



Seeing below the ice: what progress has been made since 2012?

Steve Rintoul

Antarctic Climate and Ecosystems Cooperative Research Centre

Centre for Southern Hemisphere Oceans Research

CSIRO Oceans & Atmosphere

Hobart, Tasmania , Australia

steve.rintoul@csiro.au

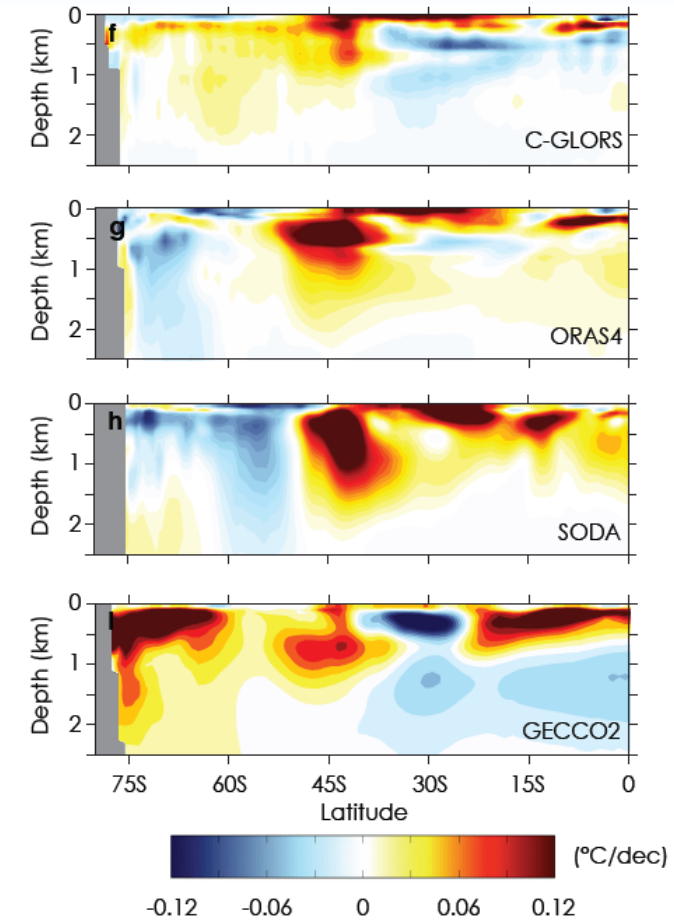
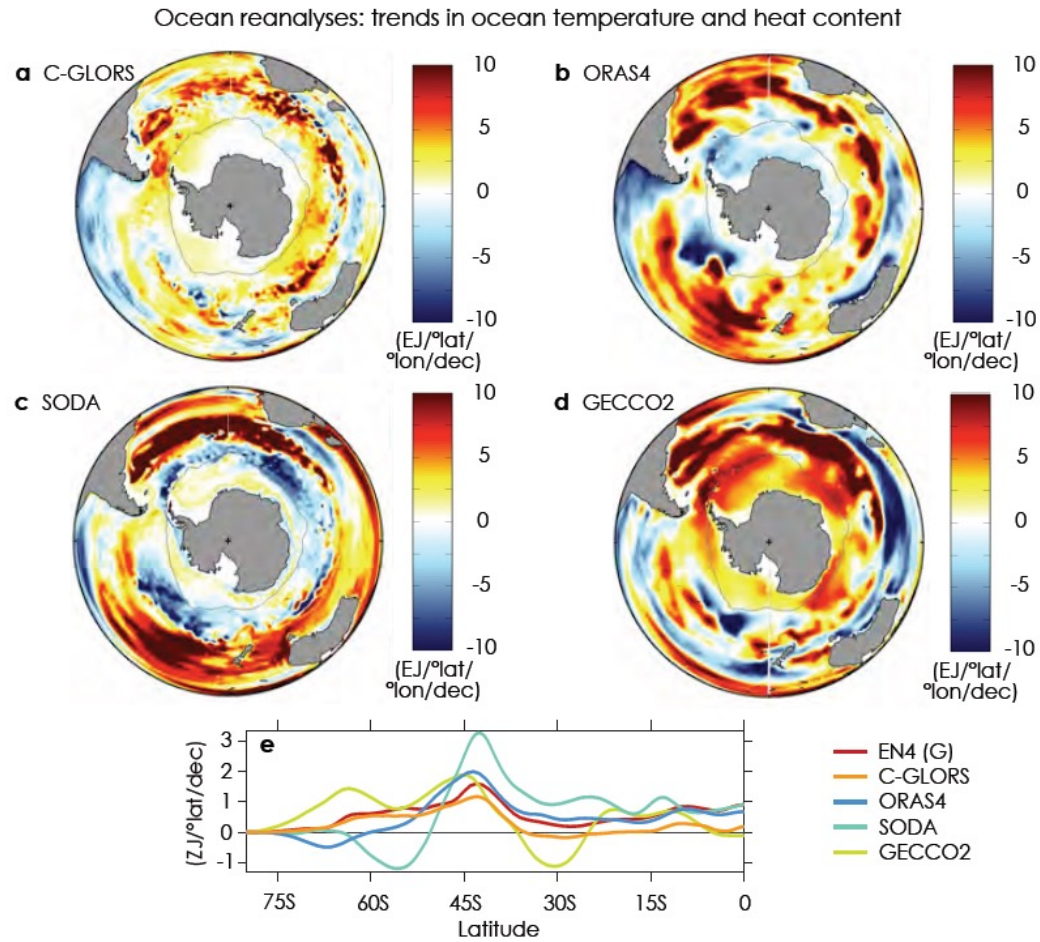


www.csiro.au

Theme 1: Circulation and heat/freshwater/carbon inventory

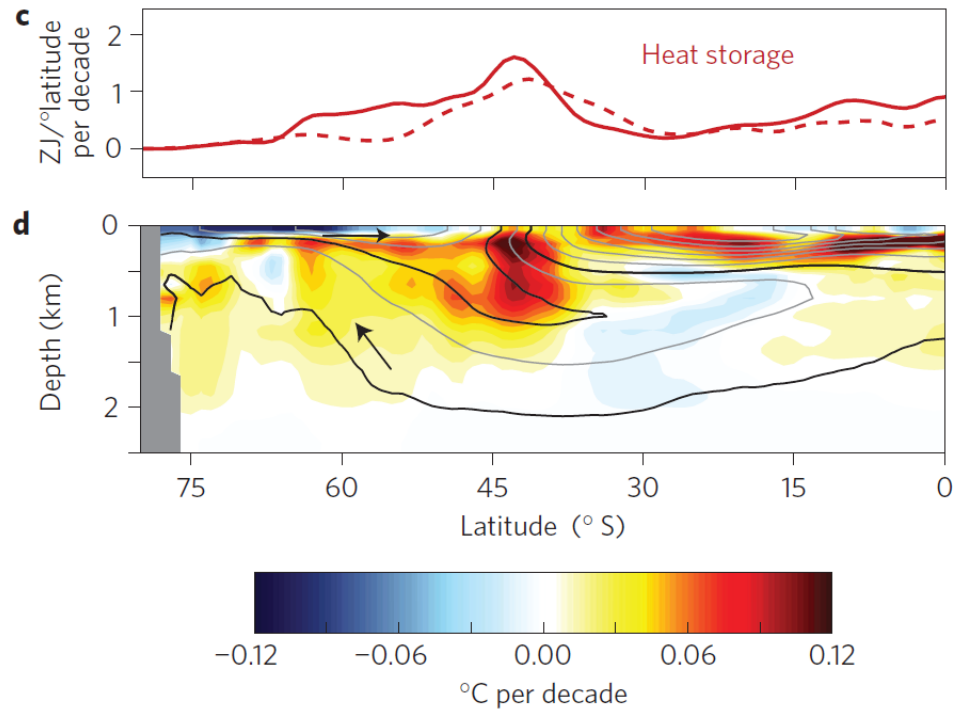


Southern Ocean heat content & temperature trends

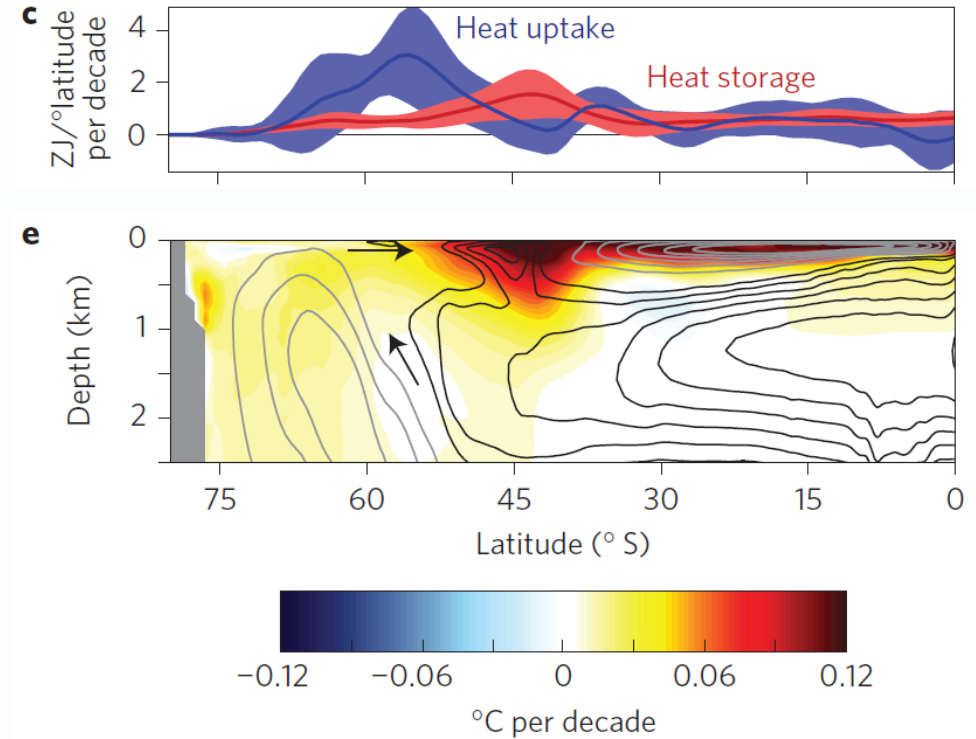


Heat uptake south of ACC, heat storage north of ACC

Observations

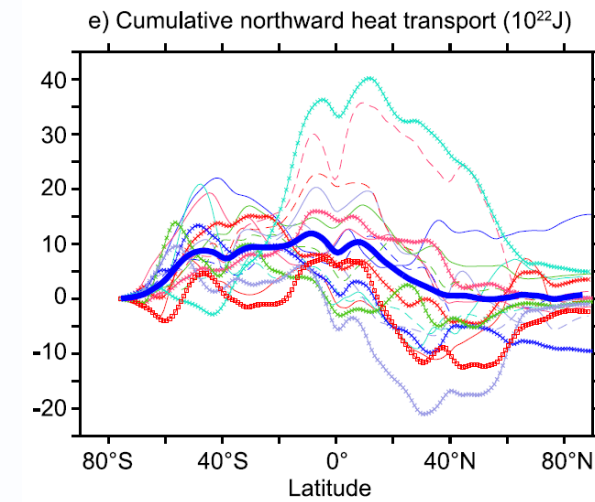
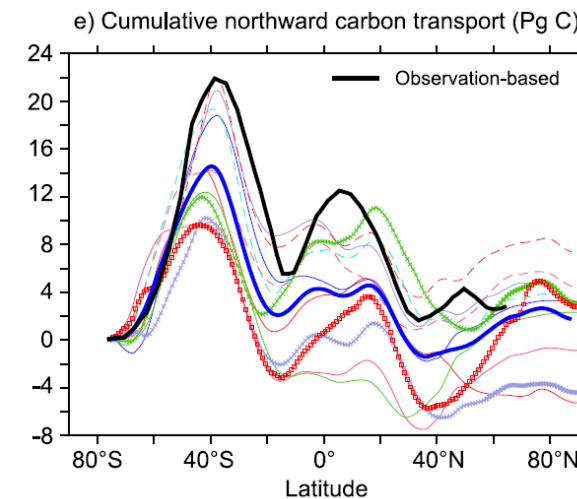
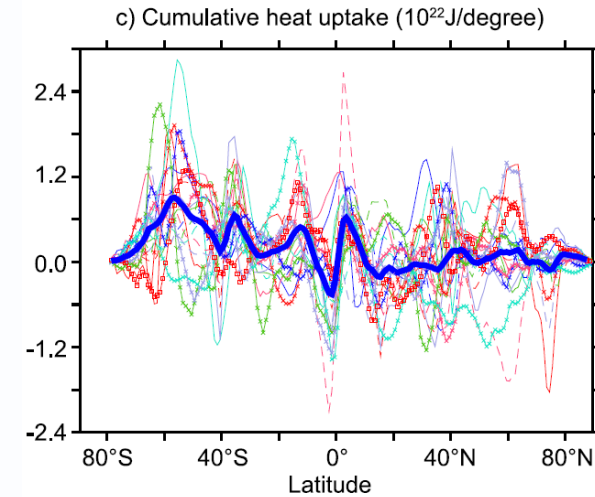
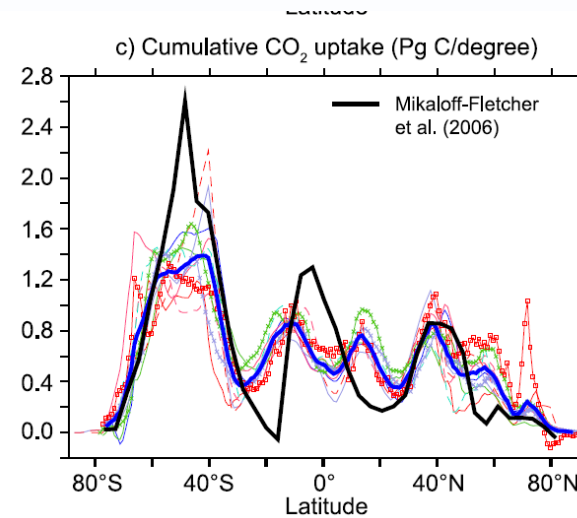


Models

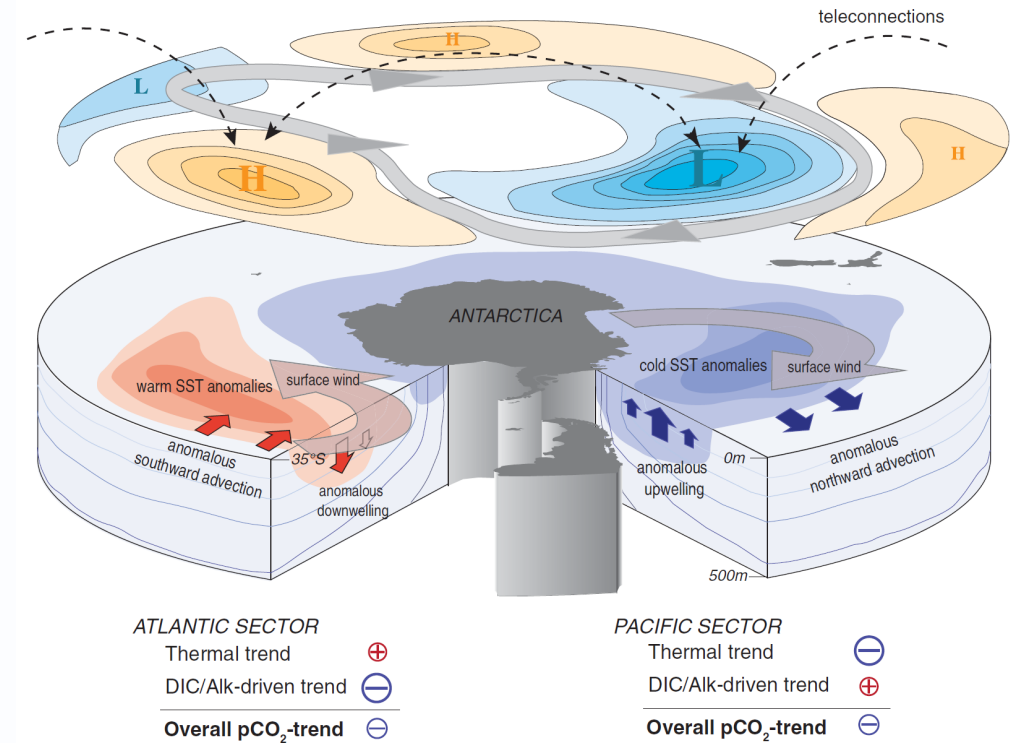
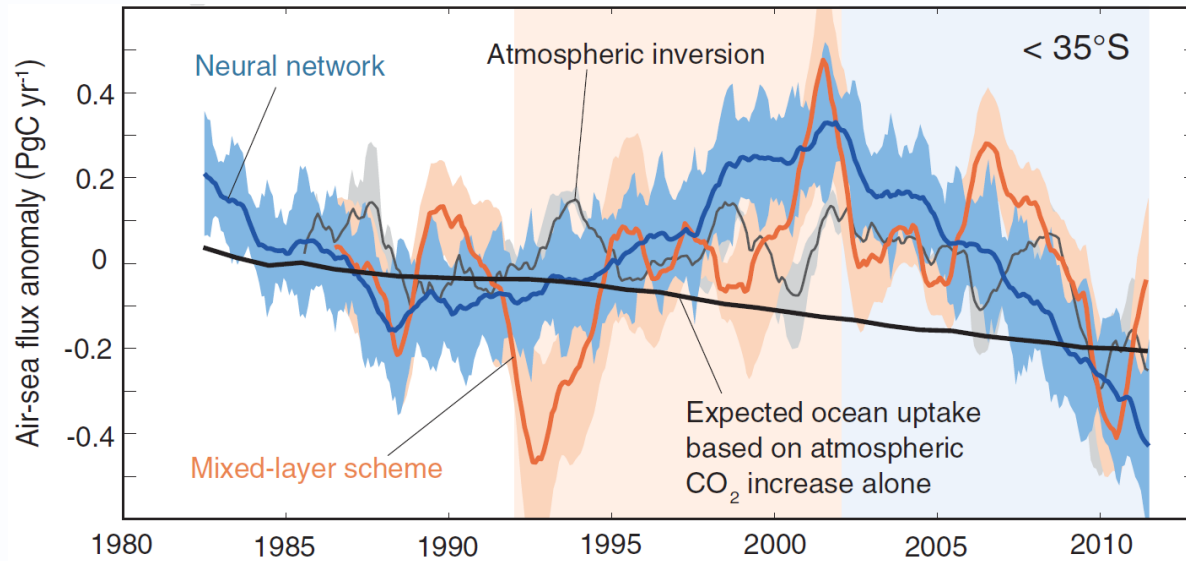


Southern Ocean dominance in ocean heat and carbon uptake

Southern Ocean accounts for 42% of anthropogenic carbon uptake, and 75% of anthropogenic heat uptake.

