measurement of changes in ice-shelf thickness by phase-sensitive radar to determine basal melt rates. GRL, 29(8), 73-1-74-4.



Figures from Jenkins et al. (2010)



Model fit: Jenkins et al. (2010)

- Found Best-fit parameters for Site 3

Symbol	Value	Description
$C_d^{1/2}\Gamma_T$	0.0011	Thermal Stanton number
$C_d^{1/2}\Gamma_S$	3.1×10^{-5}	Diffusion Stanton number
$C_d^{1/2}\Gamma_{\{TS\}}$	5.9×10^{-4}	Stanton number
C_d	0.0097	Drag coefficient
Γ_T	0.011	Heat transfer coefficient
Γ_S	3.1×10^{-4}	Salt transfer coefficient
$\Gamma_{\{TS\}}$	0.006	Transfer coefficient

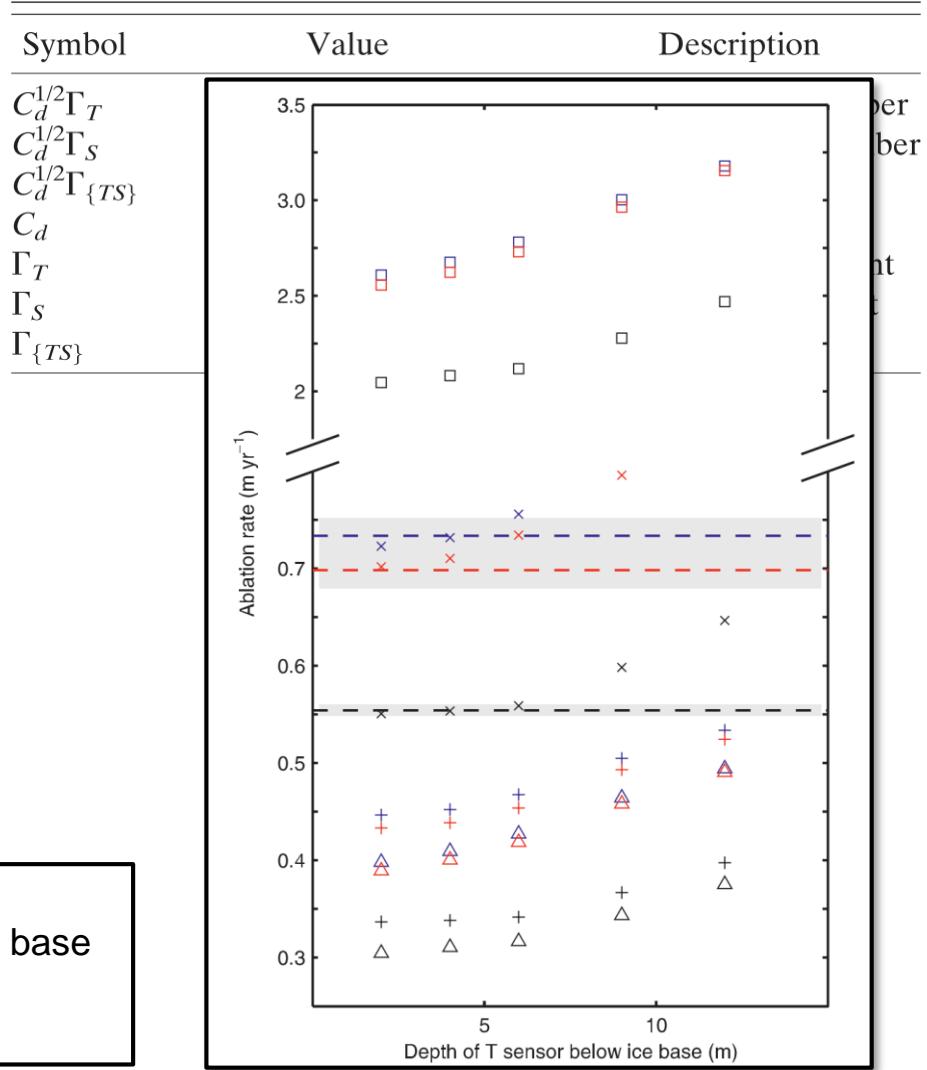
Jenkins, A., Nicholls, K. W., & Corr, H. F. J. (2010). Observation and parameterization of ablation at the base of Ronne Ice Shelf, Antarctica. *Journal of Physical Oceanography*, 40(10), 2298–2312.



Model fit: Jenkins et al. (2010)

- Found Best-fit parameters for Site 3
- Showed McPhee et al. (1987) consistent with Ronne Site 3 observations