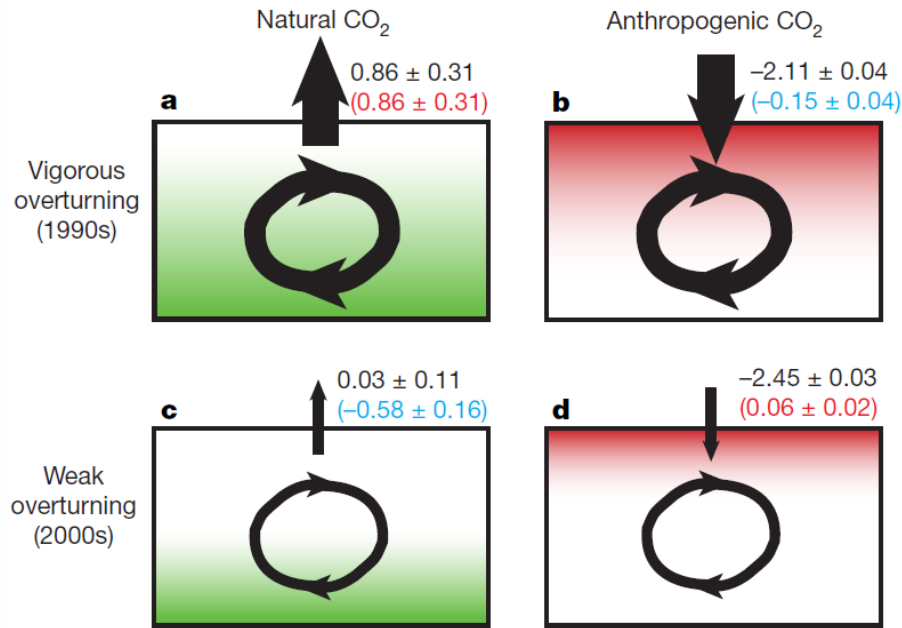


“Reinvigoration” of Southern Ocean carbon sink

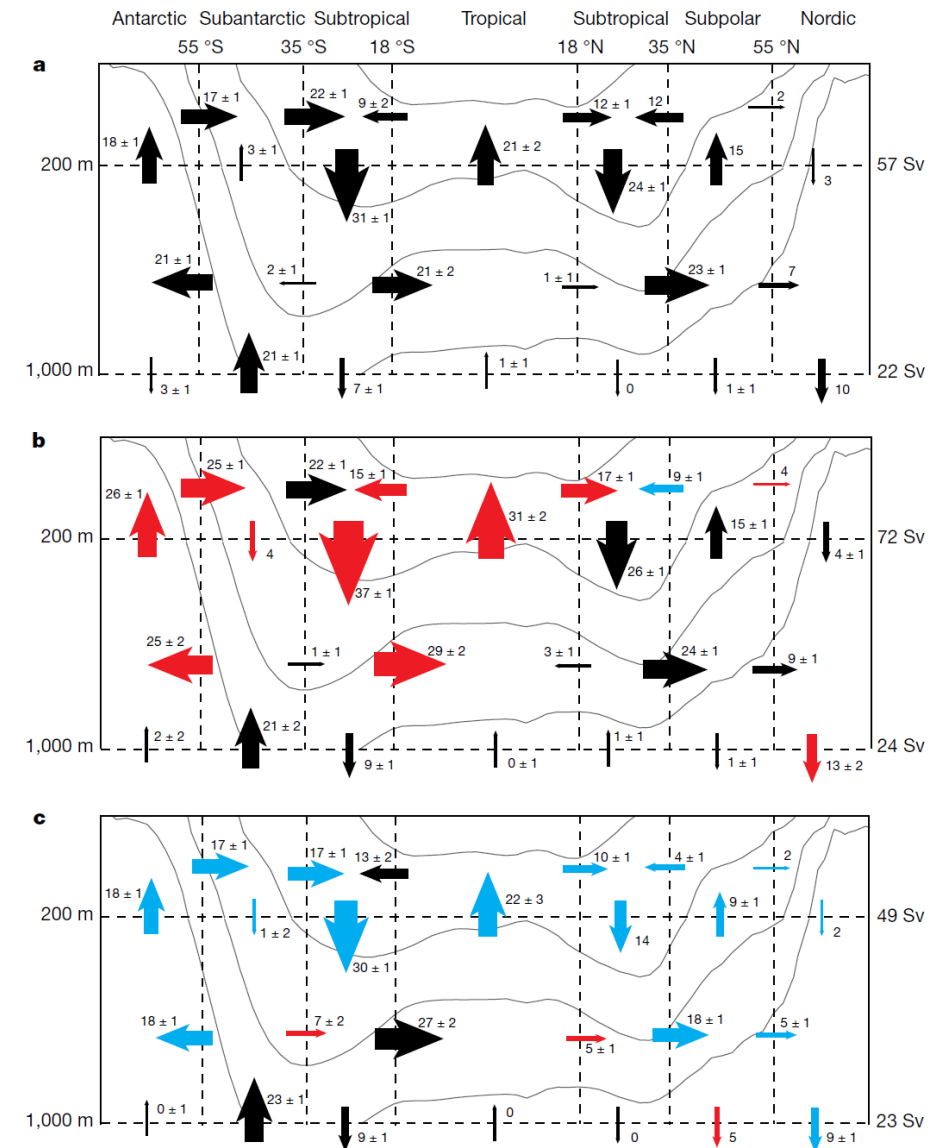


Landschutzer et al., 2016

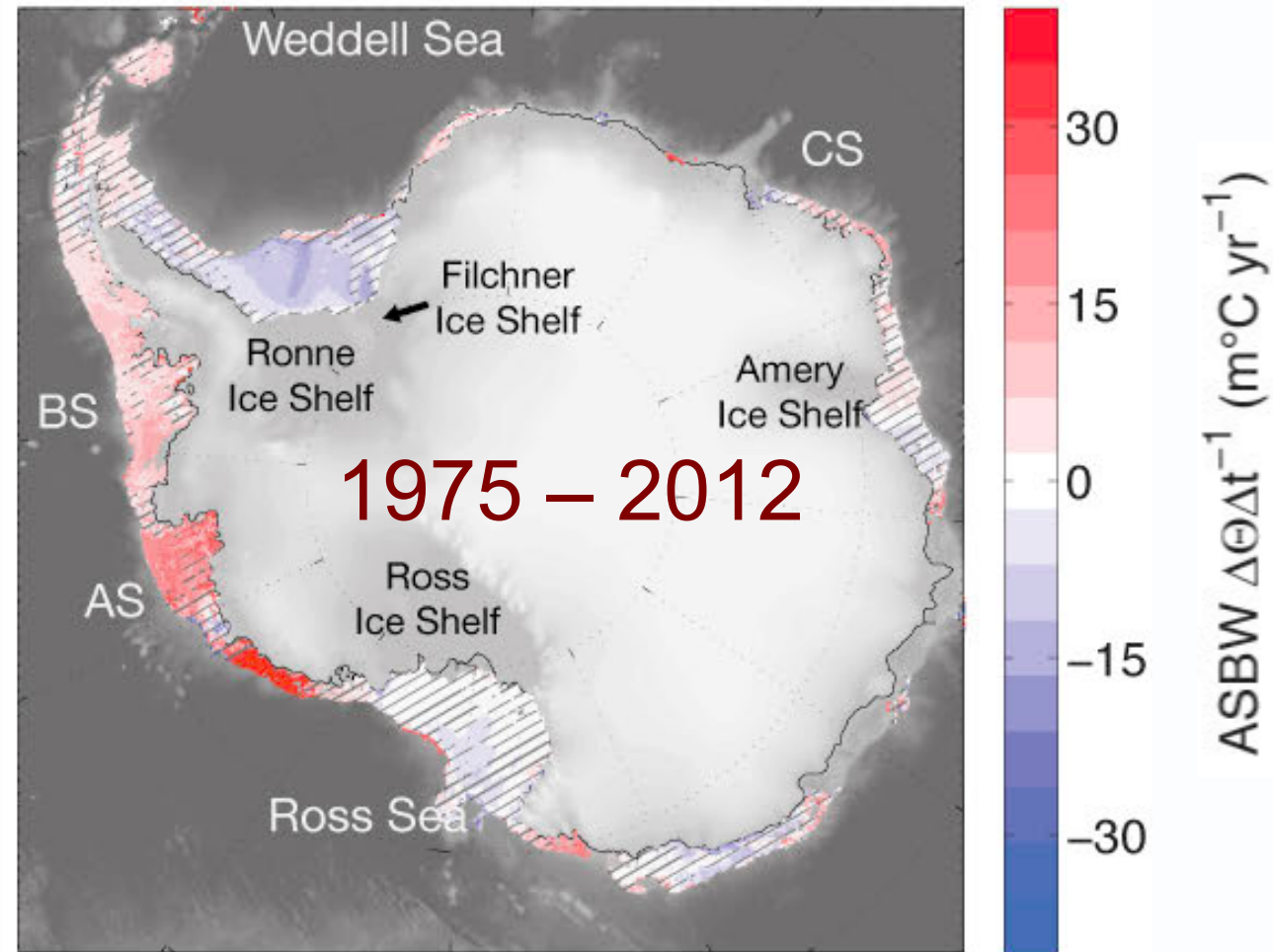
Weaker upper cell drives increase in ocean carbon storage in 2000s



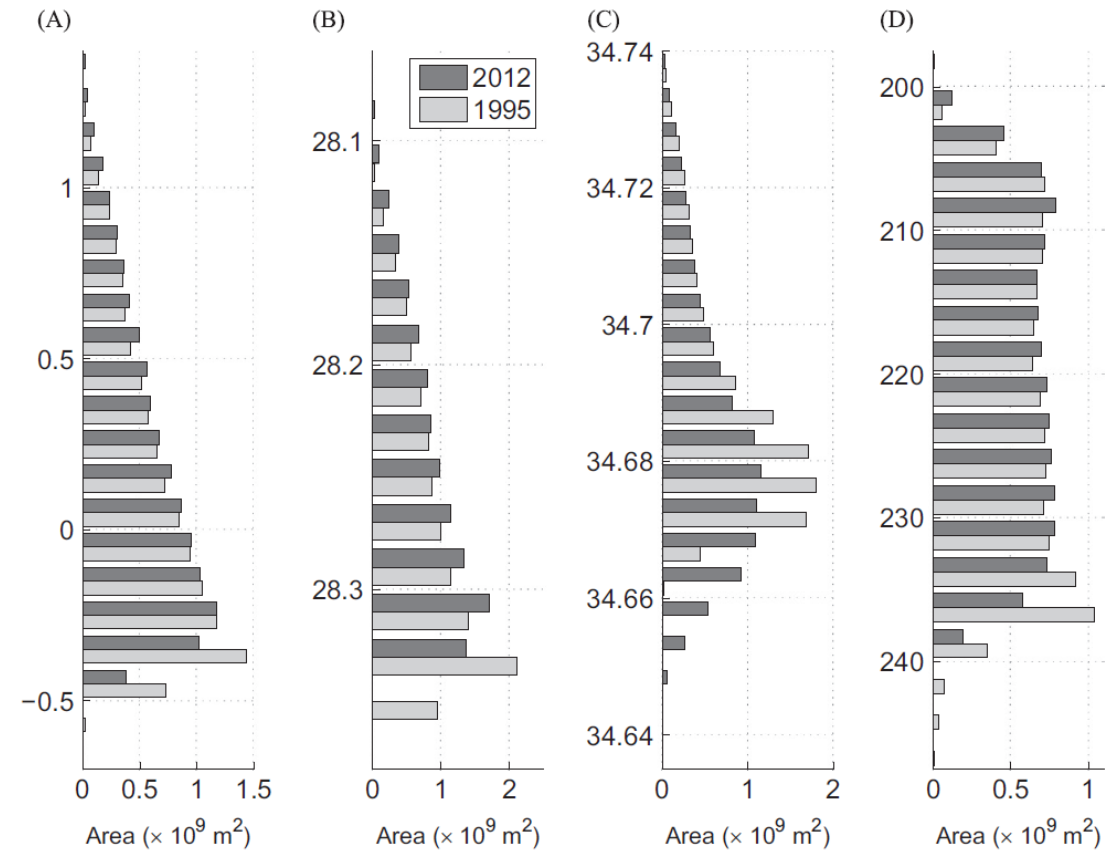
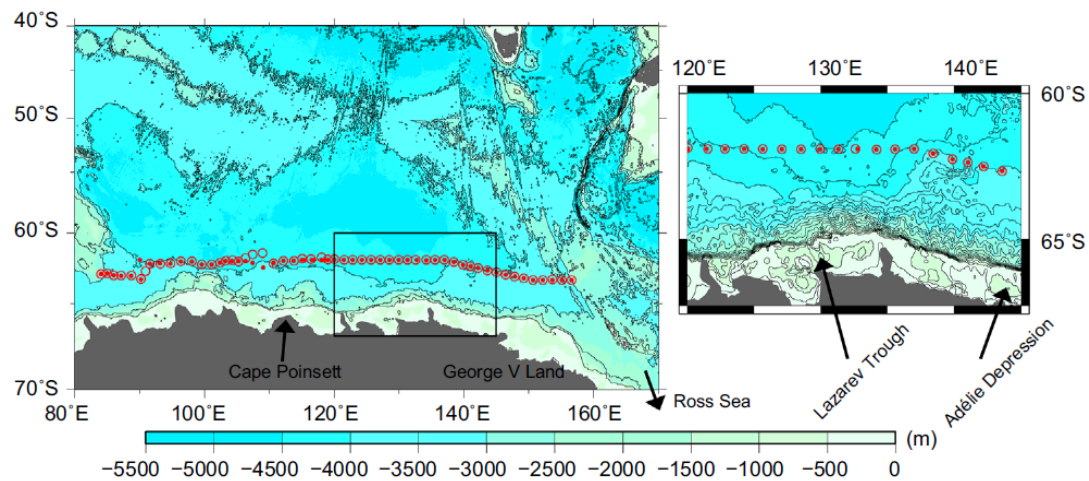
“To track the timing and magnitude of these effects, the variability of the ocean circulation and CO₂ uptake must be monitored carefully over time.”



Trend in Antarctic continental shelf bottom water

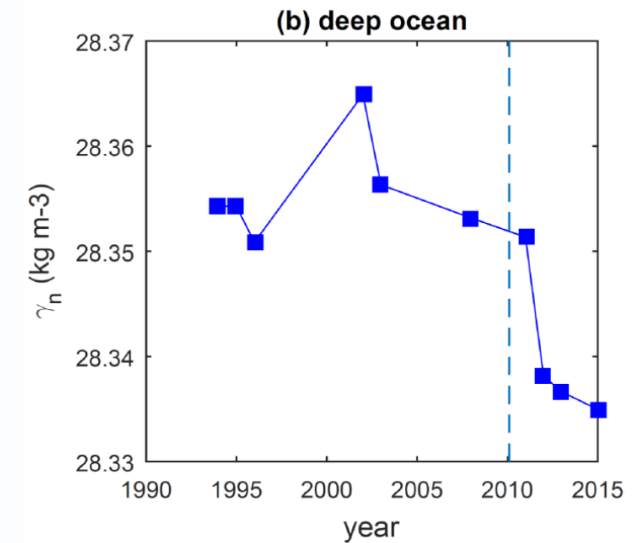
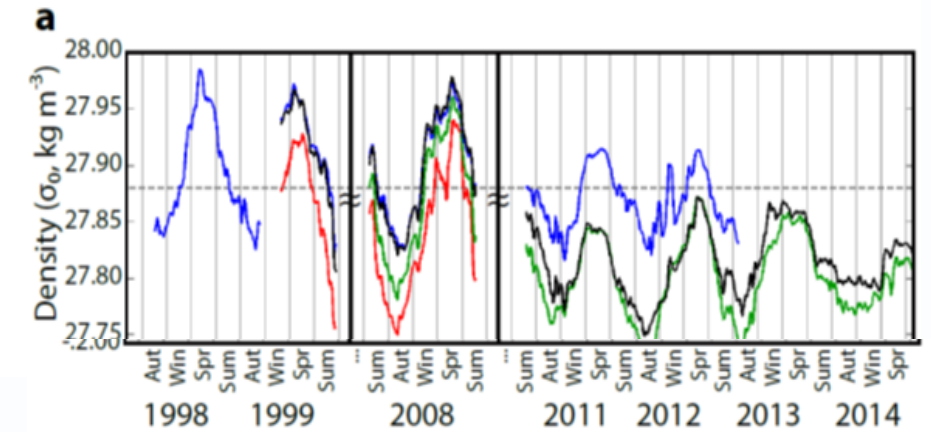
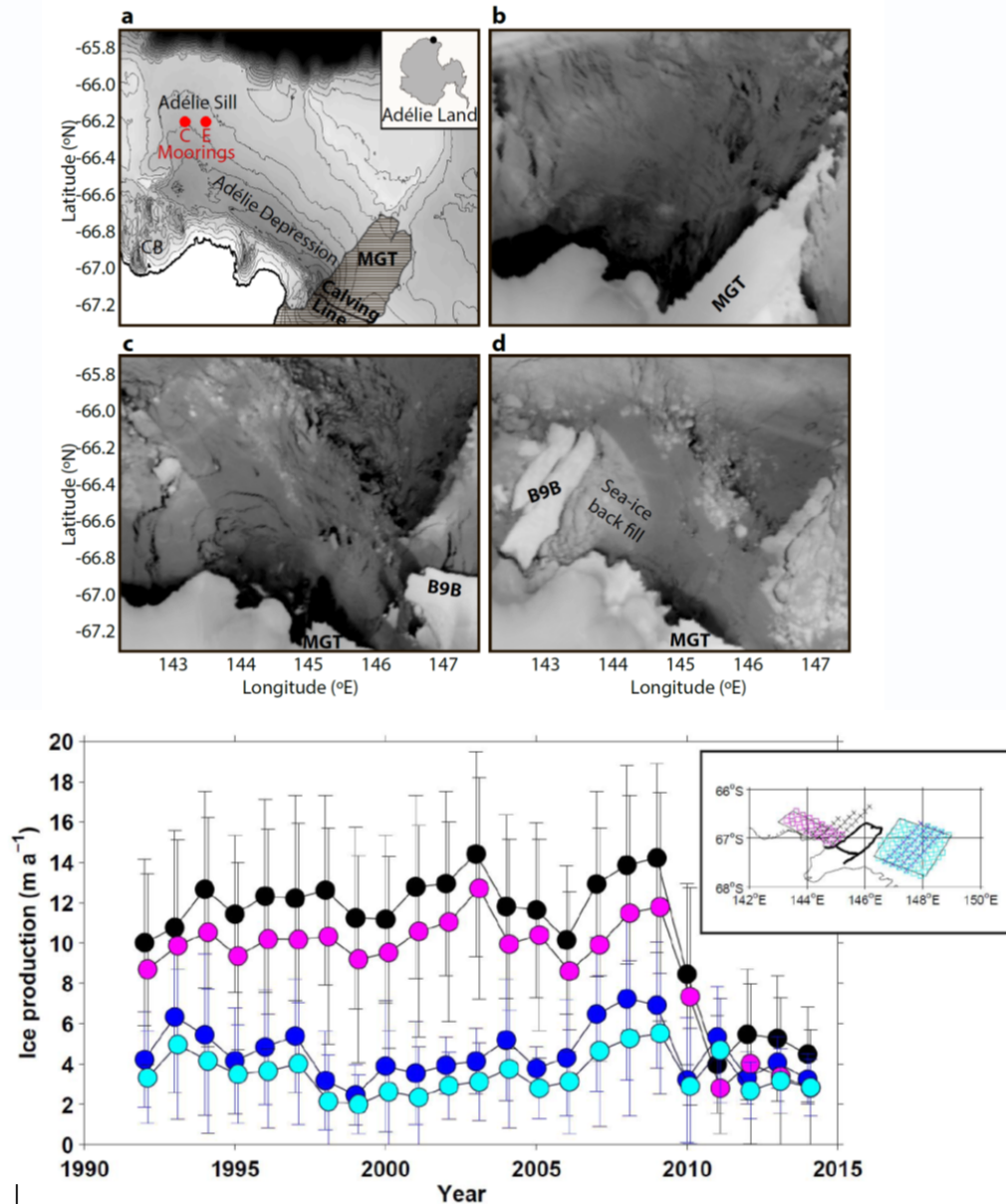


Warming, freshening and contraction of AABW



Katsumata et al., DSR-II, 2015

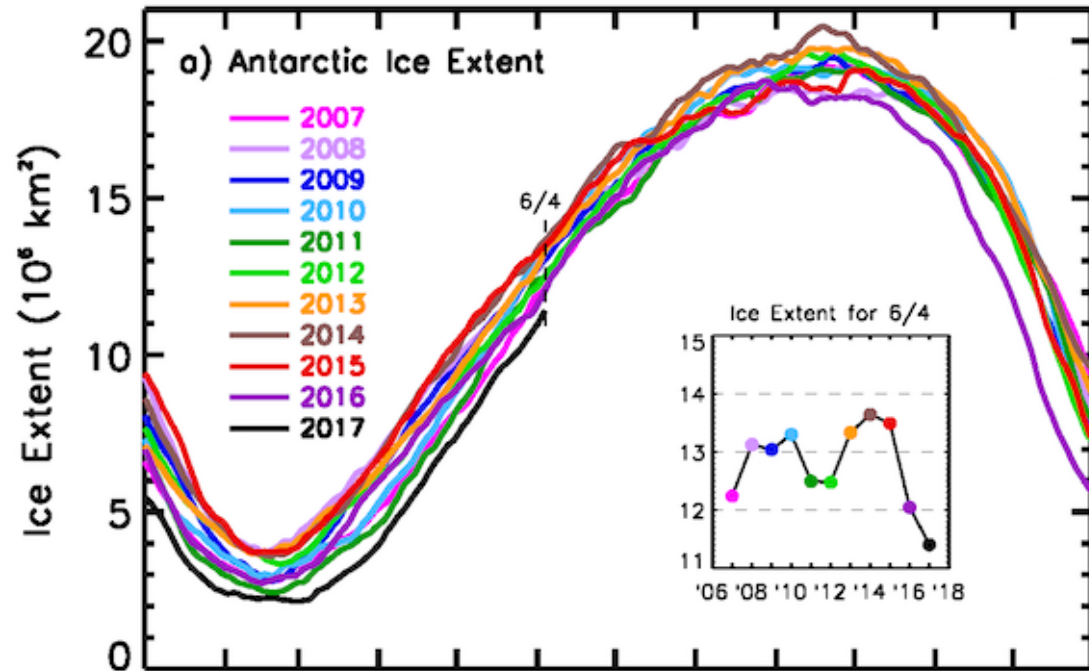
Sensitivity of bottom water formation to forcing