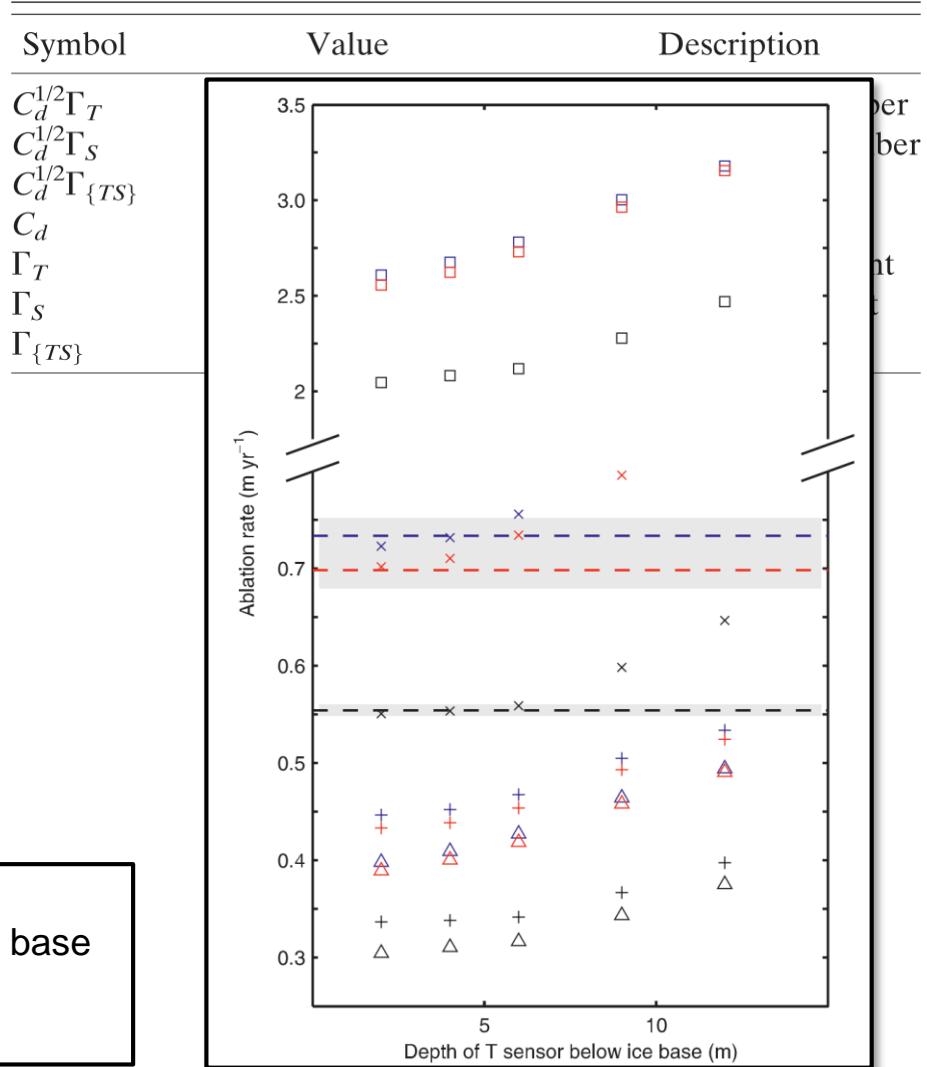
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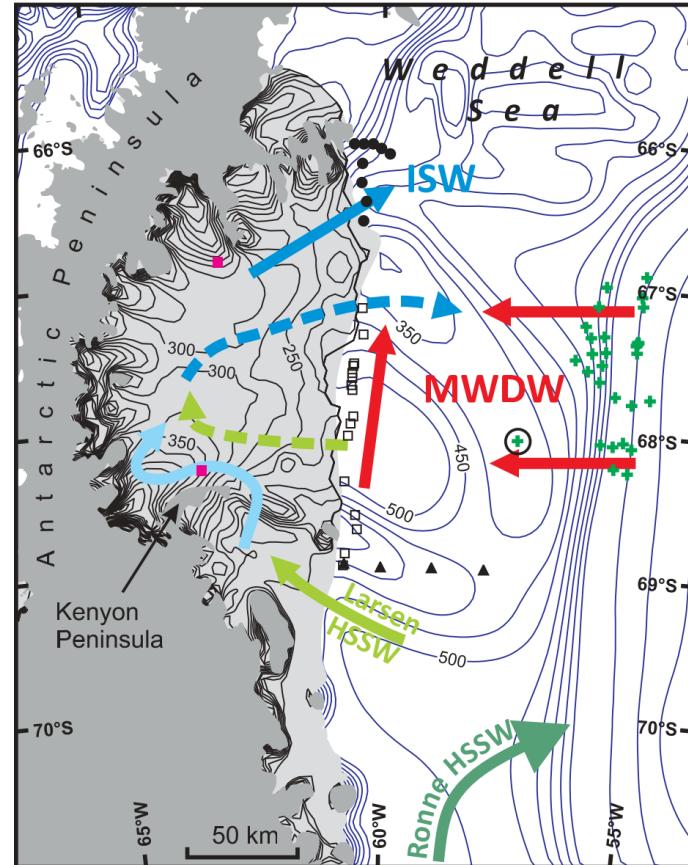
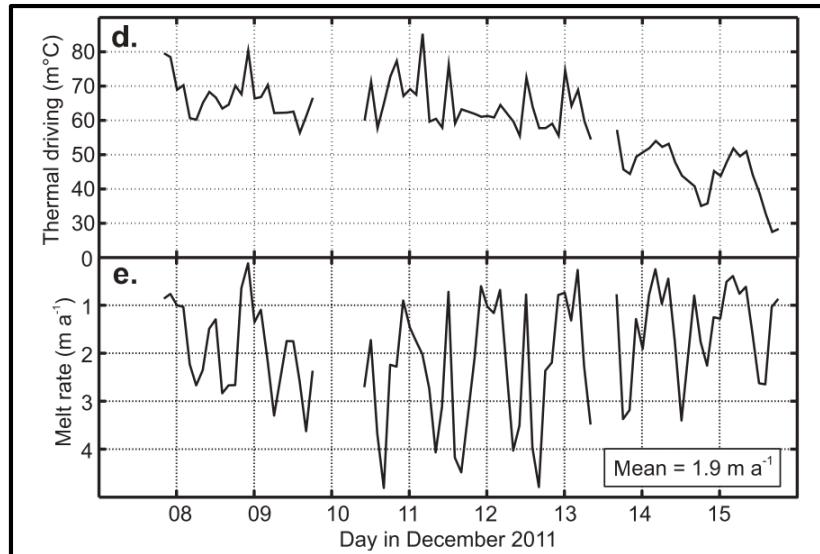
- Found Best-fit parameters for Site 3
- Showed McPhee et al. (1987) consistent with Ronne Site 3 observations
- Indicate that best-fit Γ_T , Γ_s and C_d vary with z_w
 Γ_T , Γ_s - transfer coefficients
 C_d - drag coefficient
 z_w - depth of CTD or grid cell

Jenkins, A., Nicholls, K. W., & Corr, H. F. J. (2010). Observation and parameterization of ablation at the base of Ronne Ice Shelf, Antarctica. *Journal of Physical Oceanography*, 40(10), 2298–2312.



Observations: Nicholls et al. (2012)

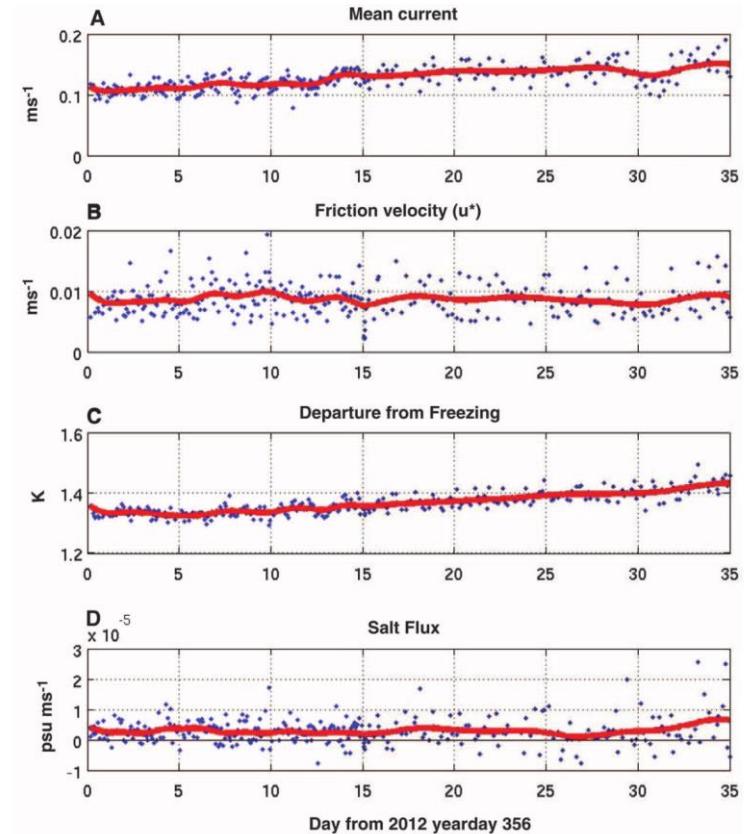
- Larsen C
- CTD, ADCP and pRes
- Best fit like Jenkins et al. (2010)



Nicholls et al. (2012). Ocean circulation beneath Larsen C
Ice Shelf, Antarctica from in situ observations.
Geophysical Research Letters, 39(19), L19608.

Observations: Stanton et al. (2013)

- We heard al about it yesterday
- PIG
- CTD
- 4Hz flux package (u , v , w , T , p , S)
 - Turb. heat, salt and mom. fluxes
- Acoustic backscatter (melt measurements)
- pRes (more melt measurements)
- Comparison with parameterizations in progress



Stanton et al. (2013). Channelized ice melting in the ocean boundary layer beneath Pine Island Glacier, Antarctica. Science (New York, N.Y.), 341(6151), 1236–9.

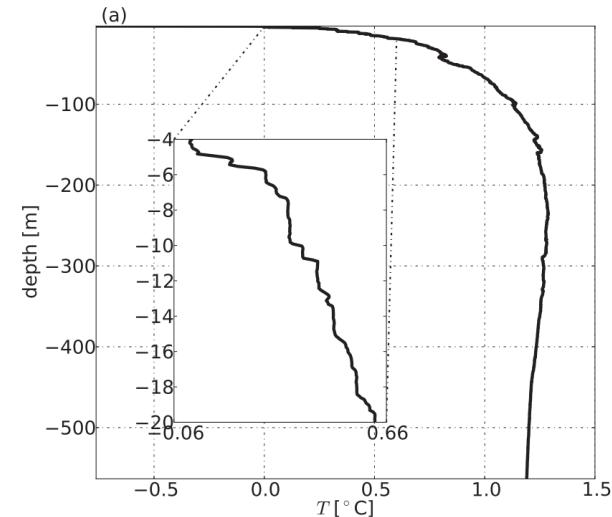


Name, Research Domain

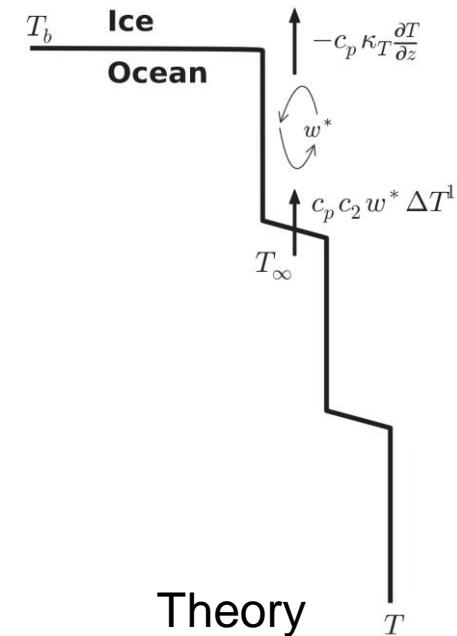
Theory/Obs: Kimura et al. (2015)

- Thermal staircase under George VI
- Outer layer isn't uniformly turbulent
- Instead, diffusive convection
 - Double diffusion
 - Freshening from melt is stabilizing
 - Cooling from melt is destabilizing
 - Alternating stable, unstable regions
- New parameterizations likely needed

Kimura, S., Nicholls, K. W., & Venables, E. (2015). Estimation of Ice Shelf Melt Rate in the Presence of a Thermohaline Staircase. *Journal of Physical Oceanography*, 45(1), 133–148.



Observations



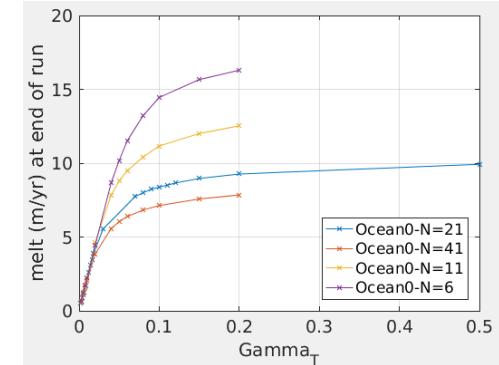
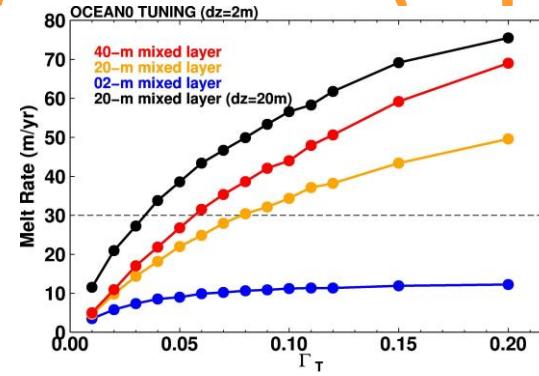
Theory



Model Intercomparison: Gwyther et al. (in prep.)

- Simulations using the ISOMIP+ framework (Asay-Davis et al. 2016)
- ROMS and COCO melt rates do not converge with vert. resolution
- MPAS-O, POP2x and NEMO melt rates converge only if BL thickness held fixed (arbitrary length scale)

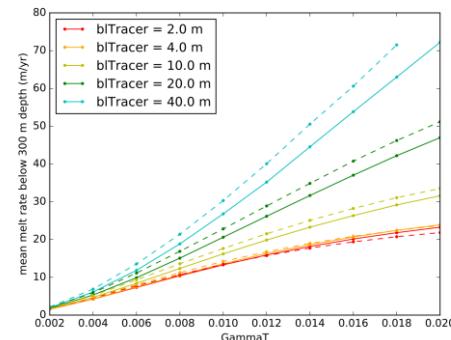
Gwyther et al. (in prep.) The importance of vertical processes and resolution on basal melt of an ice shelf: a multi-model study within the ISOMIP+ framework.
Asay-Davis et al. (2016). Experimental design for three interrelated marine ice sheet and ocean model intercomparison projects: MISMIP v. 3 (MISMIP+), ISOMIP v. 2 (ISOMIP+) and MISOMIP v. 1 (MISOMIP1). Geoscientific Model Development, 9(7), 2471–2497.