Theme 2: ocean – sea ice interaction

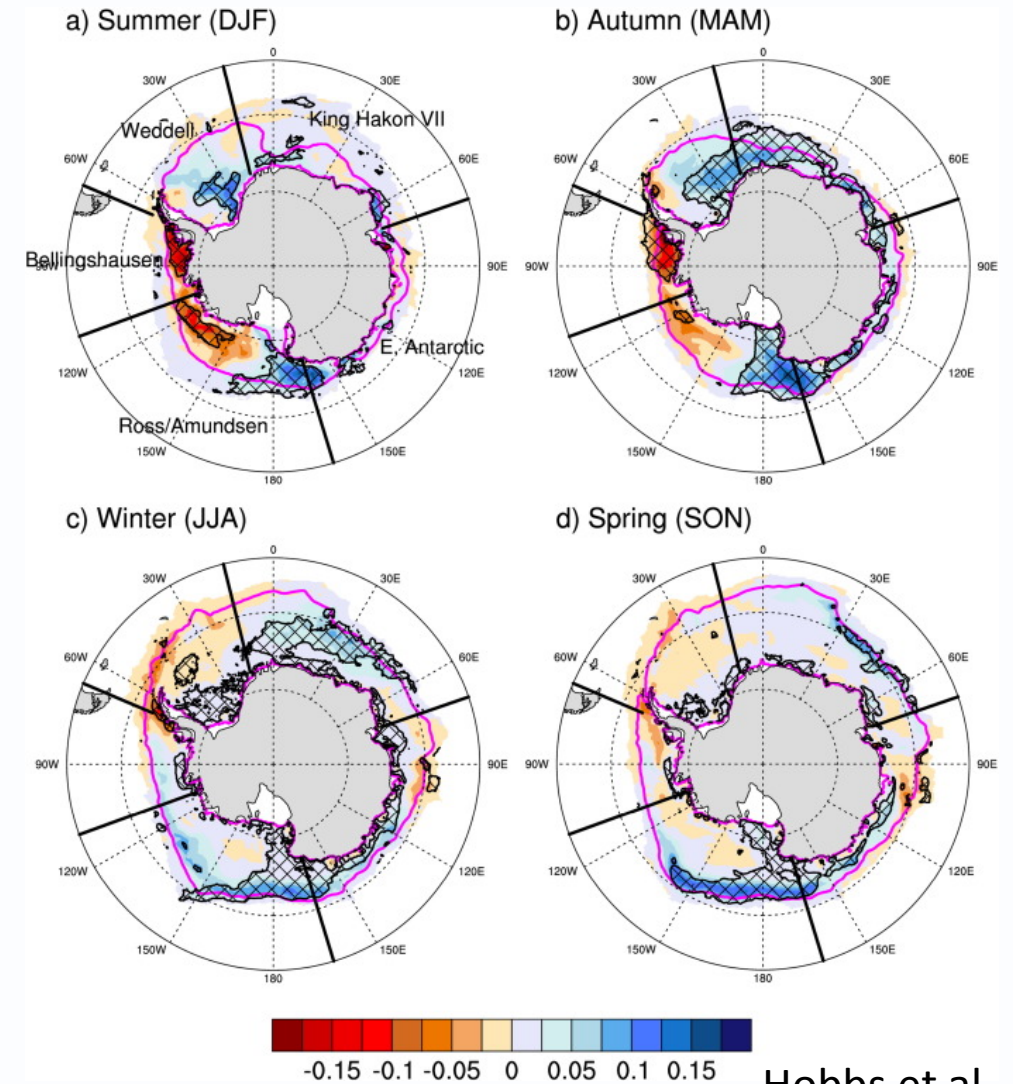


Antarctic sea ice



<https://neptune.gsfc.nasa.gov/csb/index.php?section=234>

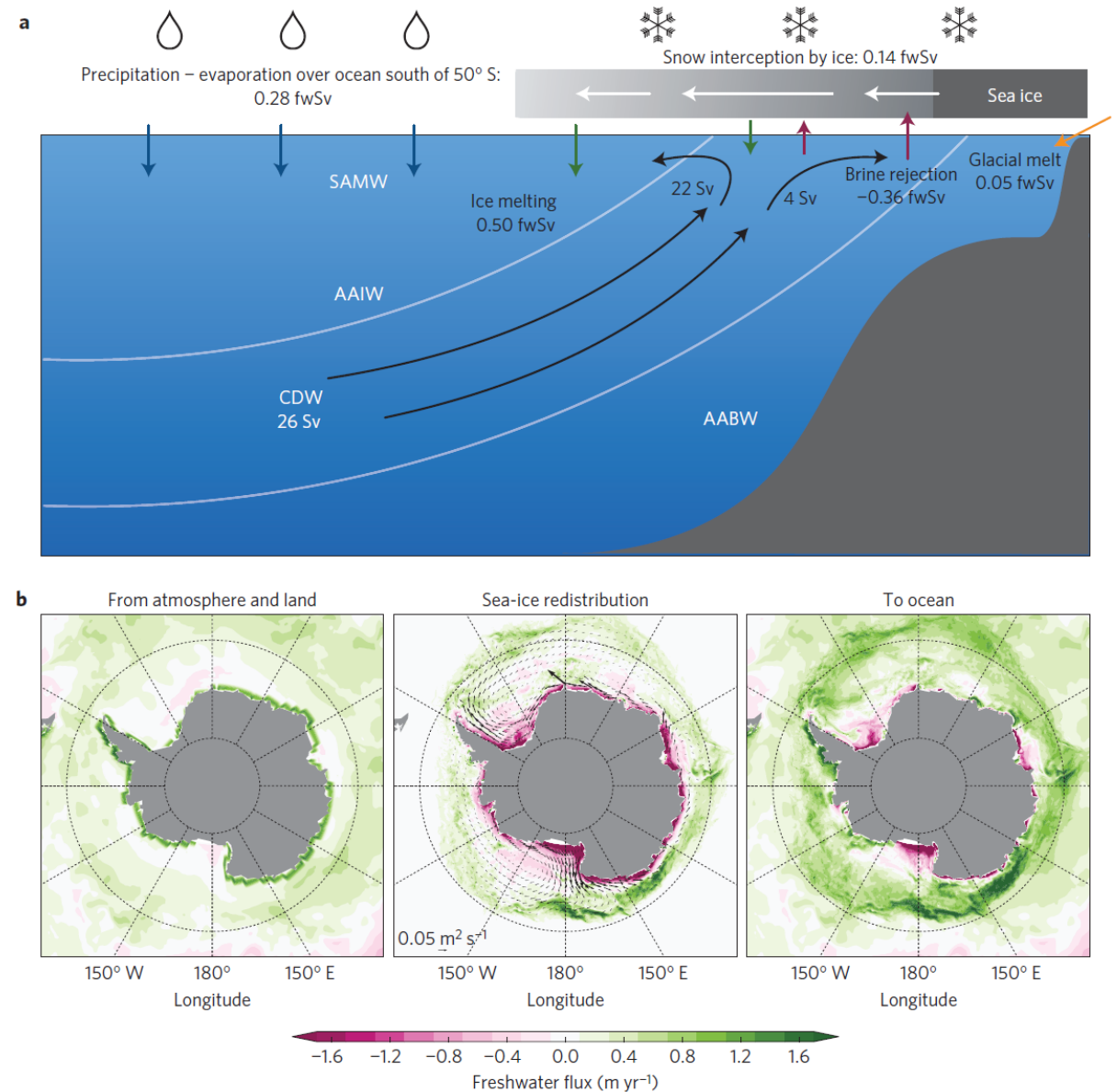
Sea ice concentration trends (fraction/decade)



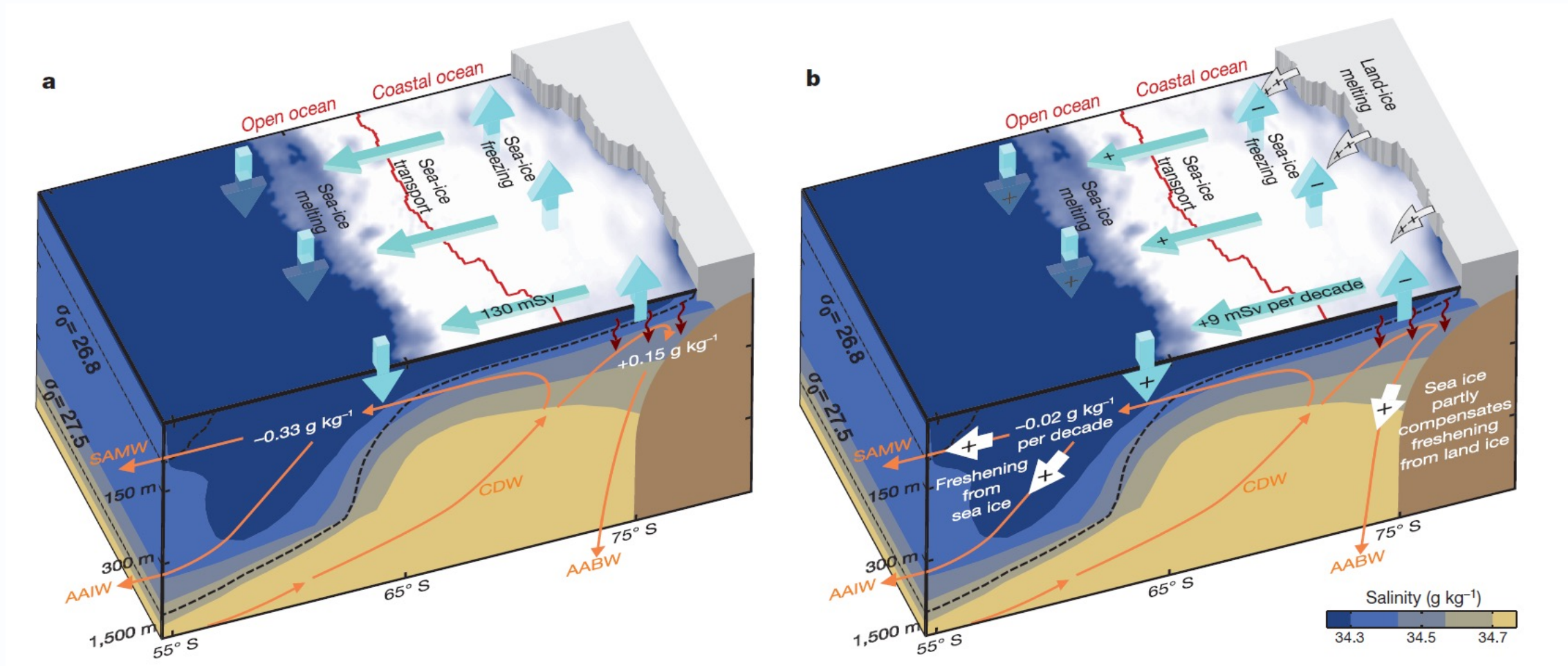
Hobbs et al., 2016

Water mass transformation by the sea ice freshwater conveyor belt

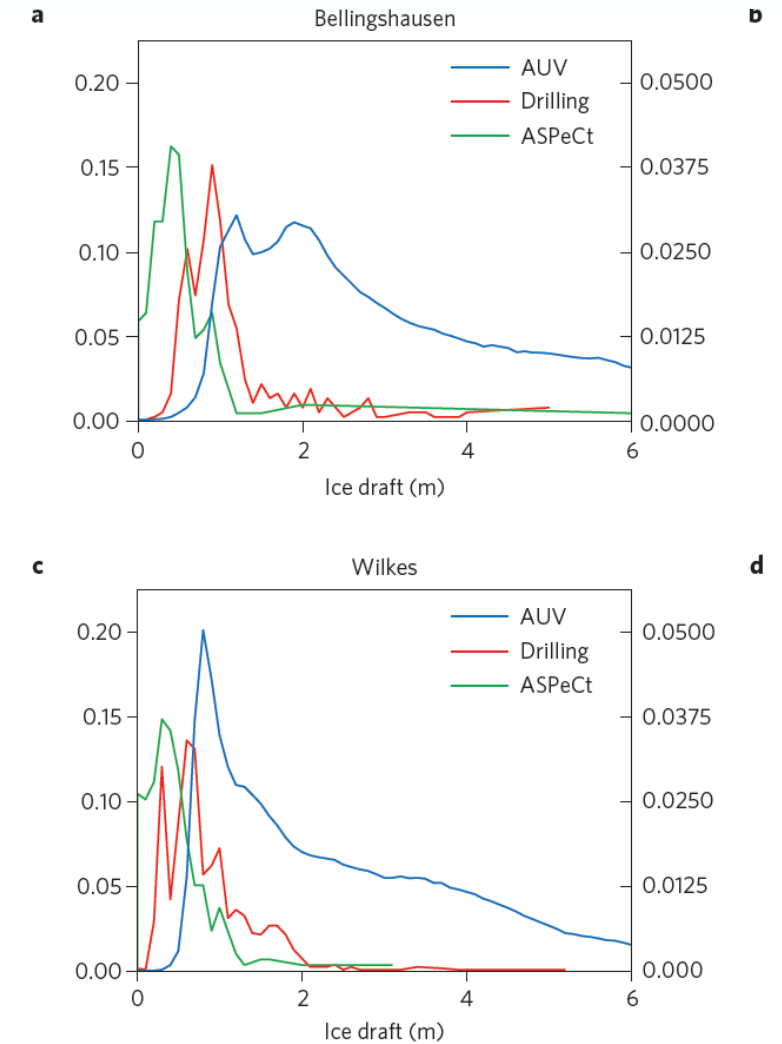
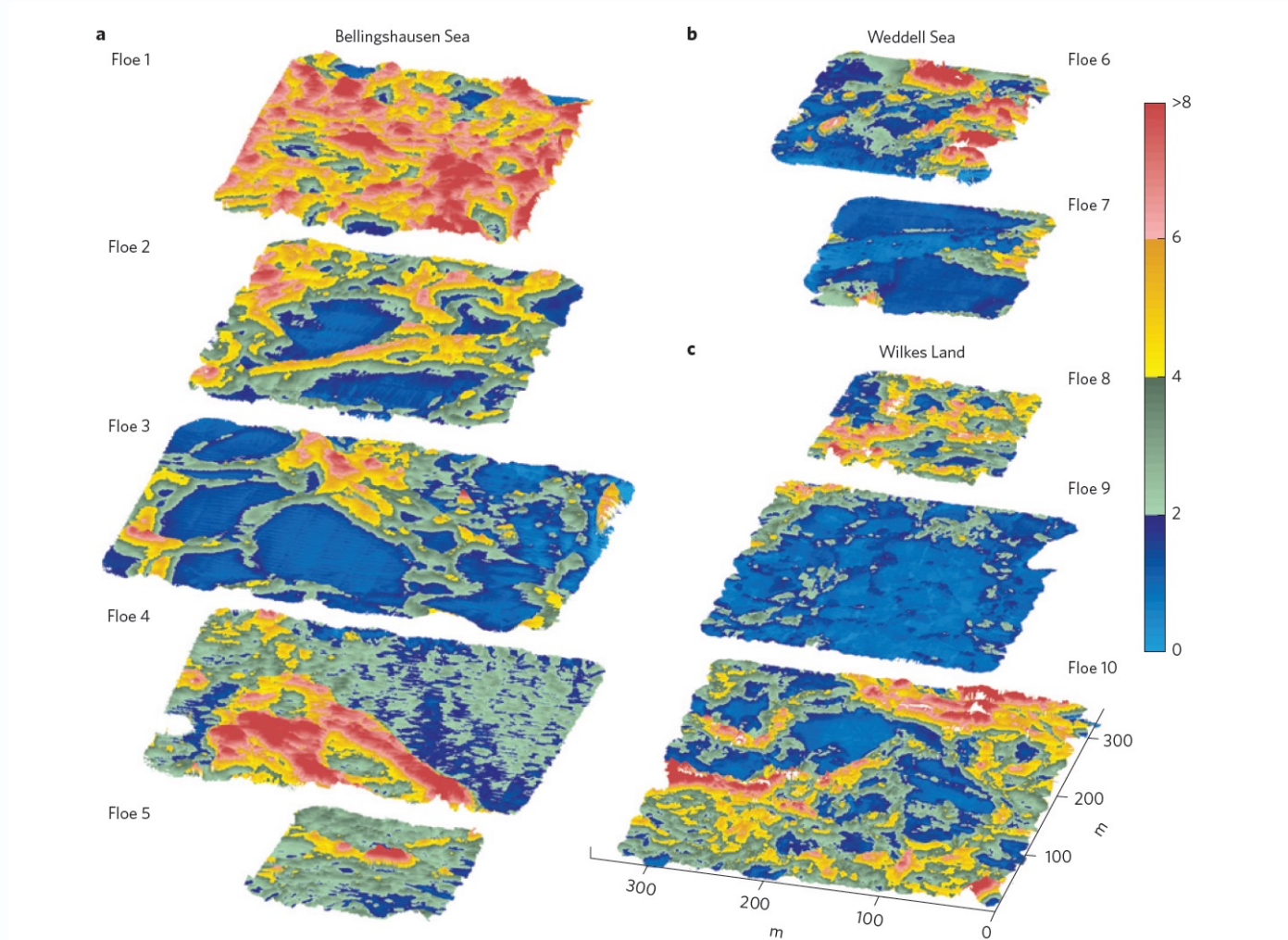
“The intricate, temporally variable relationship between ice formation, freshwater fluxes, and mixing highlights the need for year-round observations to accurately quantify the net water-mass transformation driven by sea-ice processes.”



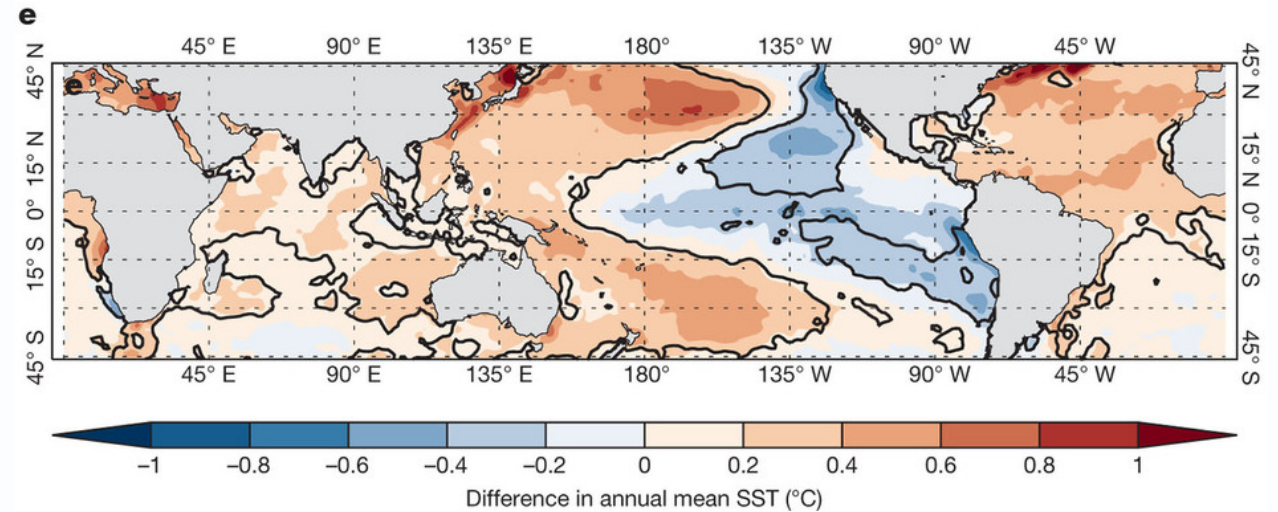
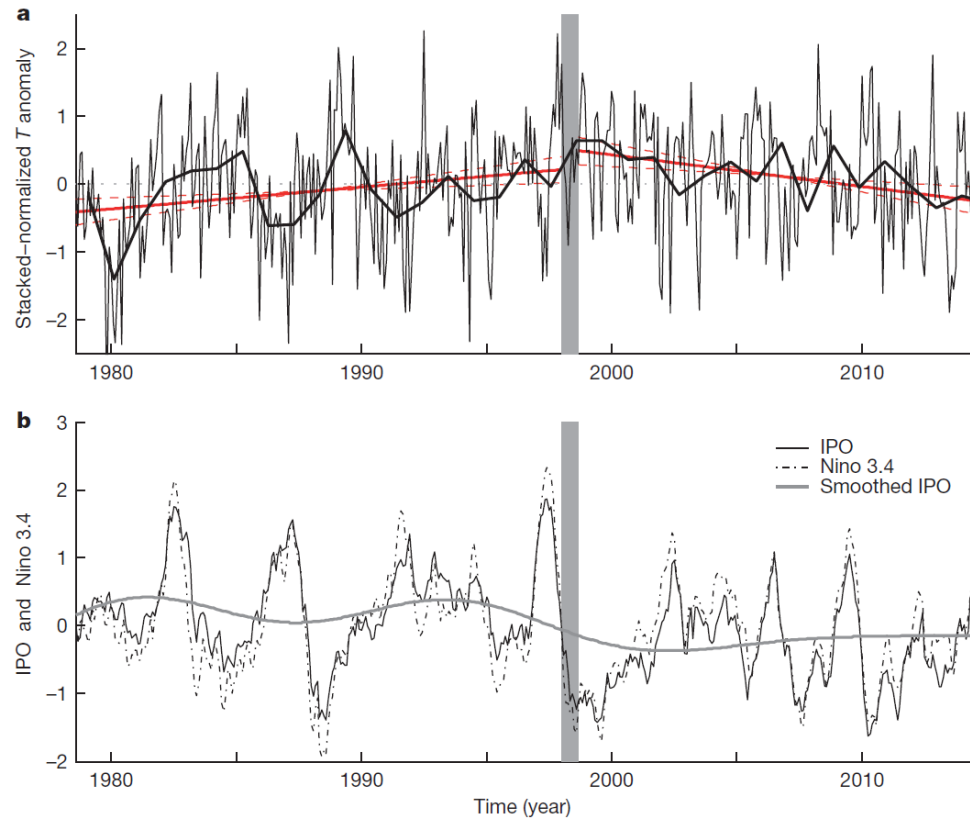
Increased sea ice transport explains freshening trends



How thick is Antarctic sea ice?



Natural variability and teleconnections



Theme 3: Ocean – ice shelf interaction



Marine Ice Sheet Instability underway in West Antarctica?

Widespread, rapid grounding line retreat of Pine Island, Thwaites, Smith, and Kohler glaciers, West Antarctica, from 1992 to 2011

E. Rignot^{1,2}, J. Mouginot¹, M. Morlighem¹, H. Seroussi², and B. Scheuchl¹ GRL, 2014

Marine Ice Sheet Collapse Potentially Underway for the Thwaites Glacier Basin, West Antarctica

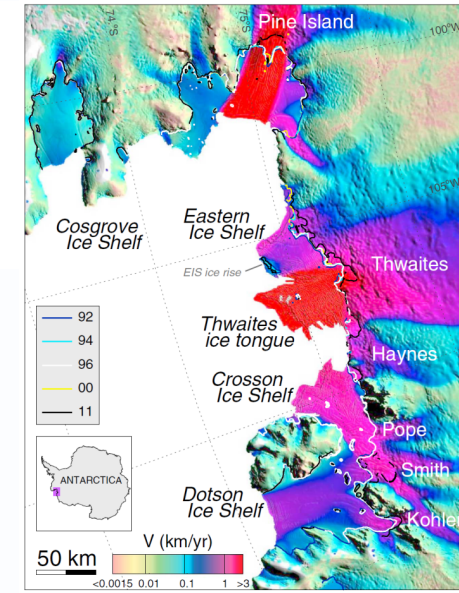
Ian Joughin*, Benjamin E. Smith, Brooke Medley

Science, 2014

Retreat of Pine Island Glacier controlled by marine ice-sheet instability

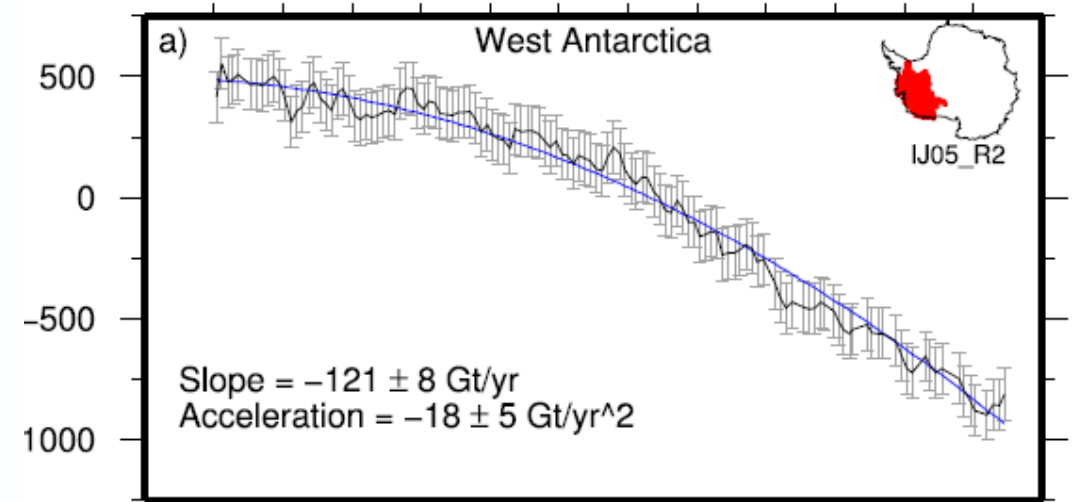
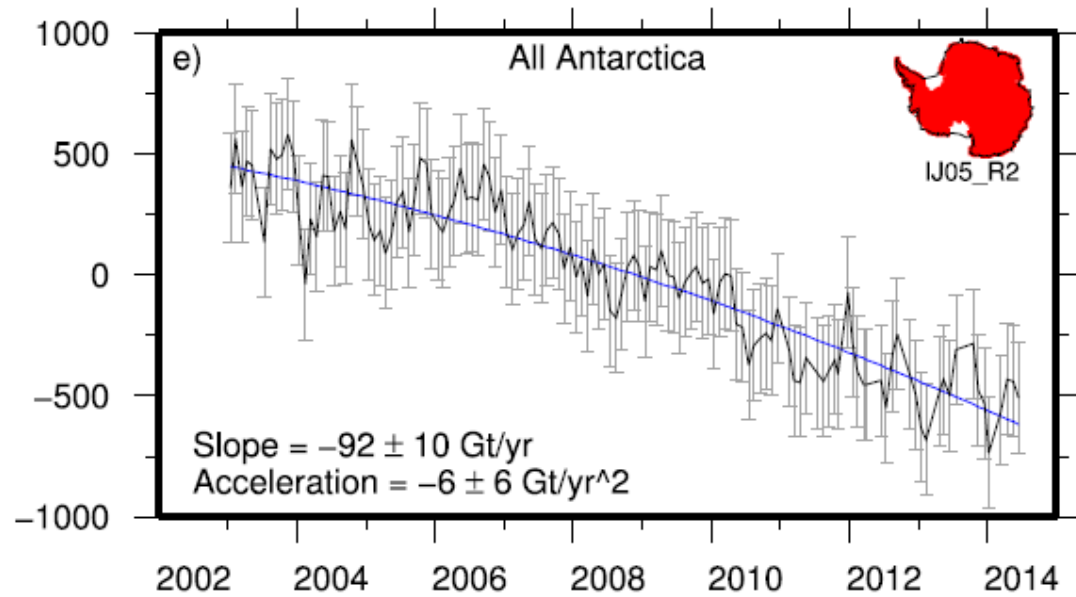
L. Favier^{1,2}, G. Durand^{1,2*}, S. L. Cornford³, G. H. Gudmundsson^{4,5}, O. Gagliardini^{1,2,6},
F. Gillet-Chaulet^{1,2}, T. Zwinger⁷, A. J. Payne³ and A. M. Le Brocq⁸