COCO

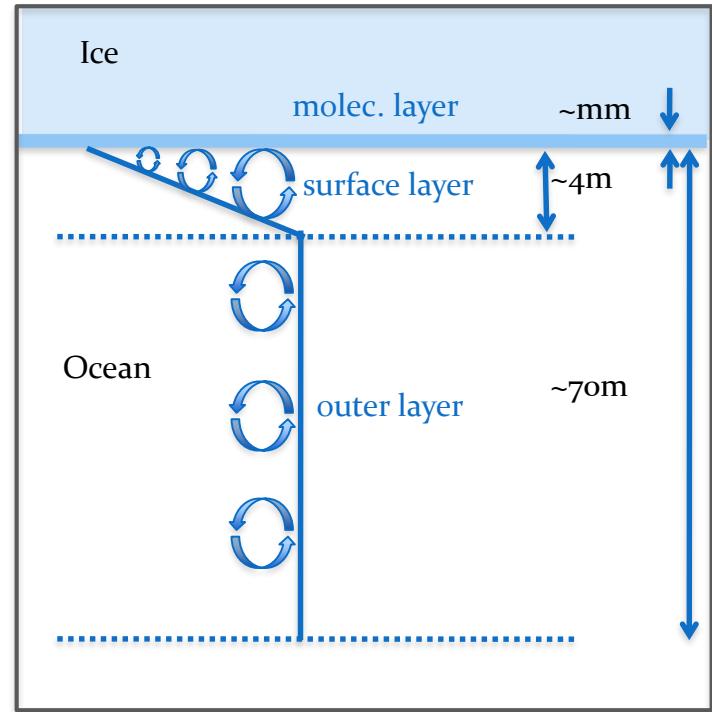
ROMS

MPAS-O



Need better BL parameterizations: convergence

- Ocean models sample T_w , S_w and u_w at depth z_w below ice-ocean interface



$$\rho_i a_b L_i = \rho_i c_i \kappa_i \left. \frac{\partial T_i}{\partial z} \right|_b - \rho_w c_w u_* \Gamma_T [T_f(S_b, P_b) - T_w],$$

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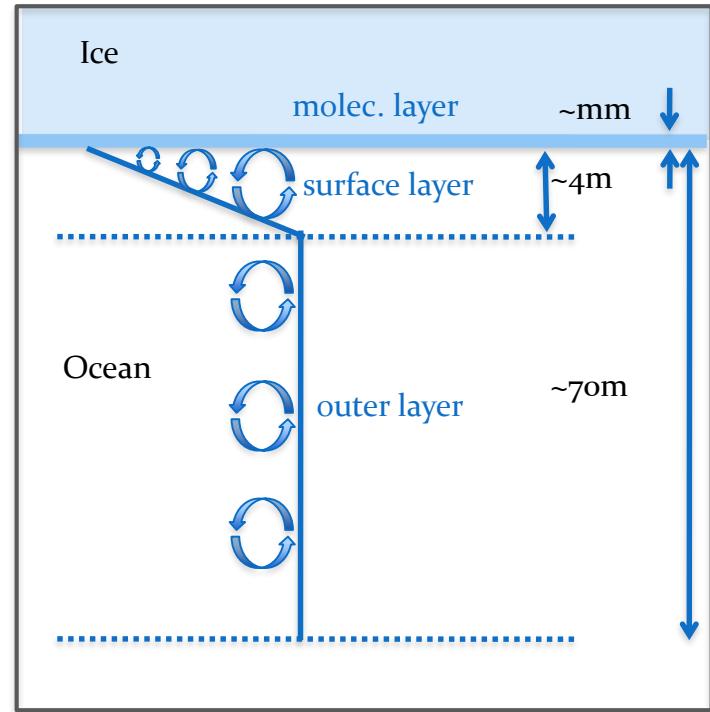
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“Three equations” using notation from
Jenkins et al. (2010)



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- To converge with vertical resolution, ocean models should:
 - Have consistent diffusion coeffs. $K_h(z)$ and $K_s(z)$ in resolved and parameterized BL



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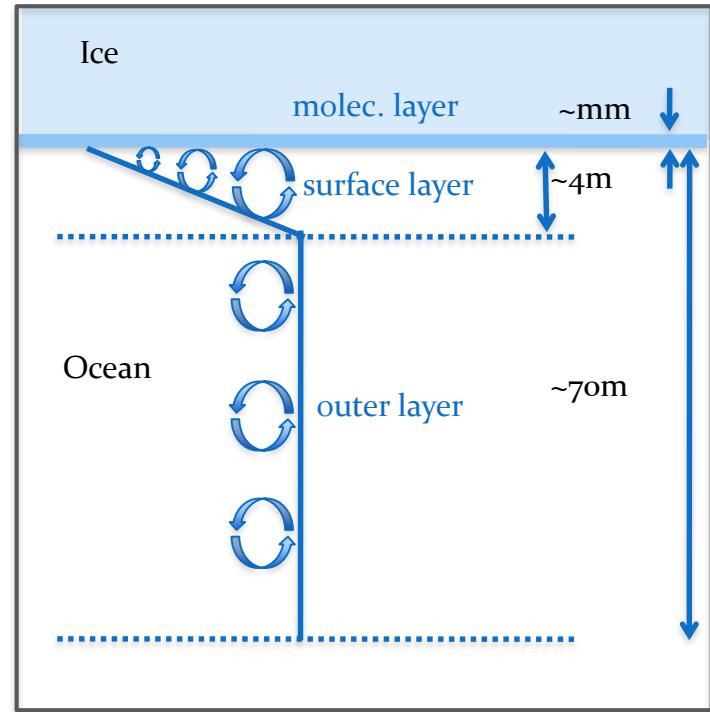
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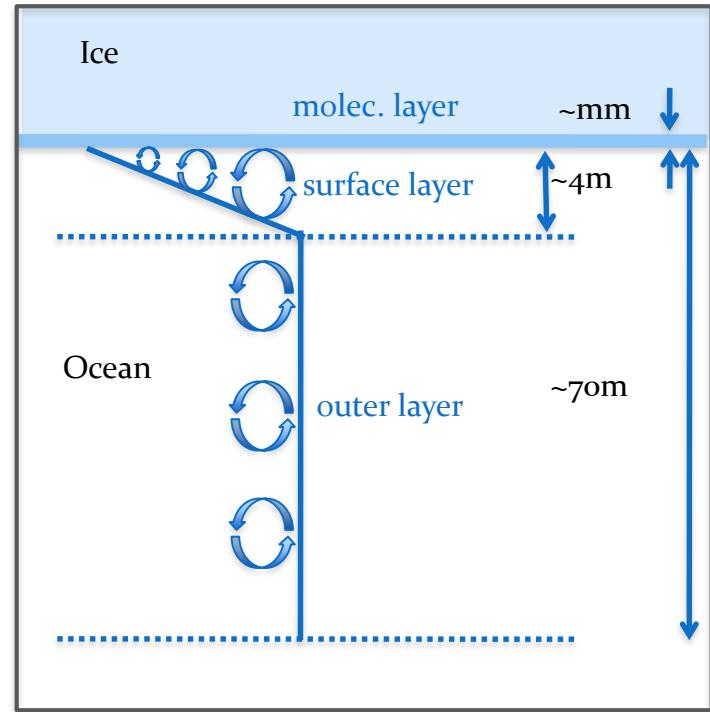
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 - Use transfer coeffs. $\Gamma_T(z_w)$ and $\Gamma_S(z_w)$ that depend on $K_h(z)$ and $K_s(z)$
 - Use a z-dependent drag coeff: $u_*^2 = C_D(z_w) |u_w|^2$