Nature Climate Change, 2014

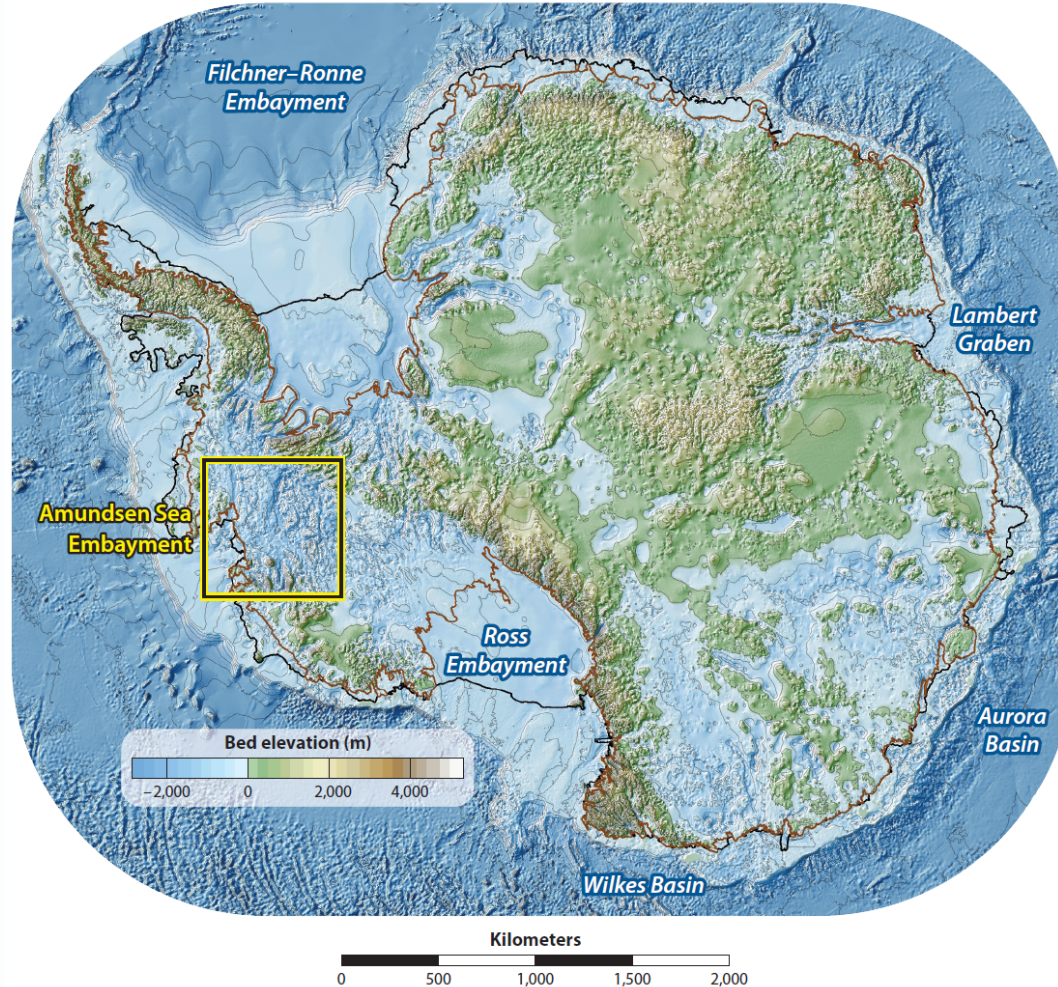


Rignot et al 2014

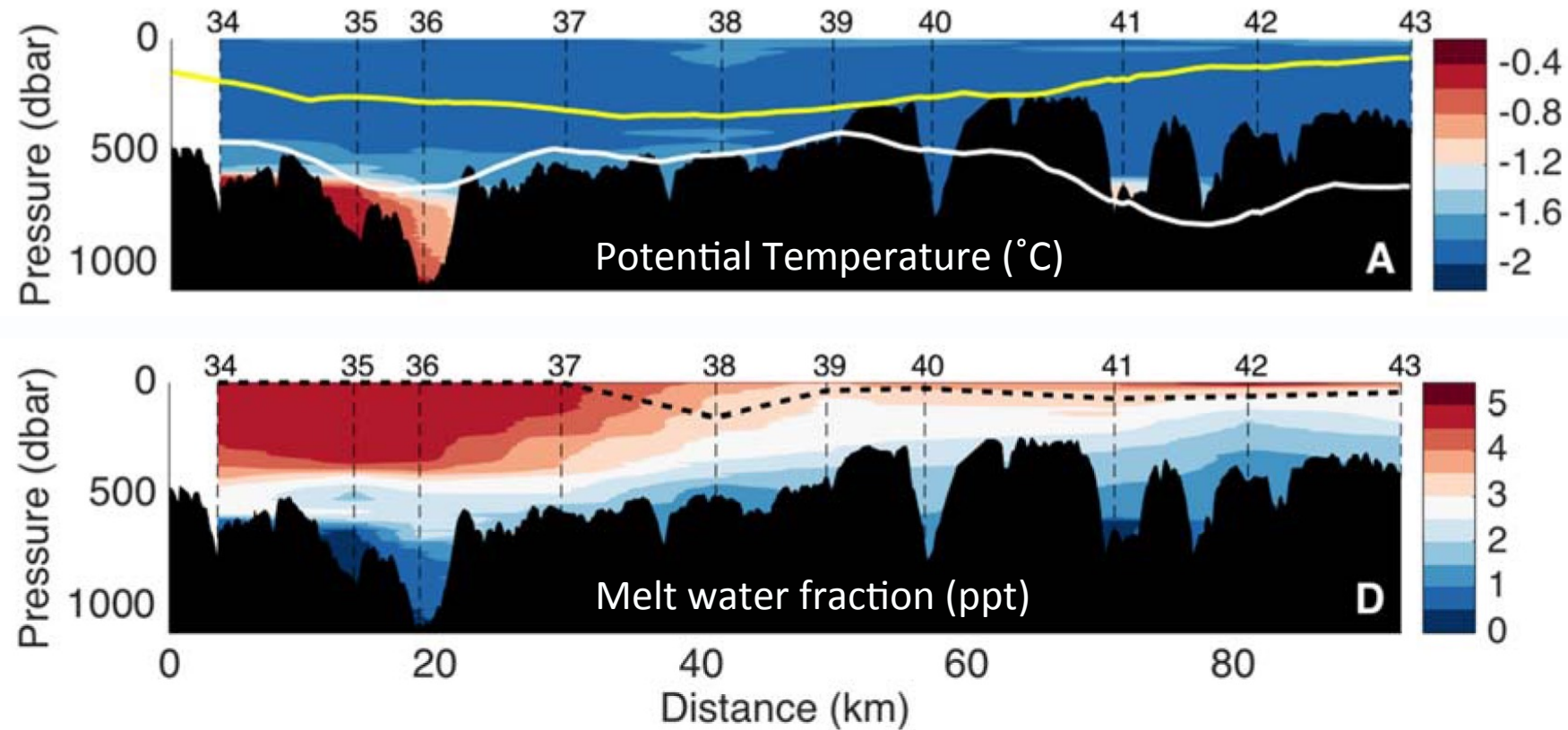
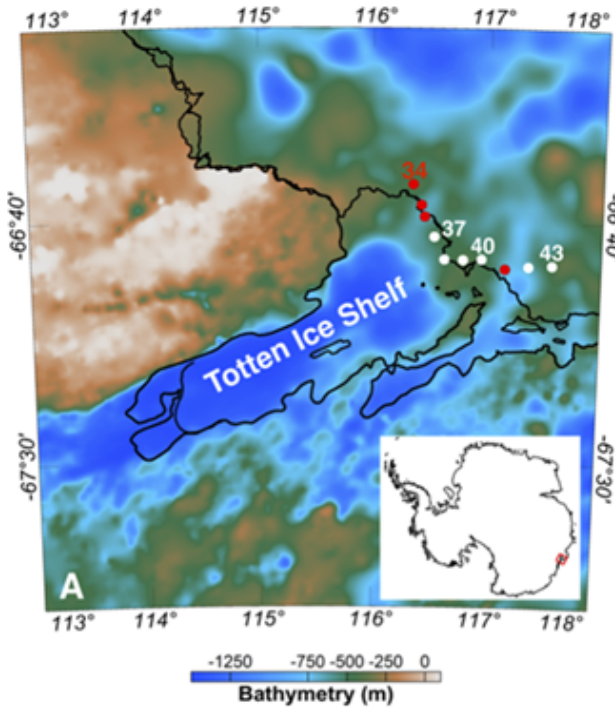
Acceleration of mass loss from ice shelves



Much of the Antarctic Ice Sheet is grounded below sea level

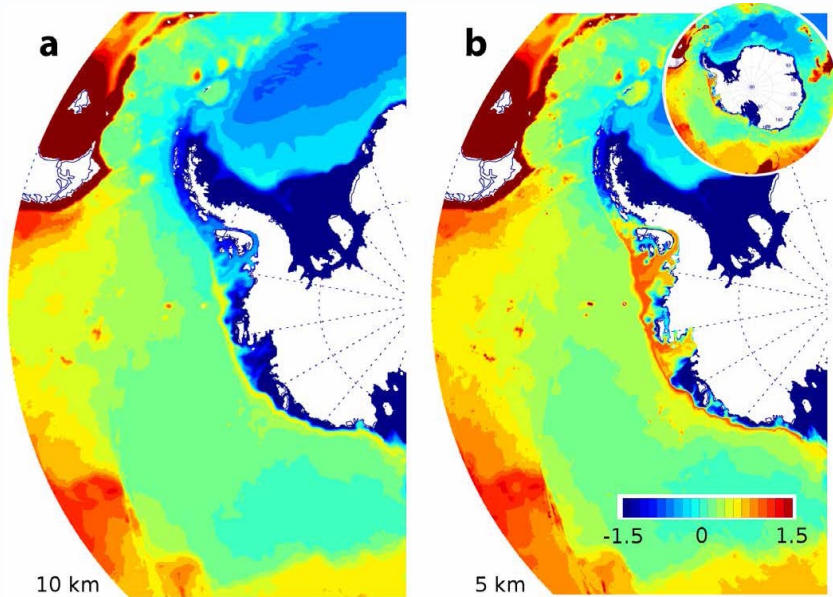


Warm water inflow, melt water outflow from ice shelf cavity

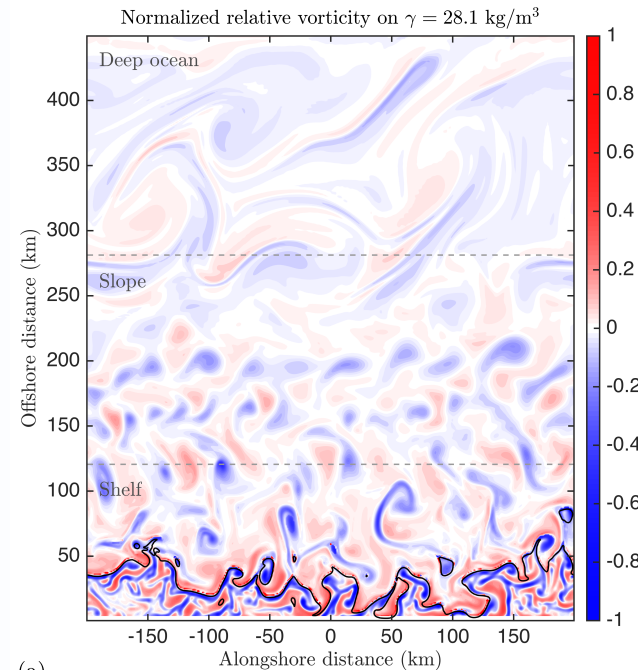


Rintoul et al., Science Advances 2016; Silvano et al., Oceanography, 2016; Silvano et al., JGR-Oceans, 2017

How does warm water reach ice-shelf cavities?

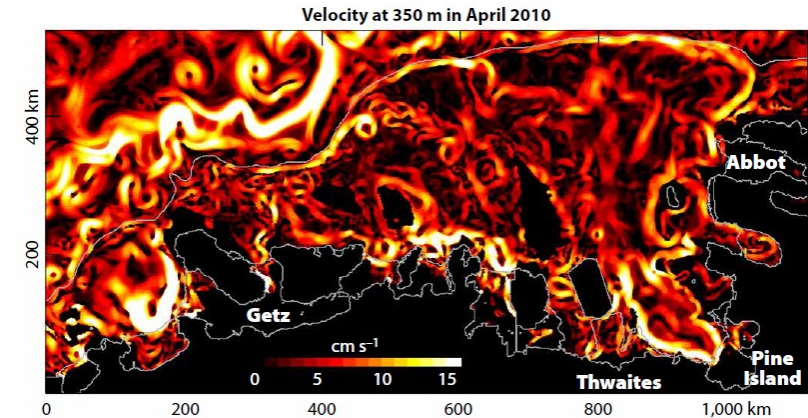


Dinniman et al., 2015; Nakayama et al., 2014



(a)

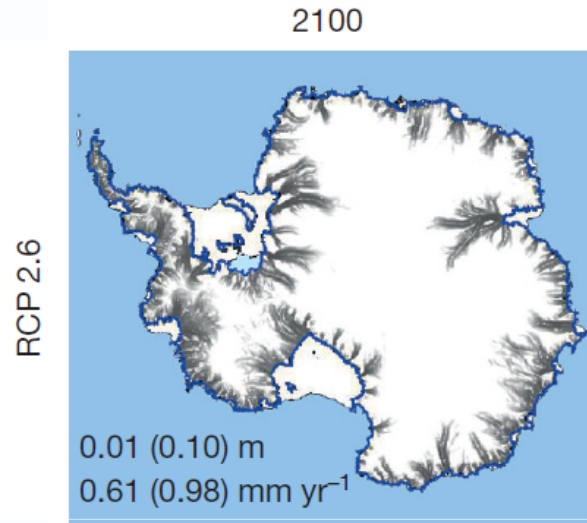
Stewart and Thompson, 2016



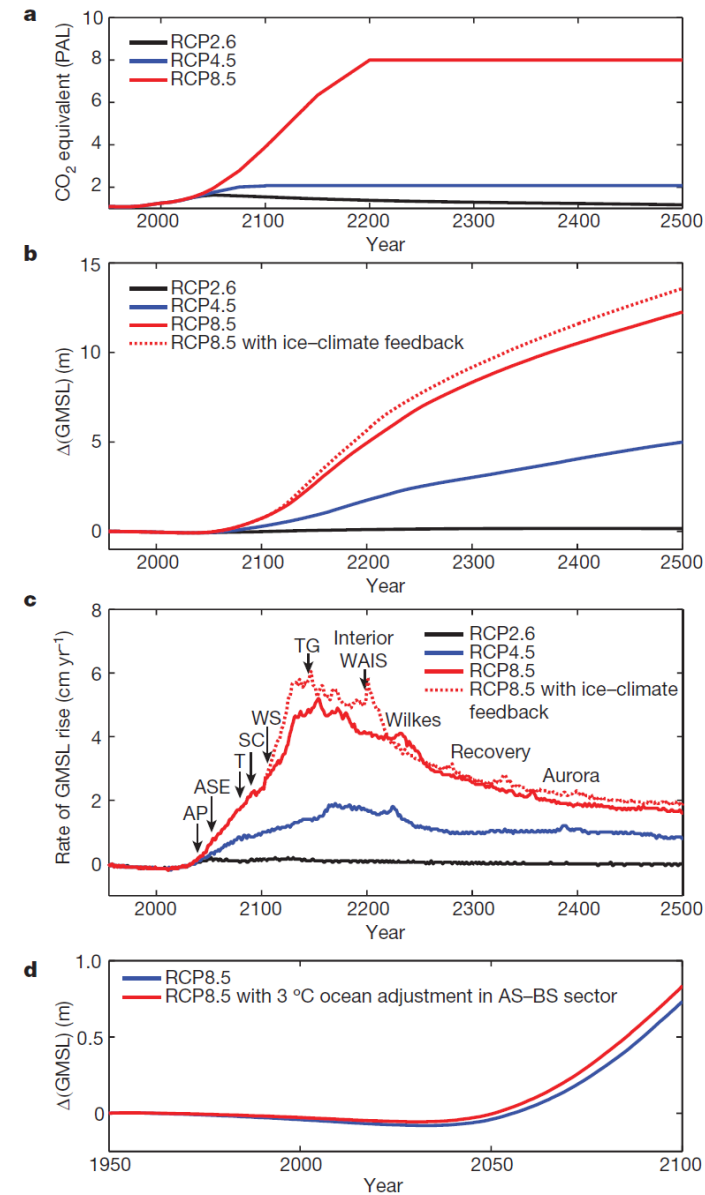
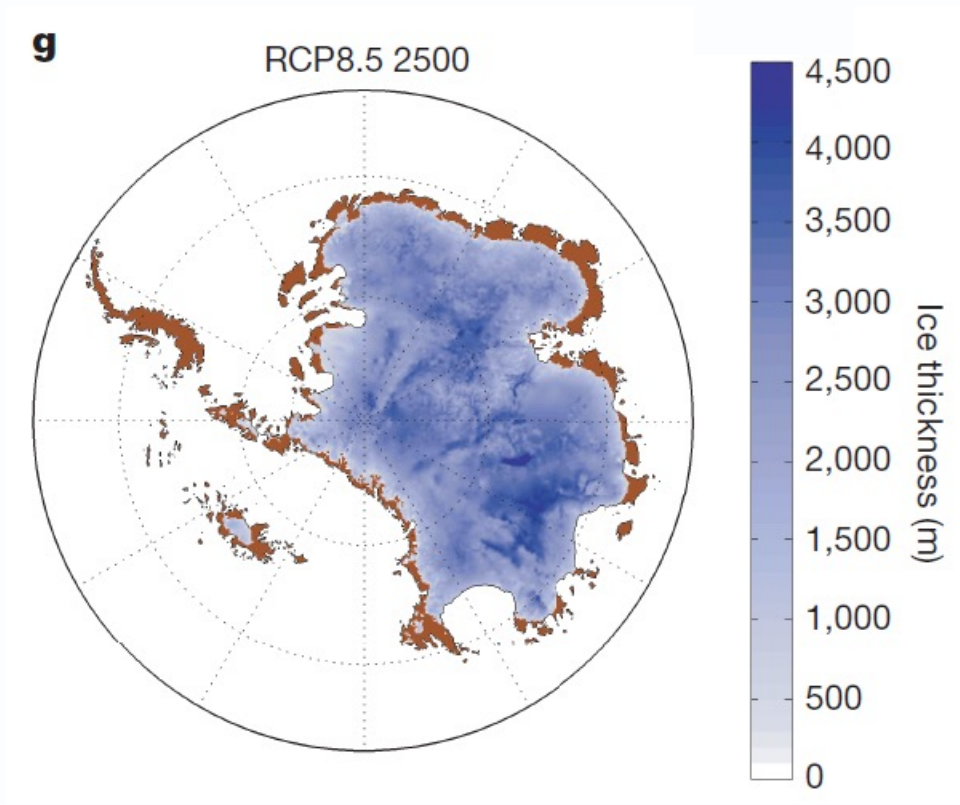
St Laurent et al., 2015

Many factors influence heat transport to ice shelf cavities: eddies, tides, wind, Antarctic Slope Front, topographic interactions, polynya dynamics, and often act on small spatial scales.

Antarctic commitment to future sea level rise



Hydrofracturing and ice cliff failure result in large and rapid contribution to sea level rise

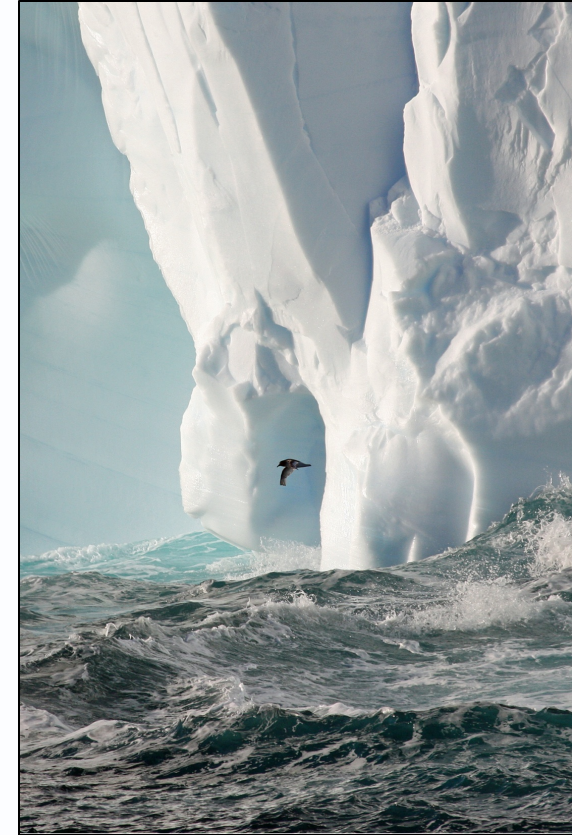


Summary

- The report from the 2012 workshop provides a strong foundation for a new strategy for under – ice observations.
- Recent advances provide even stronger motivation for under – ice observations.
- Substantial progress has been made in the past 5 years.
- Our job is to build on the previous report to articulate a compelling vision for under – ice observations in Antarctica and to provide a roadmap for practical implementation of an integrated, circumpolar effort to fill this blind spot in the global ocean observing system.