$$\rho_i a_b L_i = \rho_i c_i \kappa_i \left. \frac{\partial T_i}{\partial z} \right|_b - \rho_w c_w u_* \Gamma_T [T_f(S_b, P_b) - T_w],$$

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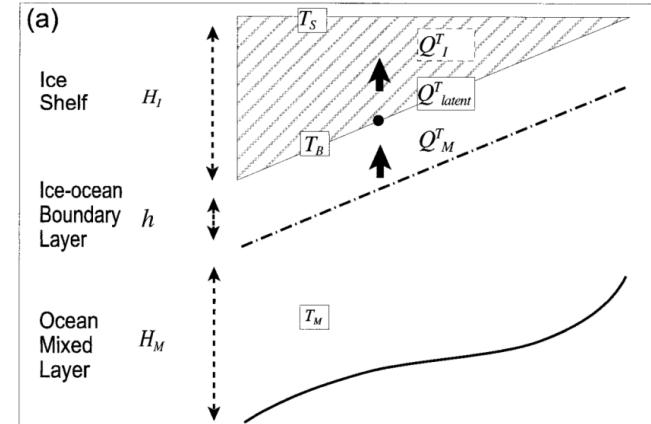
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Need better BL parameterizations: buoyancy

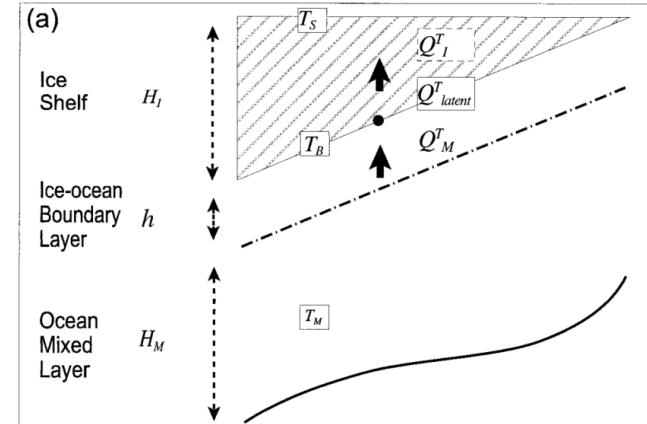
- In current BL params.
(McPhee et al. 1987):
 - Buoyancy reduces eddies via “stability parameter” η_*
(less mixing in BL)
 - But no buoyancy-driven flow in BL



$$\eta_* = \left(1 + \frac{\xi_N u_*}{f L_O R_c} \right)^{-1/2},$$

Need better BL parameterizations: buoyancy

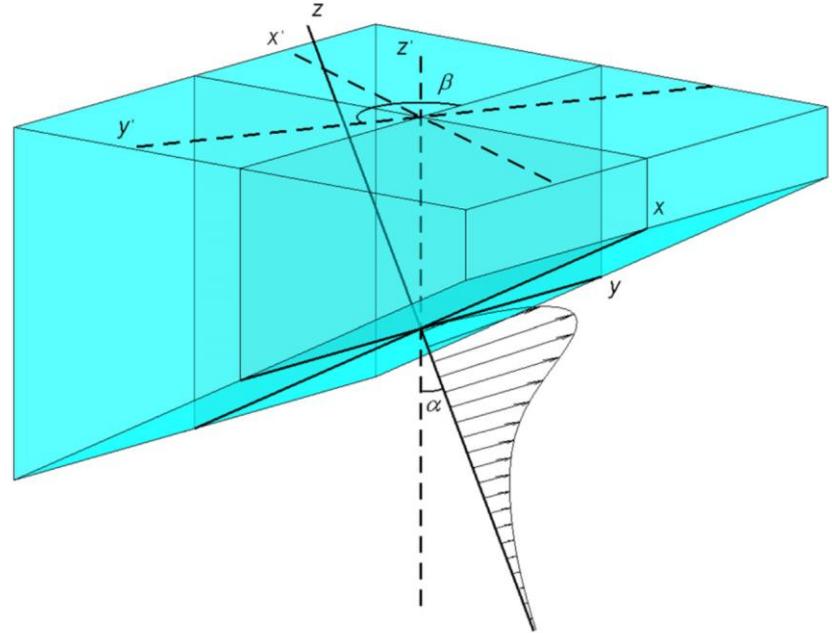
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(McPhee et al. 1987):
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(less mixing in BL)
 - But no buoyancy-driven flow in BL
- BL parameterization should include buoyancy-driven boundary current



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Theory: Jenkins (2016)