Outline

- Seeing below the ice (2012): themes & objectives
- A brief and eclectic review of new discoveries motivating under – ice observations

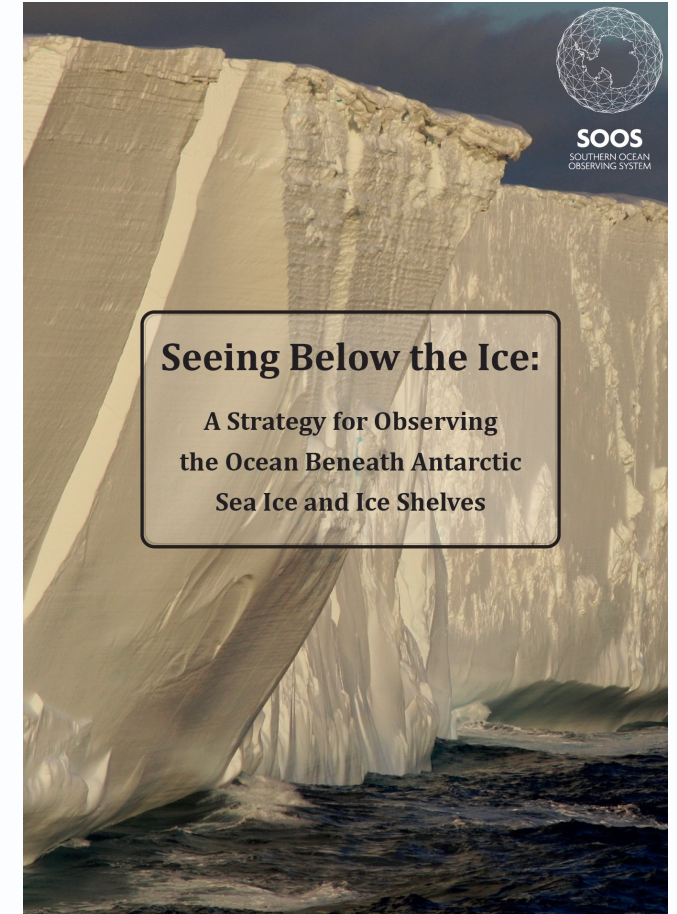


E. van Wijk

Seeing below the ice v1.0

- Result of a workshop held in Hobart in 2012
- Articulated a community view of the motivation for under-ice observations
- Identified scientific objectives and key questions in 3 themes:
 - Circulation and inventories of heat, freshwater and carbon in the sea ice zone
 - Ocean – sea ice interaction
 - Ocean – ice shelf interaction
- Outlined an integrated under – ice observing strategy and path to implementation

A partial success: raised the profile of under – ice observations, but did not fully succeed in motivating scientists and funding agencies to implement a comprehensive observing system.



Themes and Objectives

Circulation and inventories of heat, freshwater and carbon in the sea ice zone

- To quantify how much heat, freshwater and carbon are stored by the ocean between the winter sea ice edge and the Antarctic continent.
- To understand the processes responsible for ocean storage of heat, freshwater and carbon and their sensitivity to changes in forcing

Ocean – sea ice interaction:

- To determine the processes controlling the circumpolar and regional distribution of sea ice concentration and thickness.
- To determine how and why the concentration/thickness of sea ice varies over time-scales from days to millennia.
- To understand and quantify coupled interactions between Antarctic sea ice, the ocean, the atmosphere and ice shelves.

Ocean – ice shelf interaction:

- To determine the sensitivity of Antarctic ice shelves to changes in ocean circulation & temperature.
- To assess the effect of basal melt of floating ice shelves on the mass balance of the Antarctic Ice Sheet and its contribution to sea level rise.
- To determine the response of the ocean to changes in freshwater flux input from Antarctica.

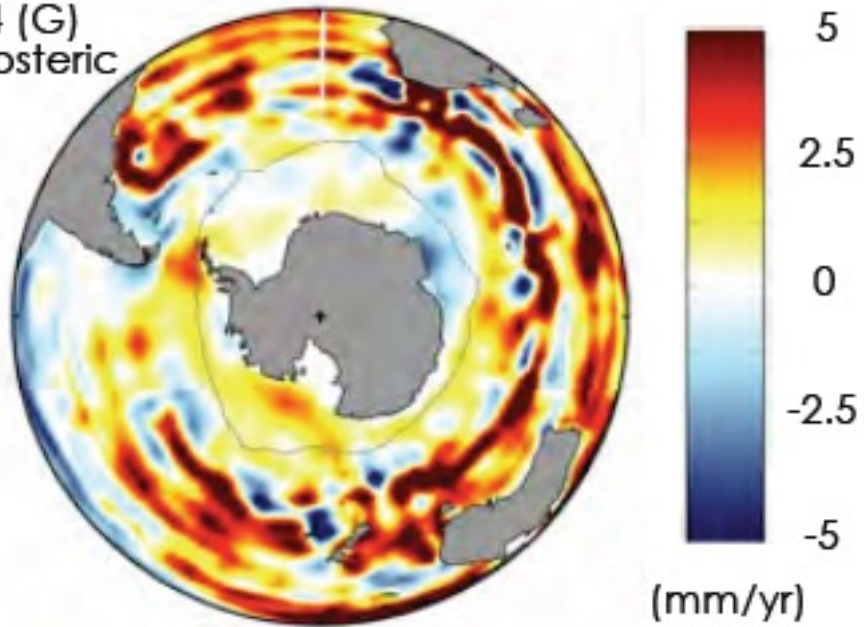
Bea current meter study (i.e. circulation in sea ice zone)

Something from Naveira – Garabato paper?

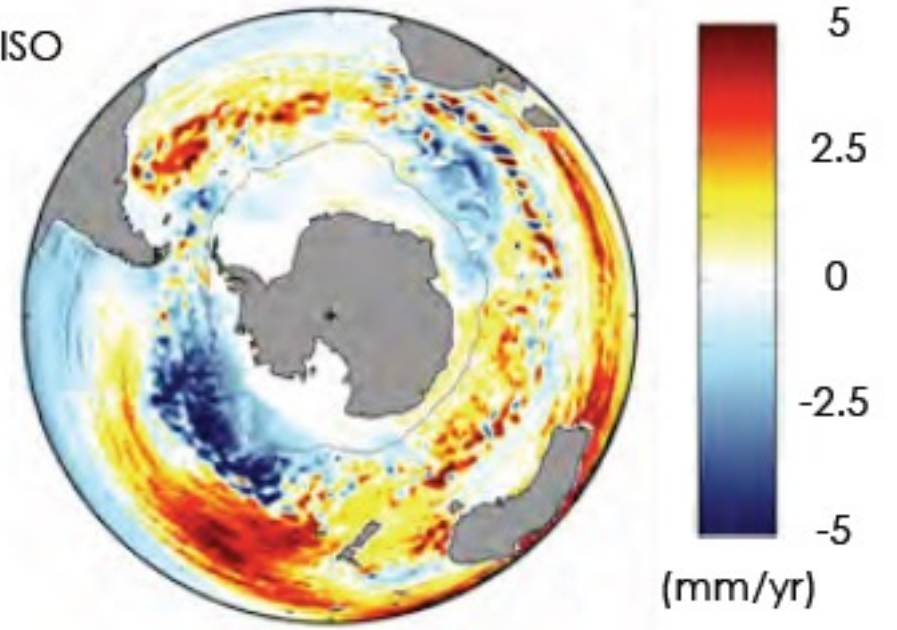
Sea-level trends

Observations: trends in sea-level

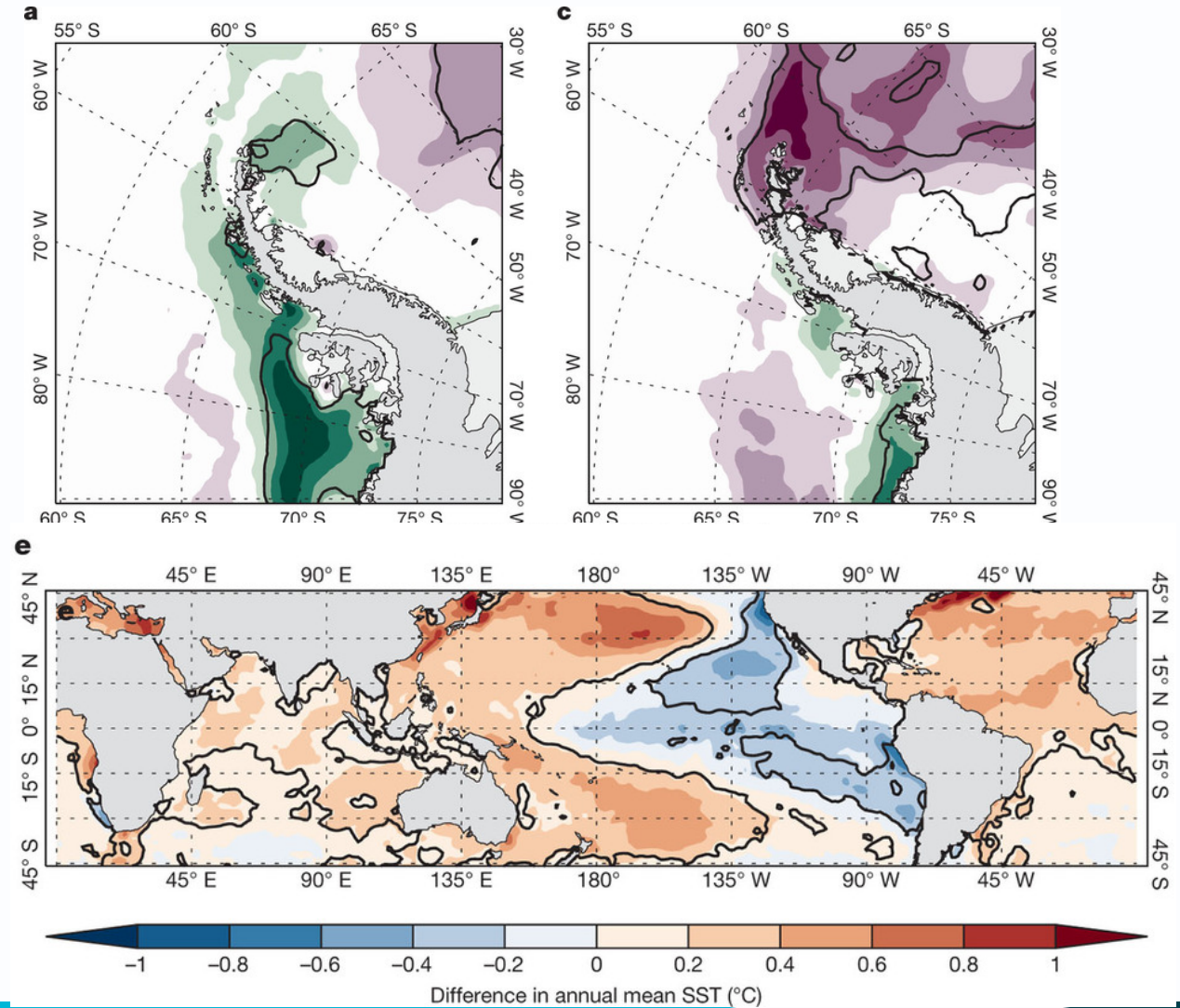
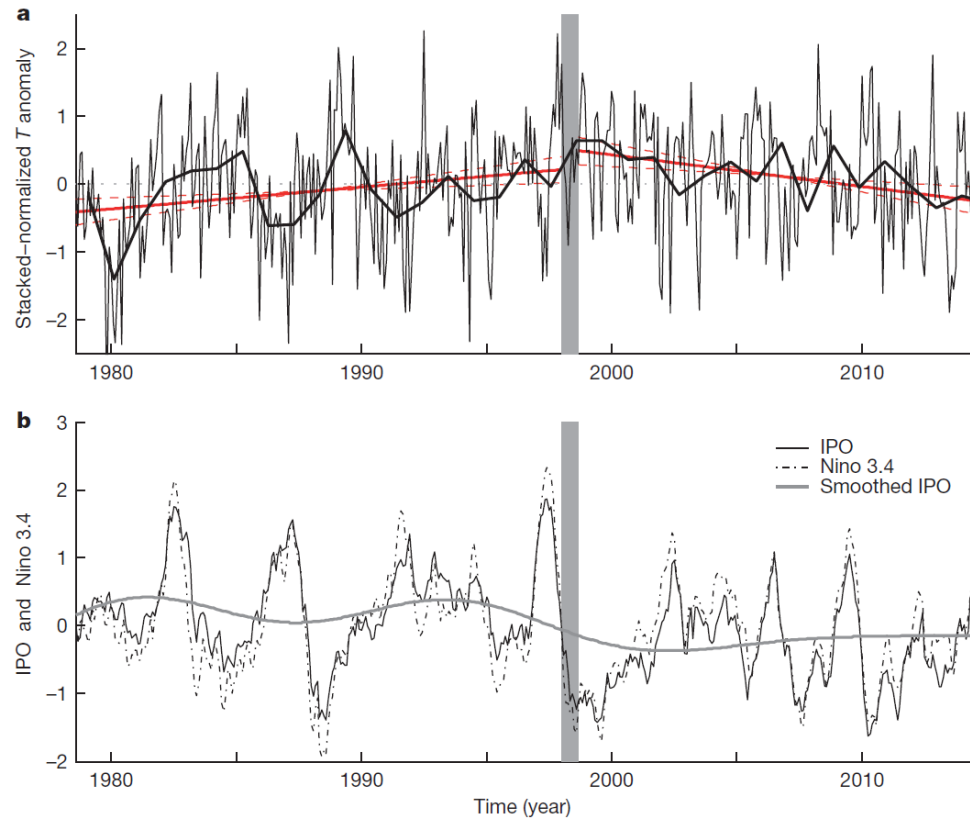
a EN4 (G)
thermosteric



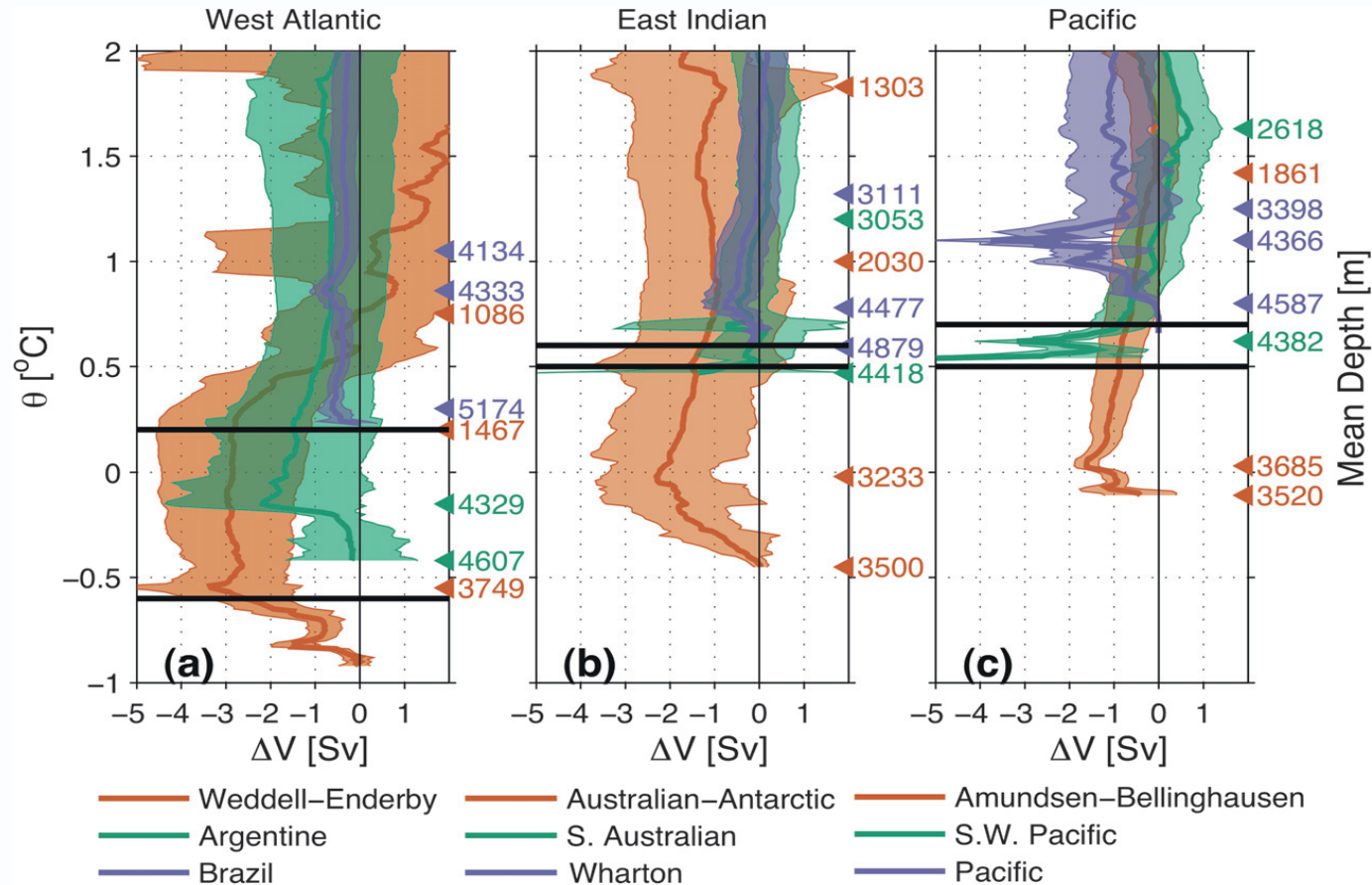
b AVISO



Natural variability and teleconnections explain trends on AP

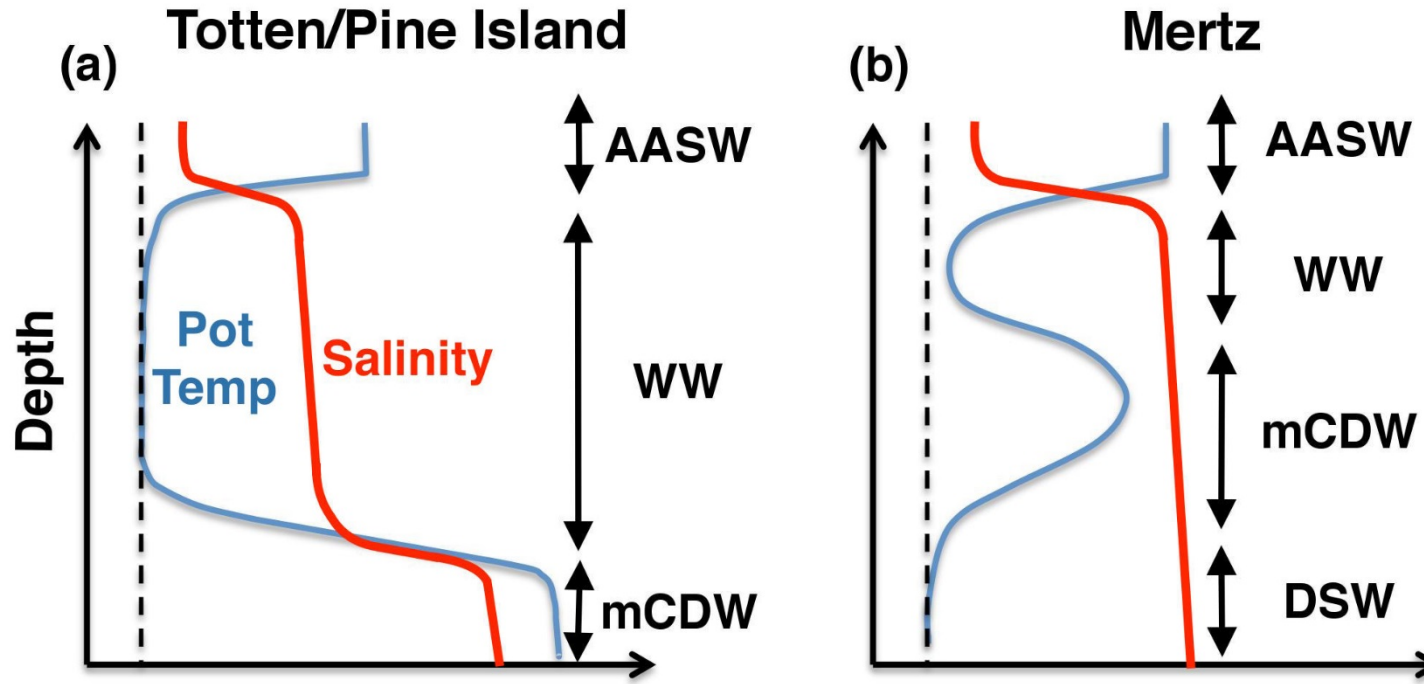


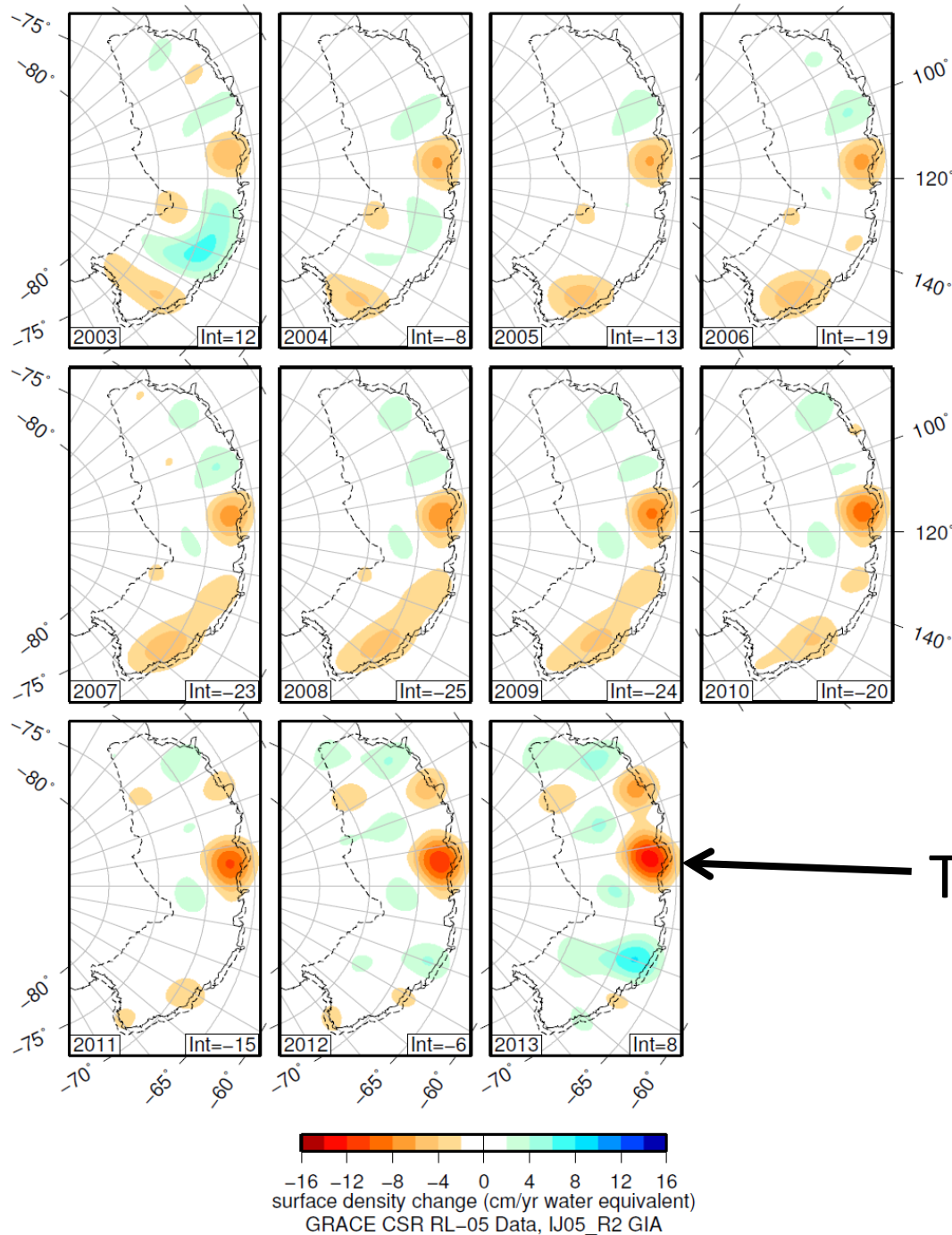
Warming, freshening and contraction of Antarctic Bottom Water