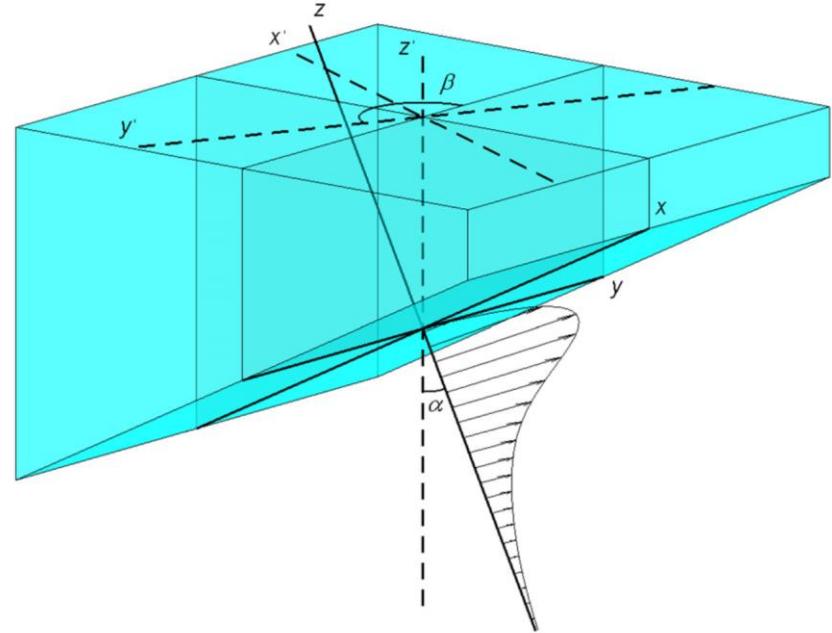
- 1D model but not just vert. diffusion



Jenkins, A. (2016). A Simple Model of the Ice Shelf–Ocean Boundary Layer and Current. *Journal of Physical Oceanography*, 46(6), 1785–1803.

Theory: Jenkins (2016)

- 1D model but not just vert. diffusion
- Allows along- and across-slope:
 - Gradients (in T, S, u, etc.)
 - Advection

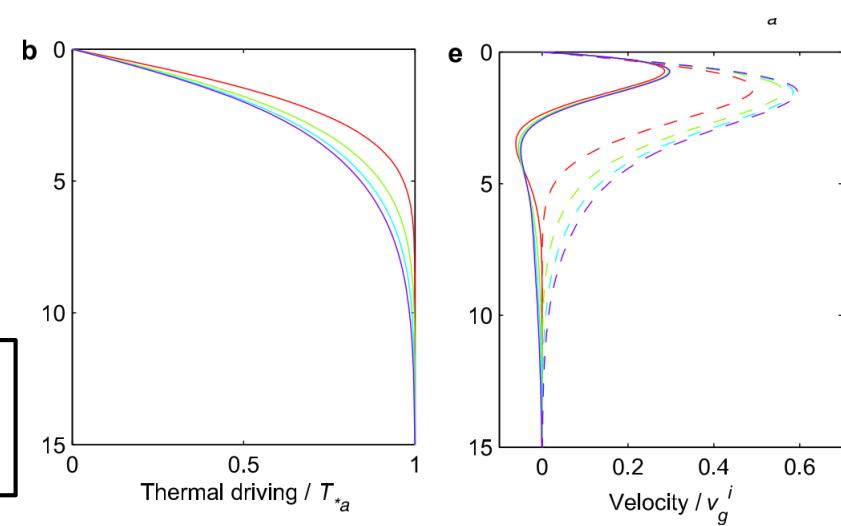
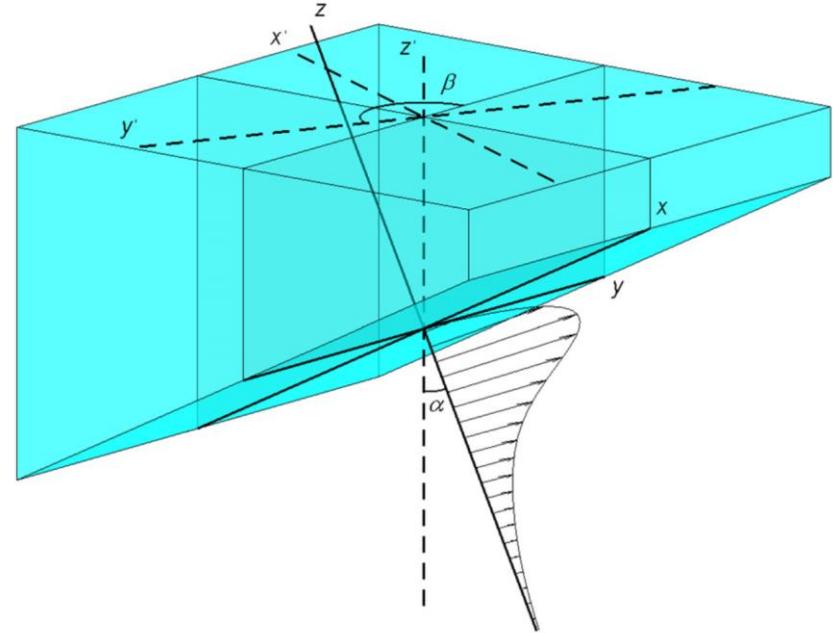


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 - Advection
- Boundary current
 - Inner Ekman boundary layer
 - Outer geostrophic region