Purkey & Johnson, 2013

Ocean stratification at Totten similar to West Antarctica



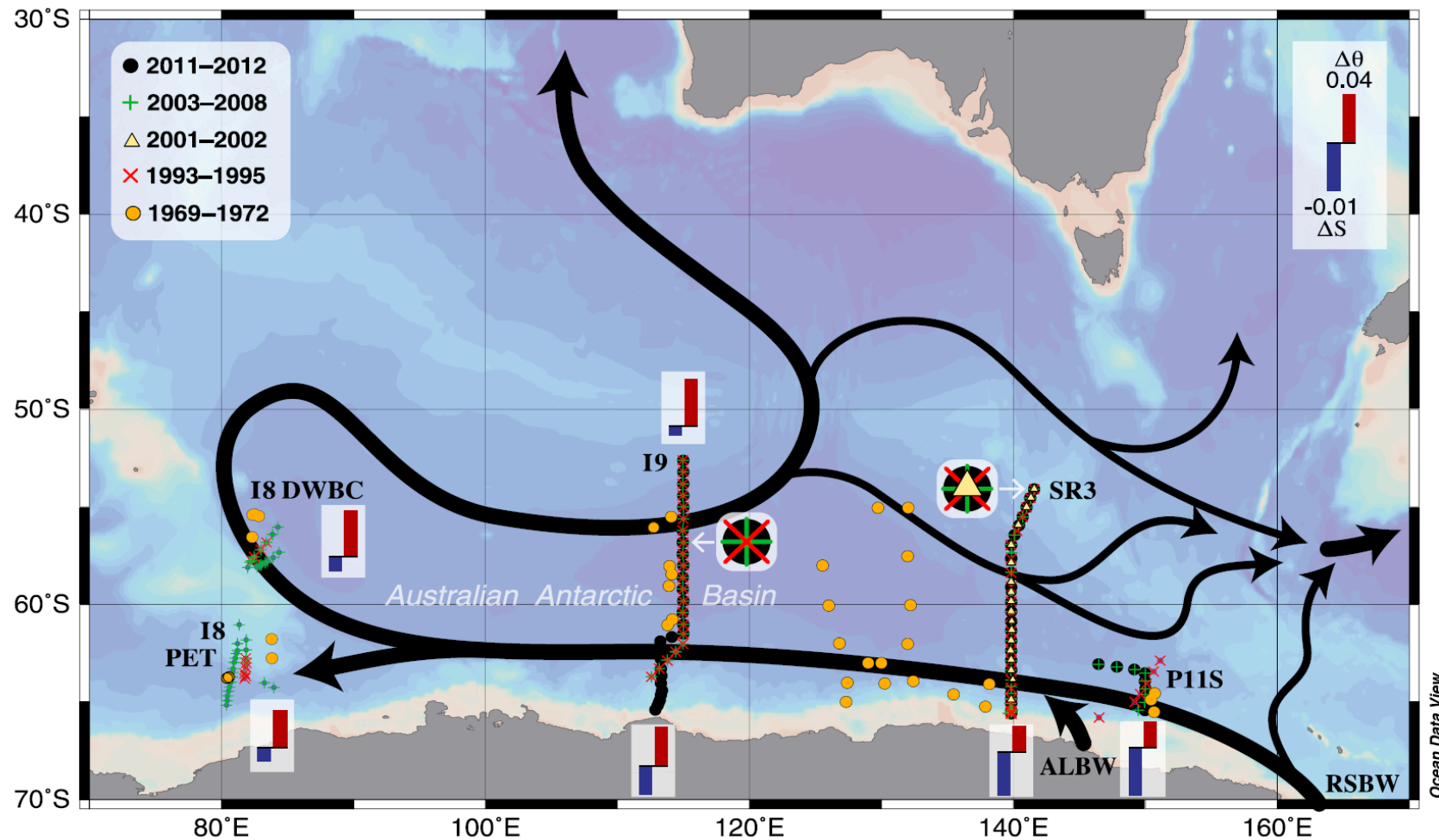


Loss of mass from East Antarctica (cm/yr) between 2003 to 2013 from satellite gravity measurements.

← Totten Glacier

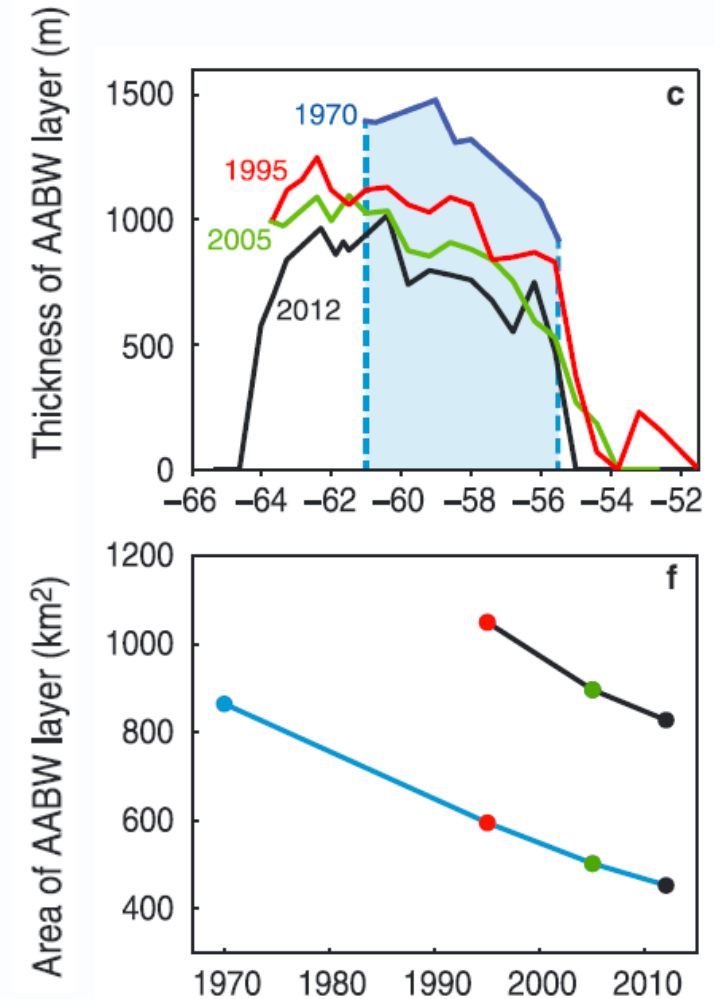
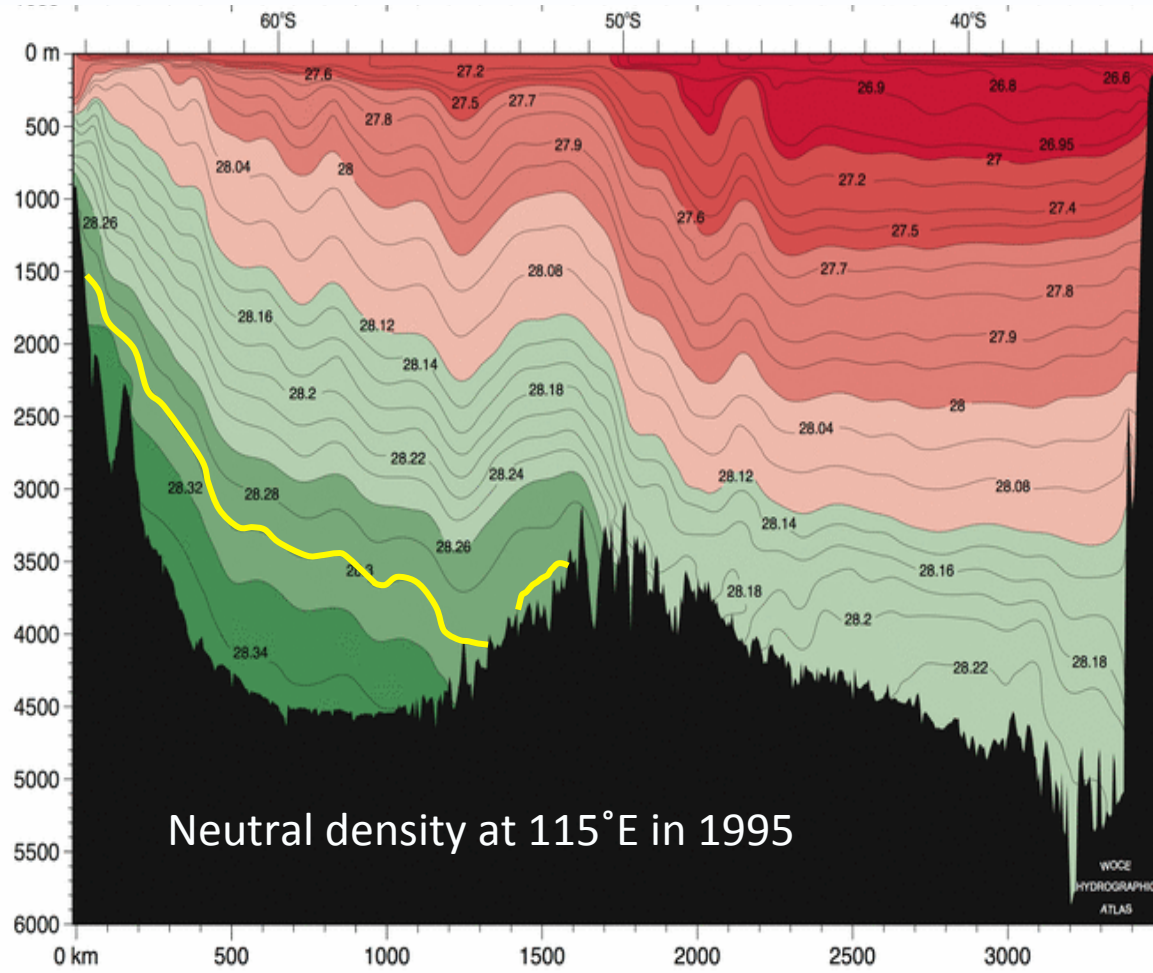
Harig & Simons 2015 (supp mat.)

Antarctic Bottom Water in the Australian – Antarctic Basin

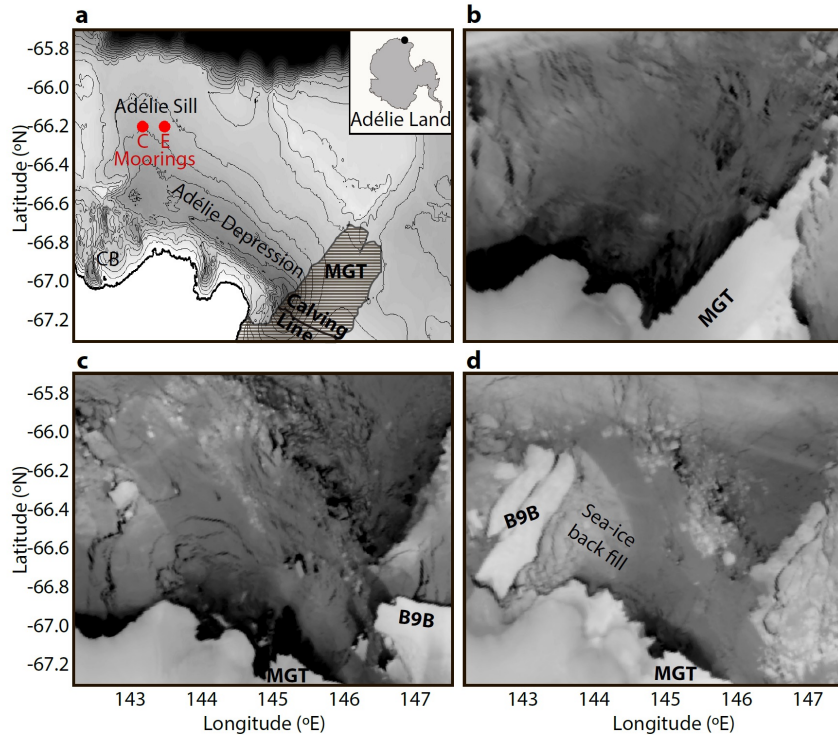


van Wijk and
Rintoul, GRL, 2014

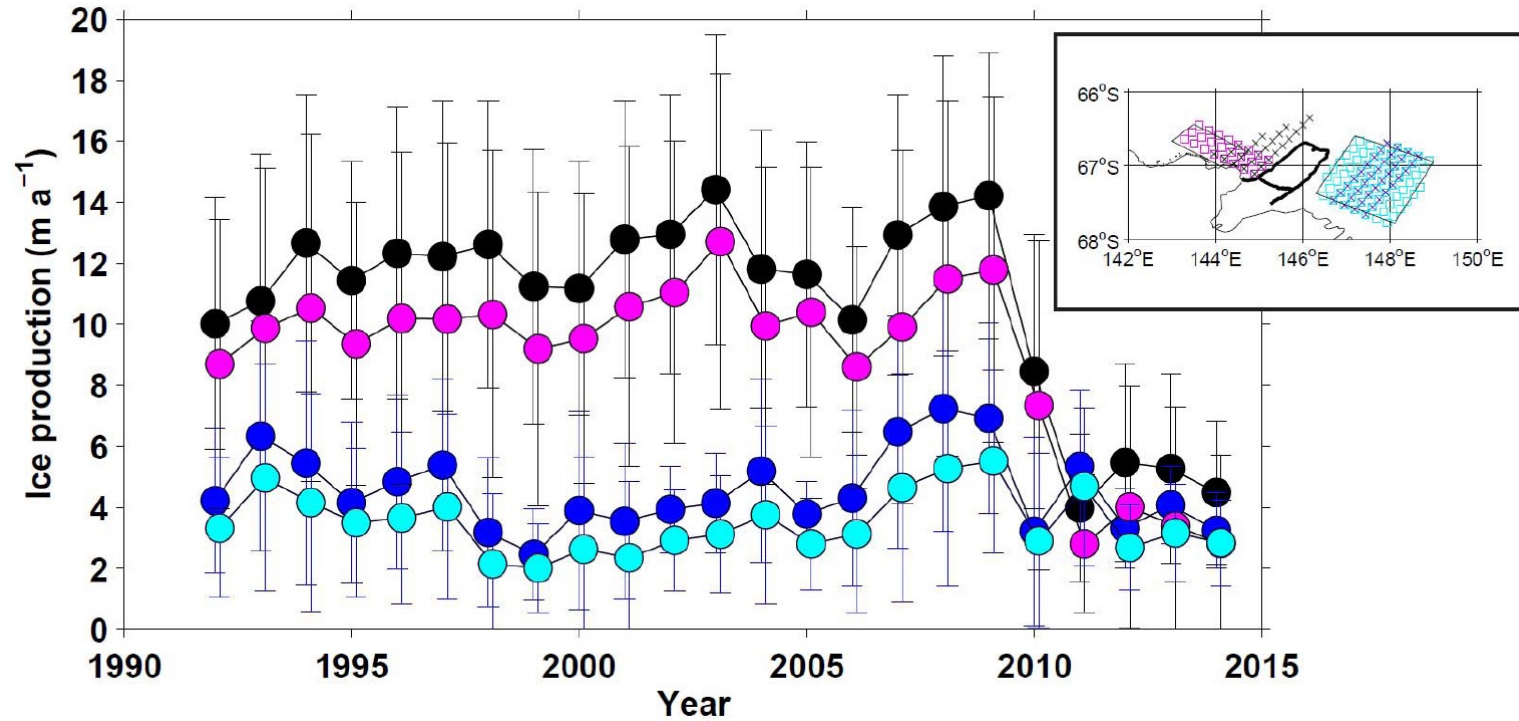
Antarctic Bottom Water layer contracted by ~50% between 1970 and 2012



A natural experiment: calving of the Mertz Glacier Tongue

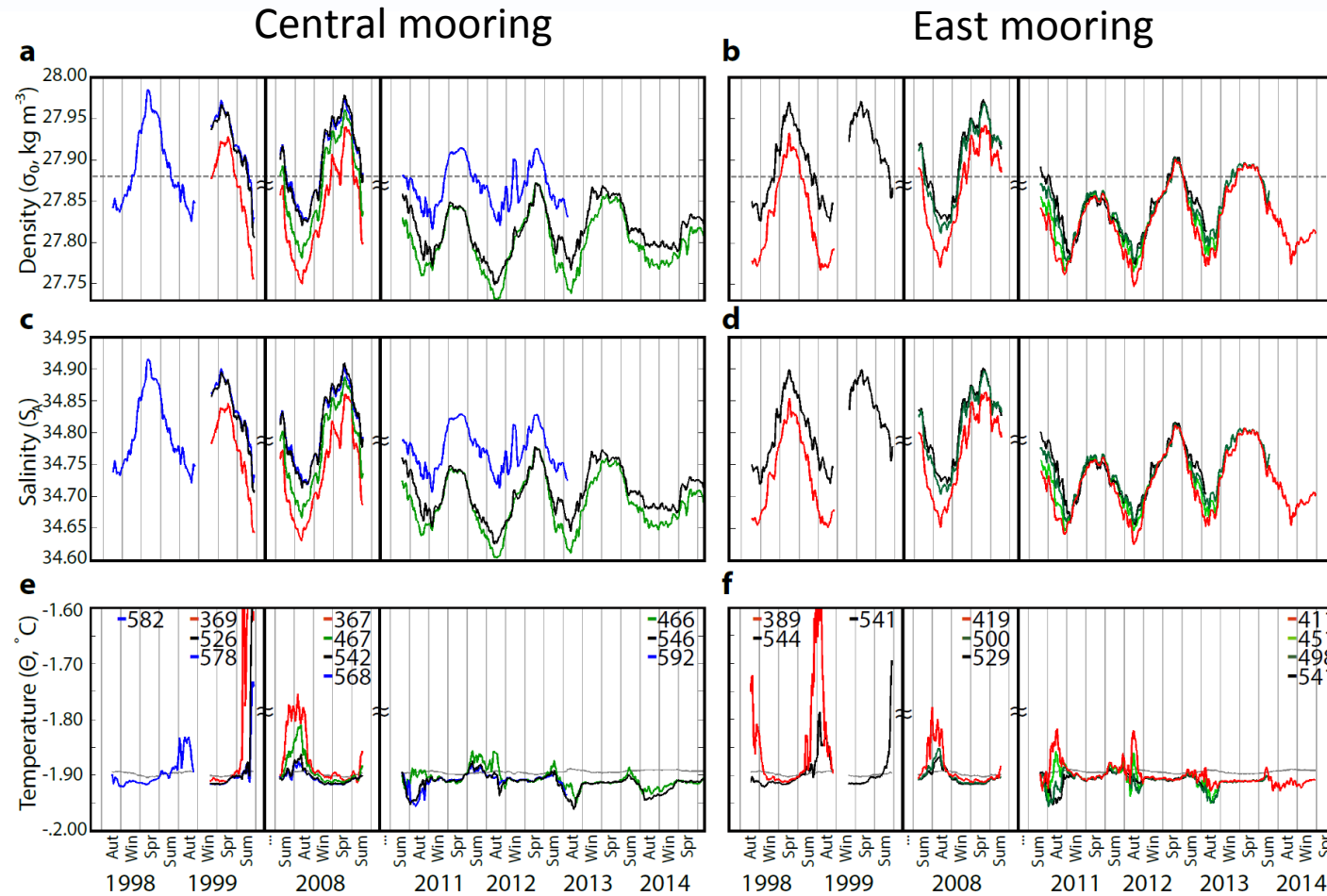


Snow et al., in prep



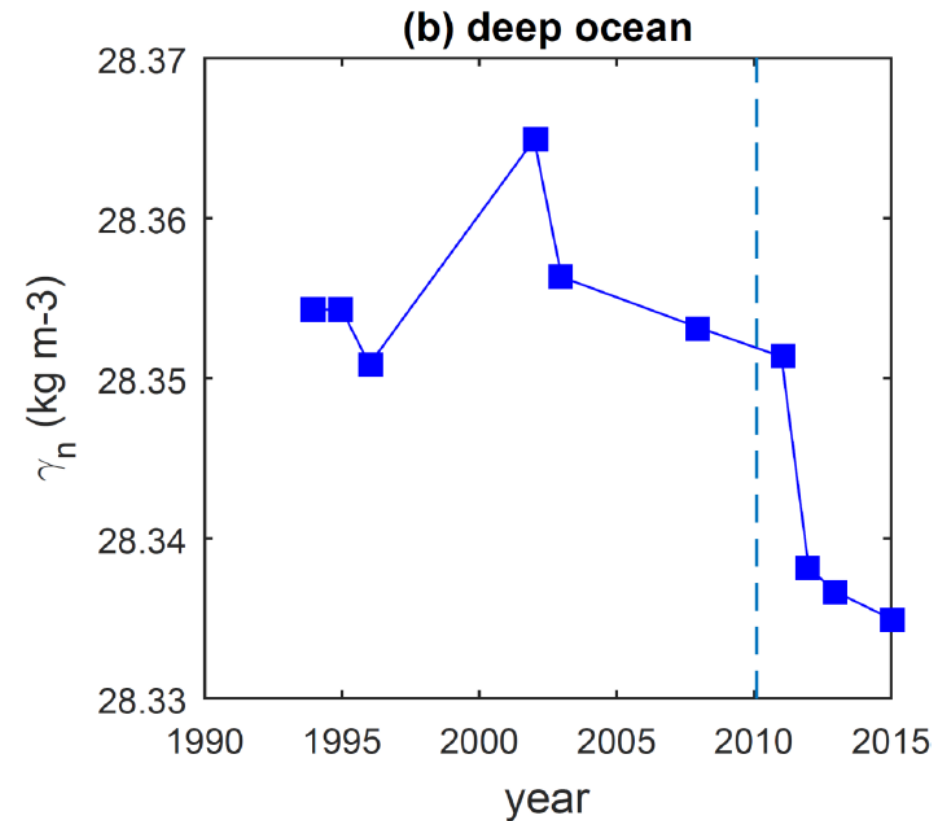
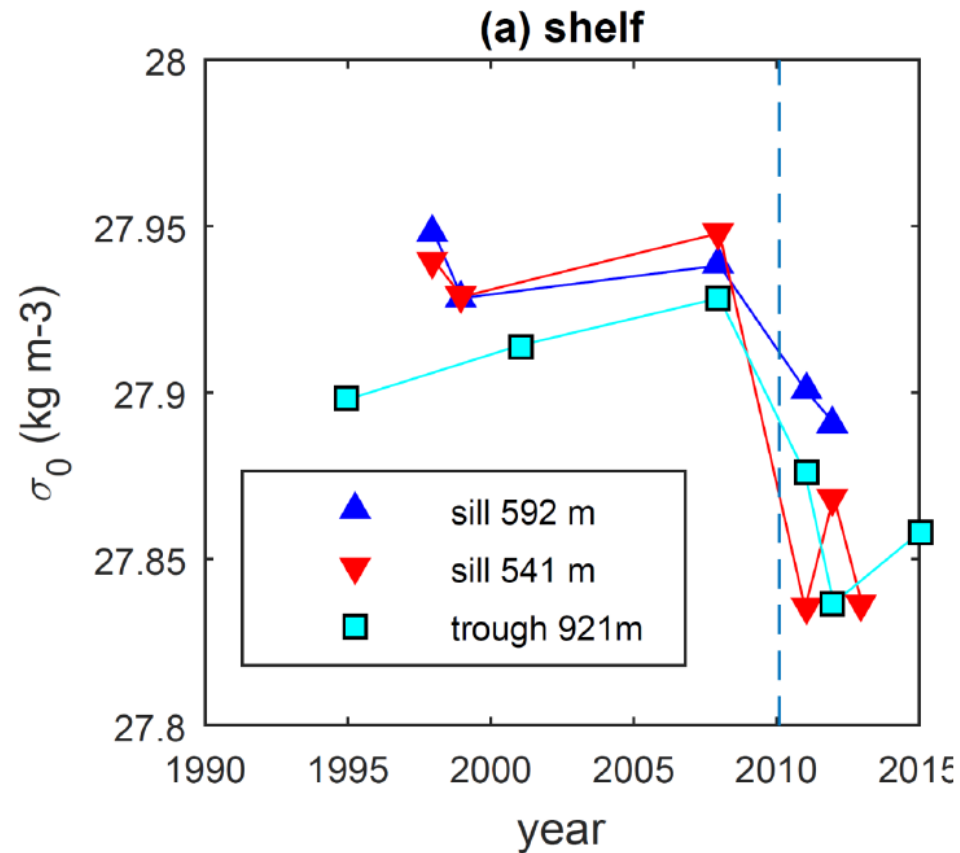
Aoki et al., in prep

Sharp reduction in density of dense shelf water export

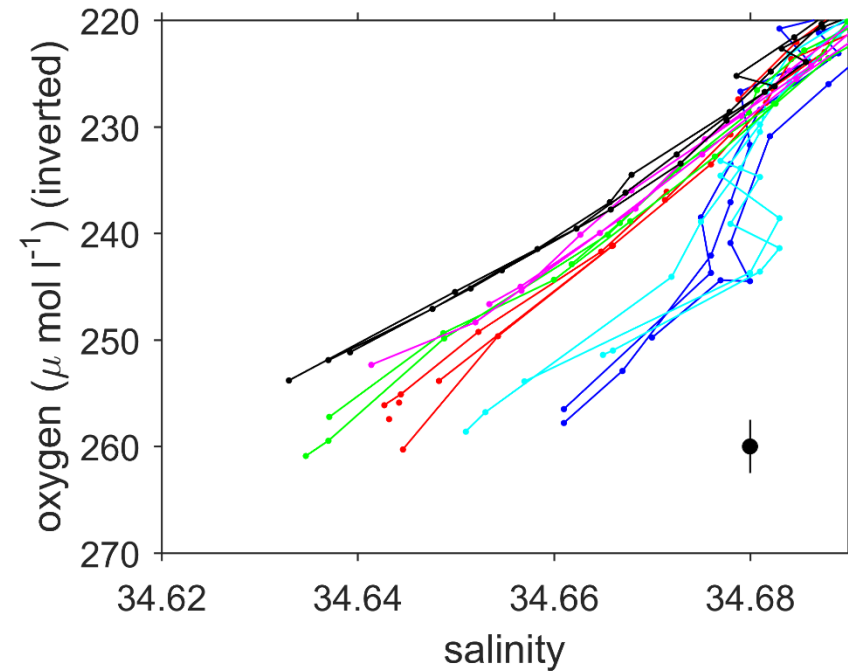
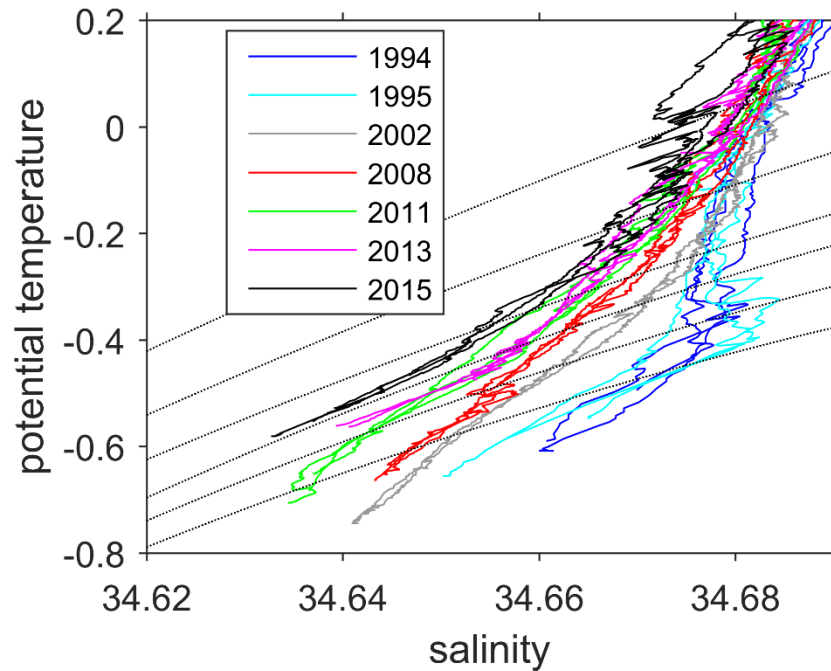


Snow et al., in prep

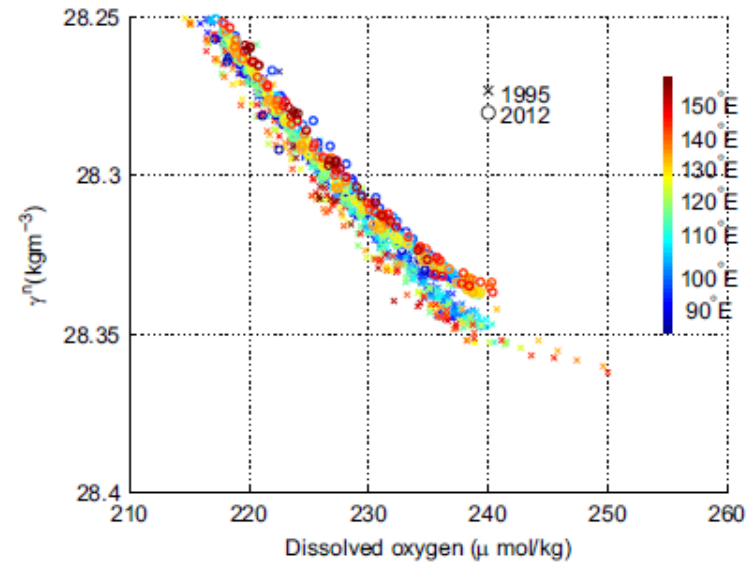
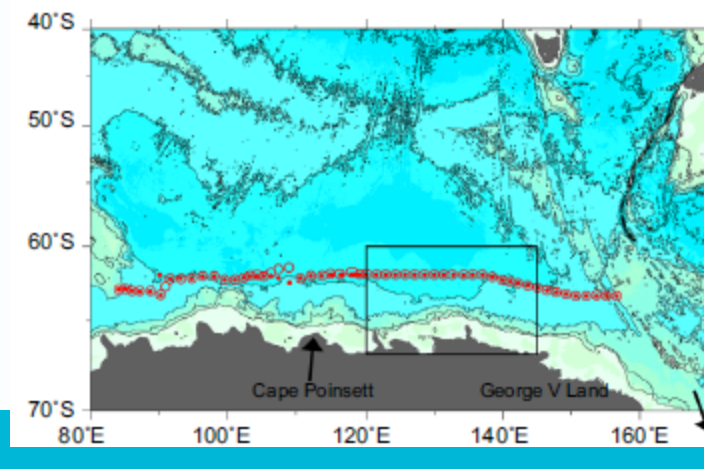
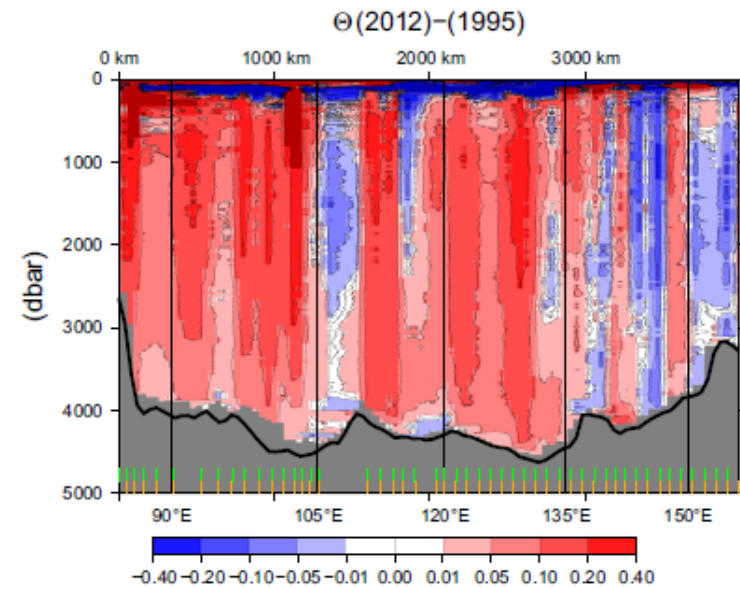
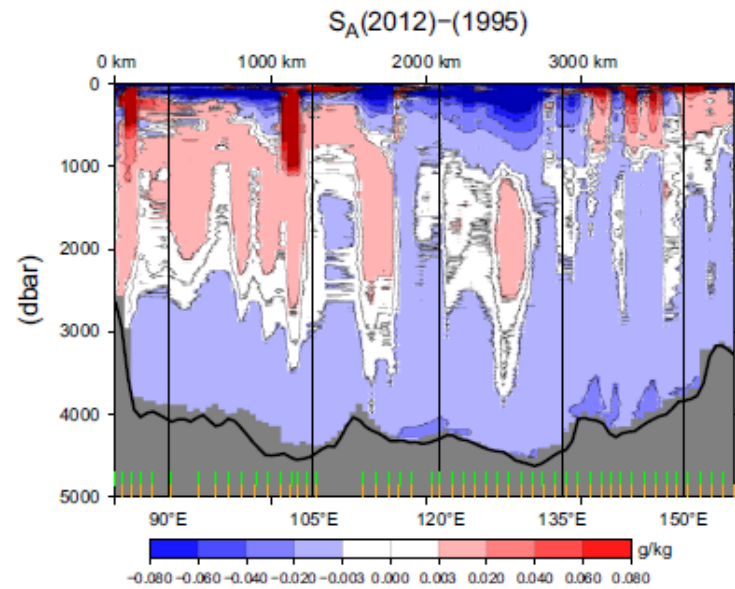
Sharp drop in density of bottom water after calving



Warmer, lighter AABW after calving, small decrease in oxygen



Change on S4 (~63S), 1995 to 2012

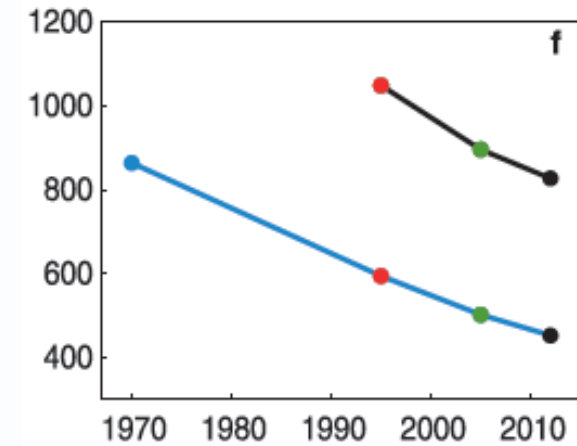
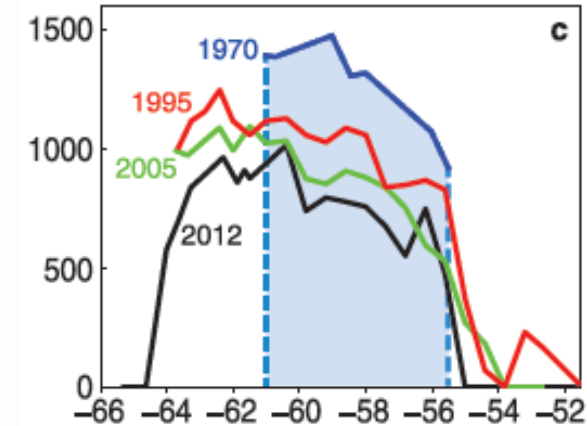
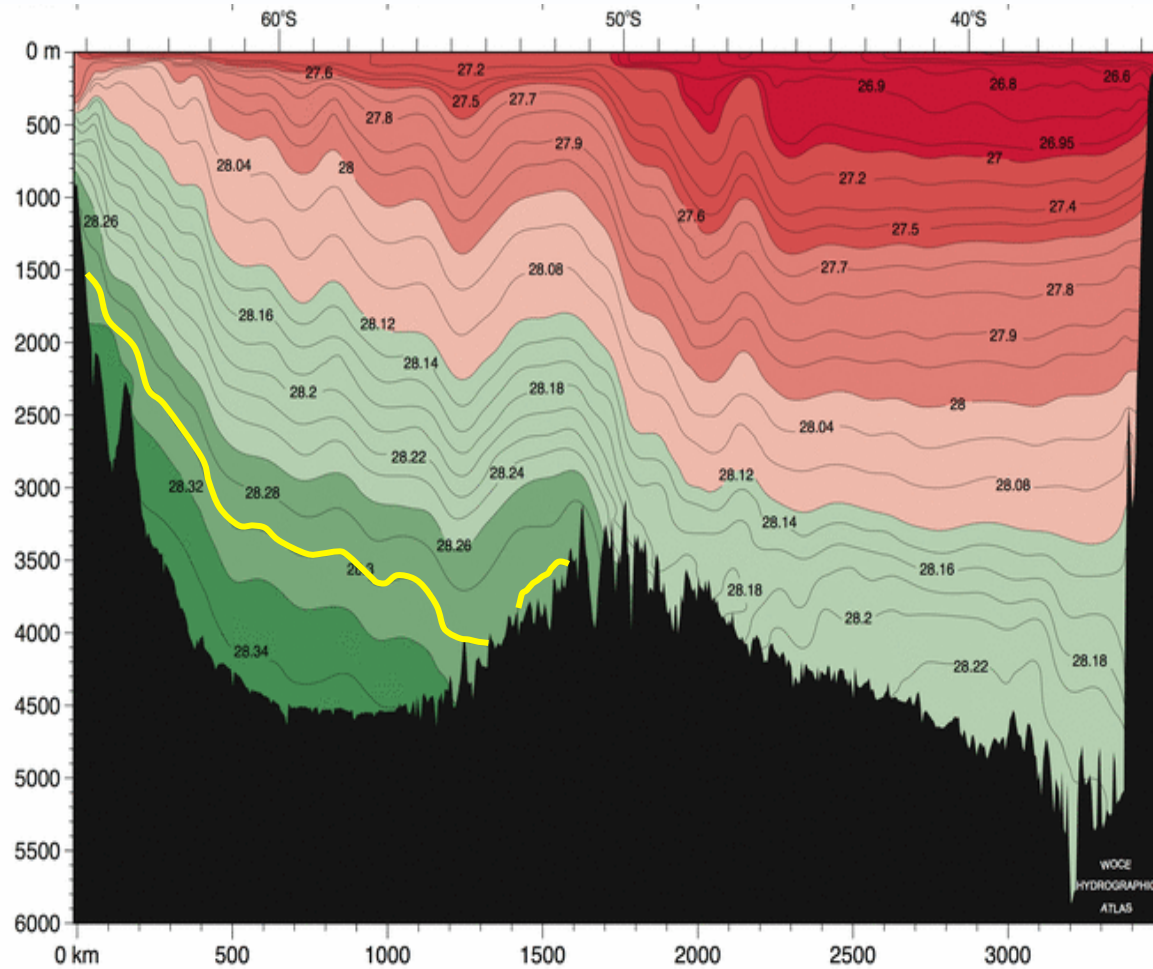