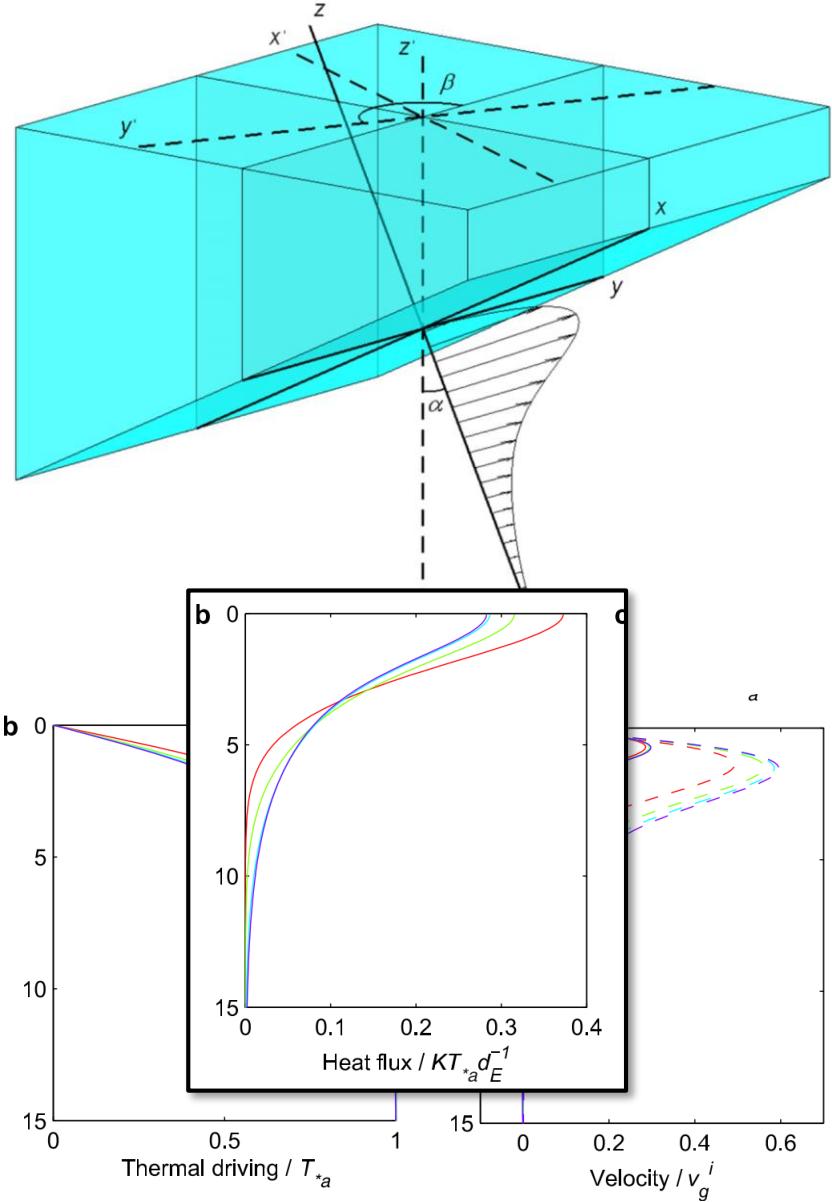


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 - Outer geostrophic region
- Heat flux not const. in BL

Jenkins, A. (2016). A Simple Model of the Ice Shelf–Ocean Boundary Layer and Current. *Journal of Physical Oceanography*, 46(6), 1785–1803.



Promising directions

Theory and Modeling

- Jenkins (2016), but needs theory for non-const. $K_h(z)$
- “Modelling the ice-shelf ocean boundary layer” (Supervised by John Taylor, Paul Holland and Keith Nicholls)
 - Very high resolution simulations
 - Modeling (not parameterizing) boundary-layer processes



Promising directions

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 - Very high resolution simulations
 - Modeling (not parameterizing) boundary-layer processes

Observations

- Tim Stanton’s observations below PIG
- Planned boreholes under FRIS, Ross, Amery... if properly instrumented
- As an end member (e.g. zero slope), similar obs. under sea ice



Summary

- BL parameterization in sub-ice-shelf ocean models:
 - Derived for sea ice (no buoyant flow)
 - Not well constrained
 - Likely needs work



Summary

- BL parameterization in sub-ice-shelf ocean models:
 - Derived for sea ice (no buoyant flow)
 - Not well constrained
 - Likely needs work
- Observations needed to:
 - Differentiate possible parameterizations
 - Constrain parameters
 - Span a range of conditions
 - Warm, cold cavities
 - Melting, freezing (including frazil)
 - All 3 modes of melting



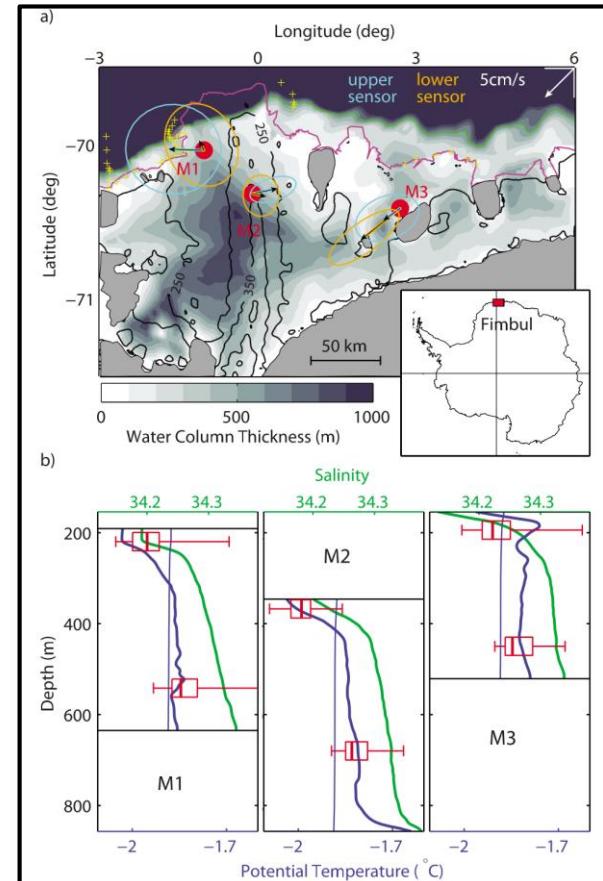
Thank you



Other related observational data sets

- Not yet used to (in)validate melt parameterizations
- Some do not include pRes (or other melt rate) measurements
- Hattermann et al. (2012)
 - Fimbul Ice Shelf
 - CTD (T, S and u)
 - no pRes (no independent melt rates)

Hattermann et al. (2012). Two years of oceanic observations below the Fimbul Ice Shelf, Antarctica. Geophysical Research Letters, 39(12), L12605.



Observations: Herraiz-Borreguero et al. (2013, 2015)

- ...among other papers
- Amery Ice Shelf
- CTD, no pRes

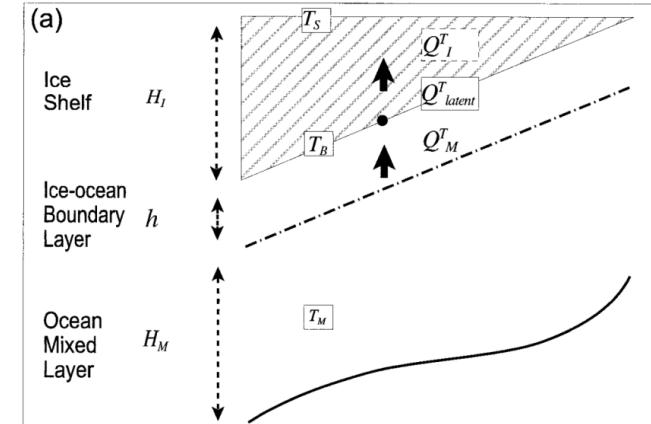
Herraiz-Borreguero et al. (2013). Ice shelf/ocean interactions under the Amery Ice Shelf: Seasonal variability and its effect on marine ice formation. *Journal of Geophysical Research: Oceans*, 118(12), 7117–7131.

Herraiz-Borreguero et al. (2015). Circulation of modified Circumpolar Deep Water and basal melt beneath the Amery Ice Shelf, East Antarctica. *Journal of Geophysical Research: Oceans*, 120(4), 3098–3112.

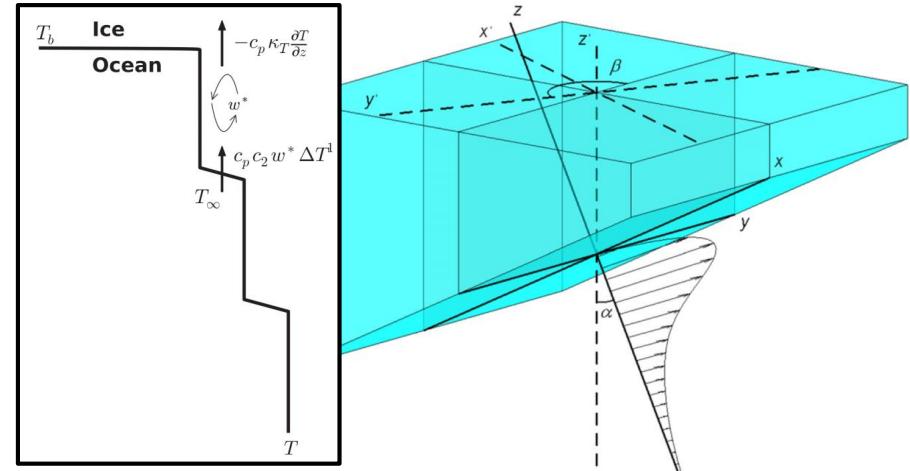


Need better BL parameterizations: buoyancy

- In current BL params. (McPhee et al. 1987):
 - Buoyancy reduces eddies via “stability parameter” η_* (less mixing in BL)
 - But no buoyancy-driven flow in BL
- BL parameterization should include buoyancy-driven boundary current (Jenkins 2016)
- Sometimes may need diffusive convection

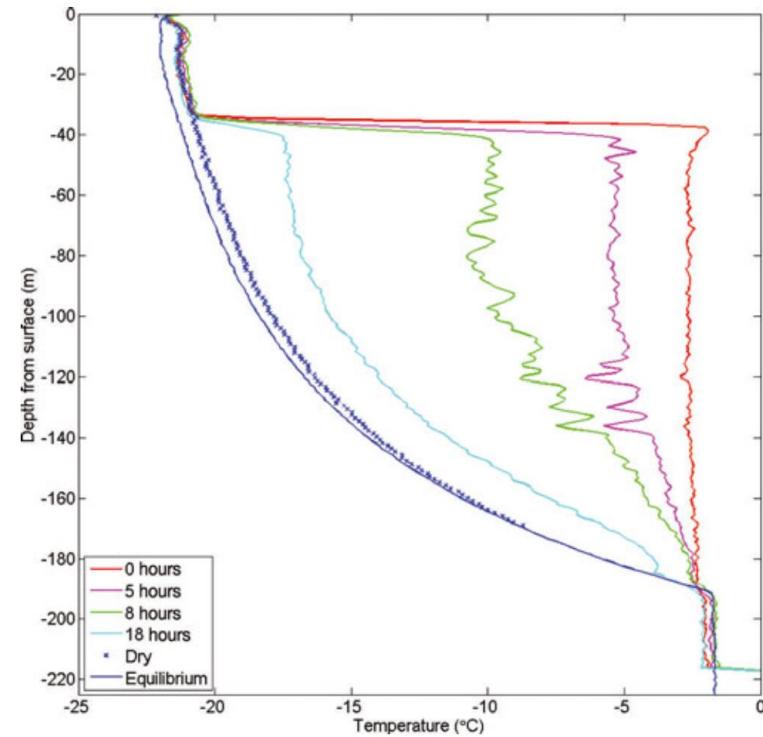
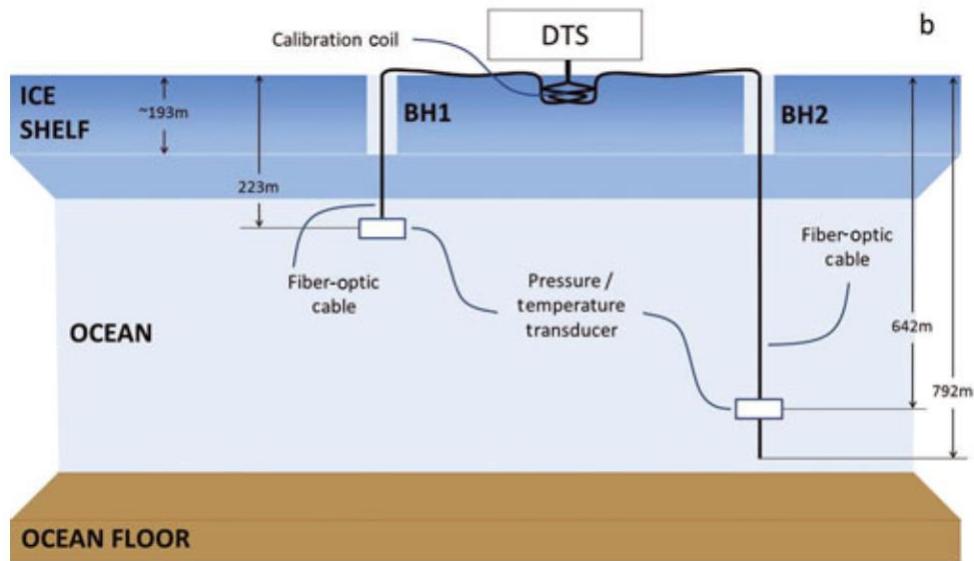


$$\eta_* = \left(1 + \frac{\xi_N u_*}{f L_O R_c} \right)^{-1/2},$$



Observations: Tyler et al. (2013)

- McMurdo Ice Shelf
- distributed temperature sensor (DTS)



Tyler et al. (2013). Using distributed temperature sensors to monitor an Antarctic ice shelf and sub-ice-shelf cavity. *Journal of Glaciology*, 59(215), 583–591.