A revolution in ocean observing



Antarctic Bottom Water layer contracted by 50% between 1970 and 2012





Much of the Antarctic Ice Sheet sits on bedrock far below sea level

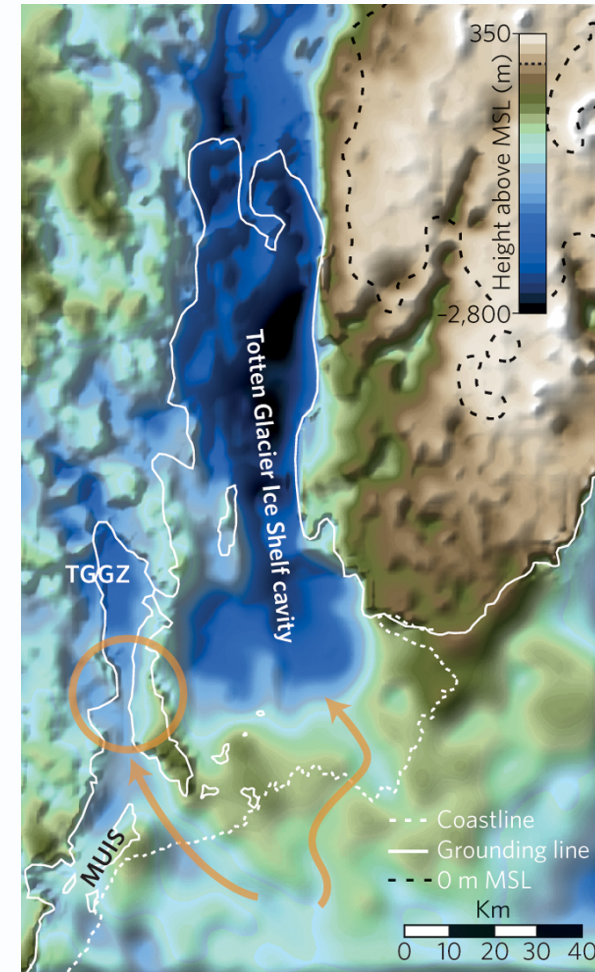


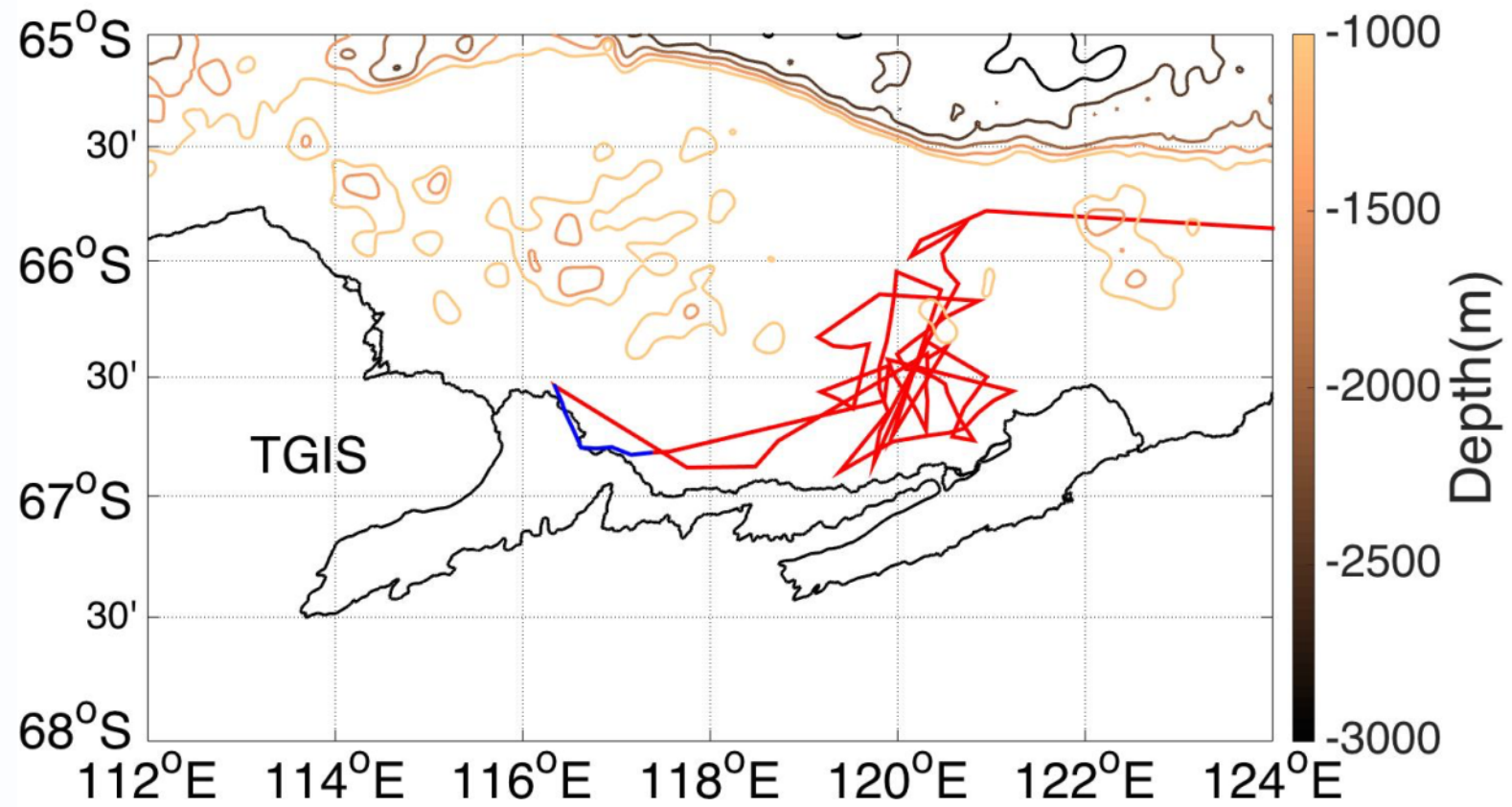
An expedition to the Totten Glacier



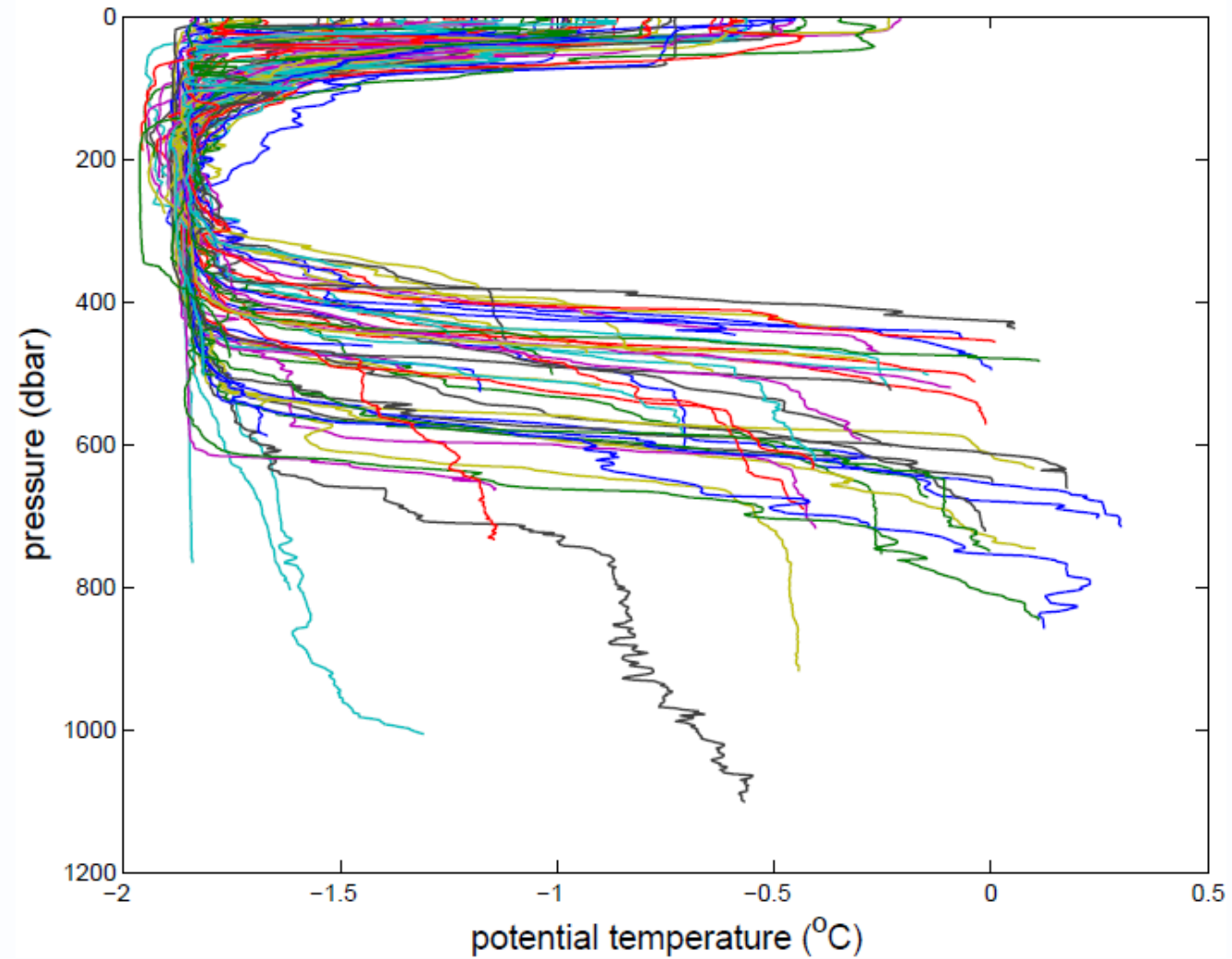
Why focus on the Totten Glacier?

- Ice volume equivalent to 3.9 m of global sea level rise
- Marine ice sheet - potentially unstable
- Likely contributed to past sea level rise
- Thinning rapidly, but we don't know why
- No oceanographic data near the glacier

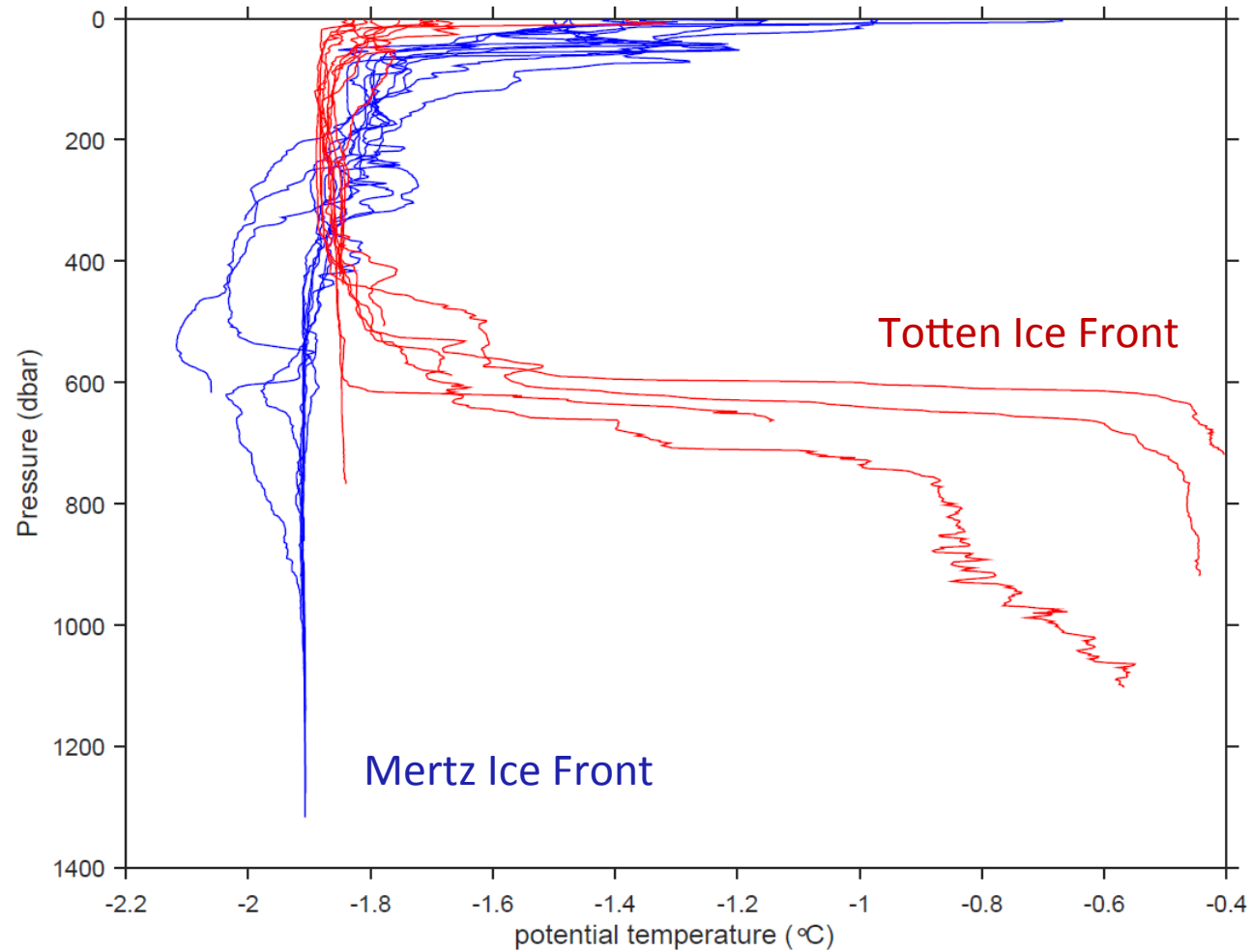




Lots of warm water on the Totten shelf



Not all East Antarctic glaciers are alike

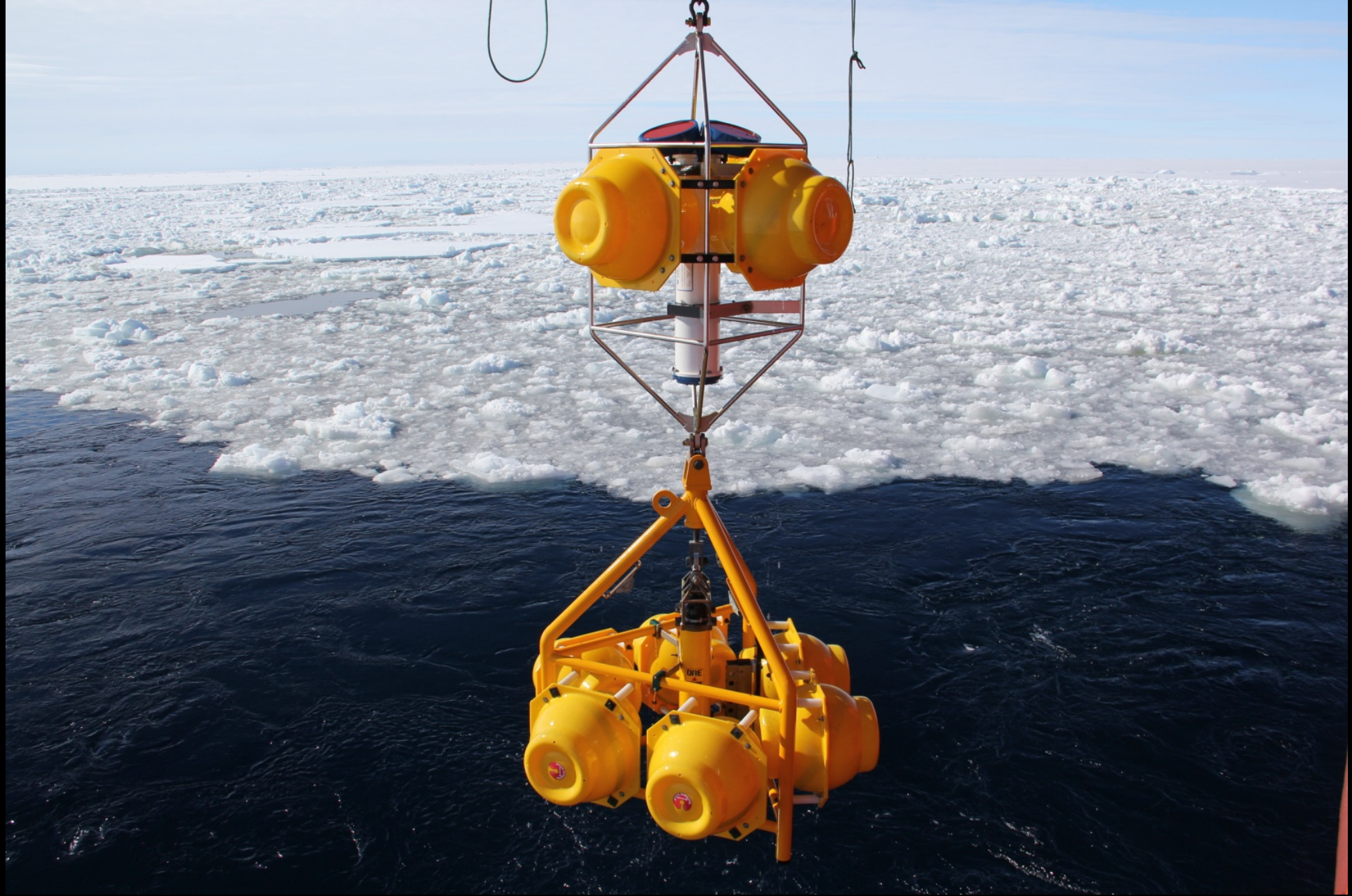




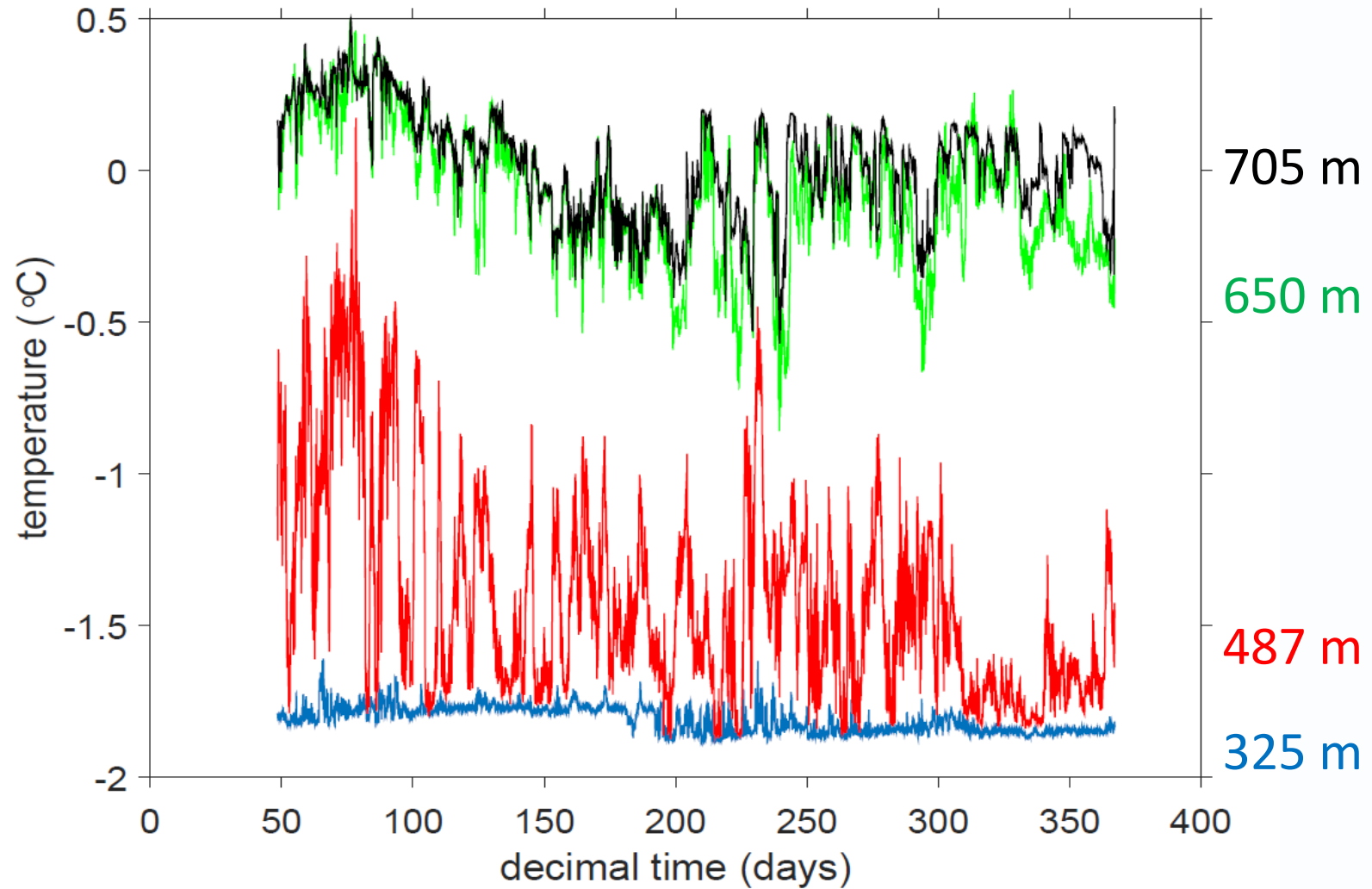
mooring







Warm water present at depth year-round





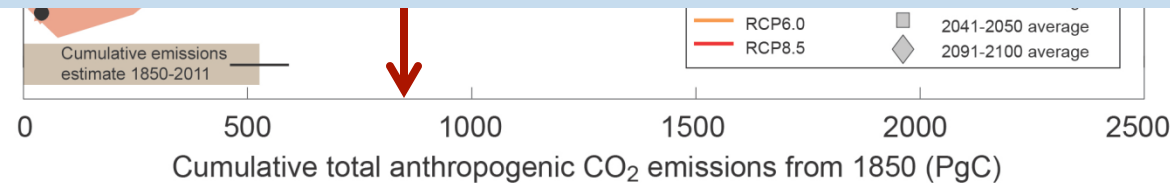
Time is running out to hit the 2°C target

We have a total of 800 billion tonnes of carbon to spend, if we want a 66% chance of keeping warming $< 2^{\circ}\text{C}$.

We have already used 550 billion tonnes.

We emit about 10 billion tonnes each year.

We run out of “safe” emissions in 25 years if we keep emitting carbon at the present rate.



Summary

The oceans have a profound influence on climate.

The Southern Ocean is particularly important – and is changing.

The fate of the Antarctic Ice Sheet lies in the ocean.

The East Antarctic Ice Sheet is likely more vulnerable to ocean change than we thought.

Therefore sea level will likely rise more than we expected.

Continued high emissions commit us to future change that we cannot reverse.



Thanks to:

Captain & crew of Aurora Australis

Students and colleagues

Antarctic Climate and Ecosystems Cooperative Research Centre

ARC Special Research Initiative for the Antarctic Gateway Partnership

Integrated Marine Observing System

Australian Antarctic Division

CSIRO



Peter Schuller

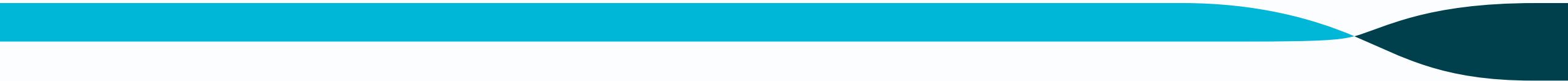
What if we don't hit the 2°C target?

Larger and more rapid change

Increased likelihood of crossing “tipping points”

For sea level, warming greater than 2°C results in an irreversible commitment to >10m of sea level rise

How long it takes to realise this commitment depends on how warm it gets (ie how much carbon we emit)



Thanks to:

- Captain + crew of Aurora Australis
- Students and colleagues
- Antarctic Climate and Ecosystems CRC
- Australian Research Council Special Research Initiative for the Antarctic Gateway Partnership
- Integrated Marine Observing System
- Australian Antarctic Division
- CSIRO

Peter Schuller



Open questions:

Has ocean heat flux to the Totten Glacier changed and caused thinning observed by satellite?

Why does so much warm water reach the Totten?

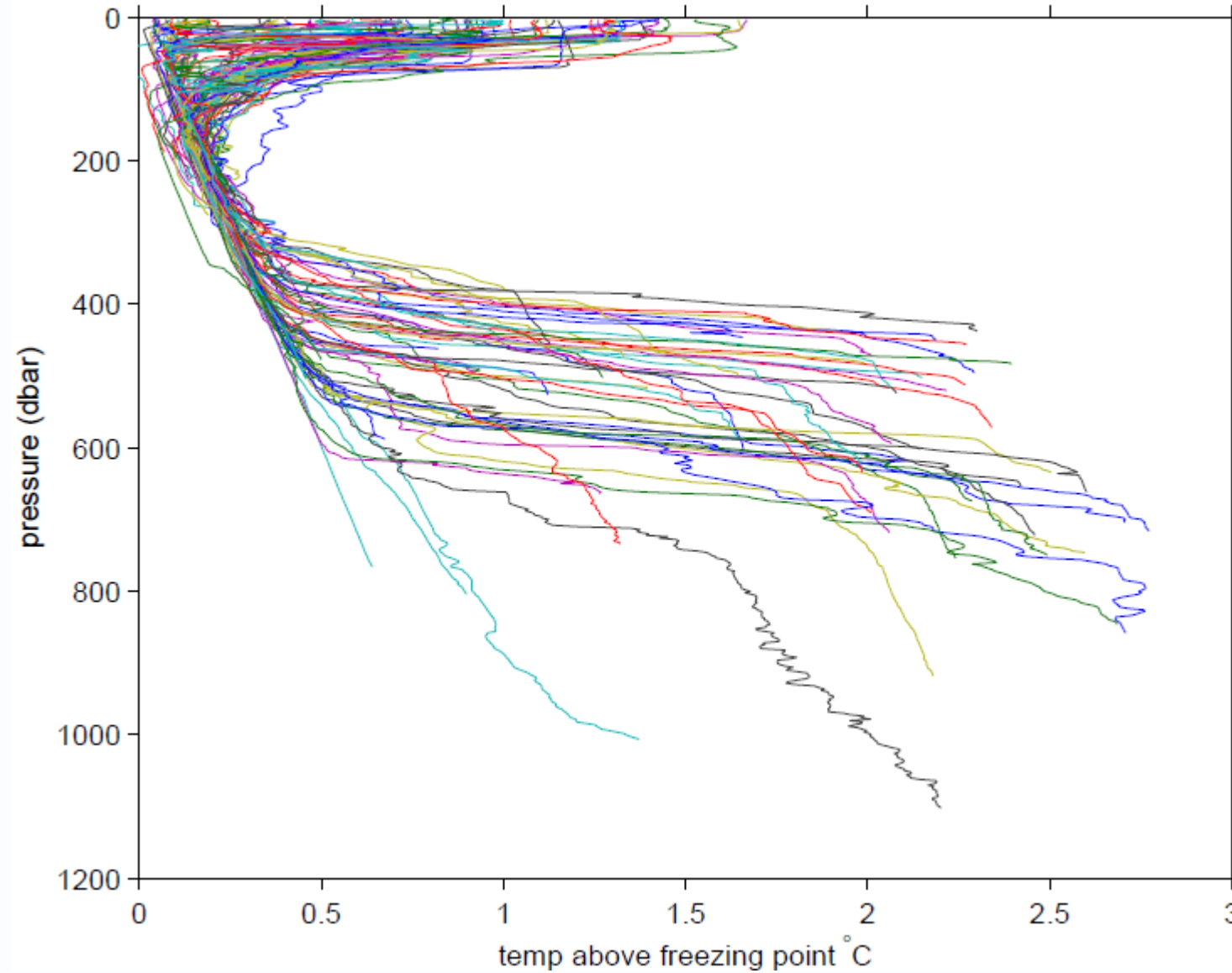
What controls the delivery of ocean heat to the ice shelf?

Does the deep trough observed at the ice front extend to the grounding line?

How stable is the Totten Glacier?

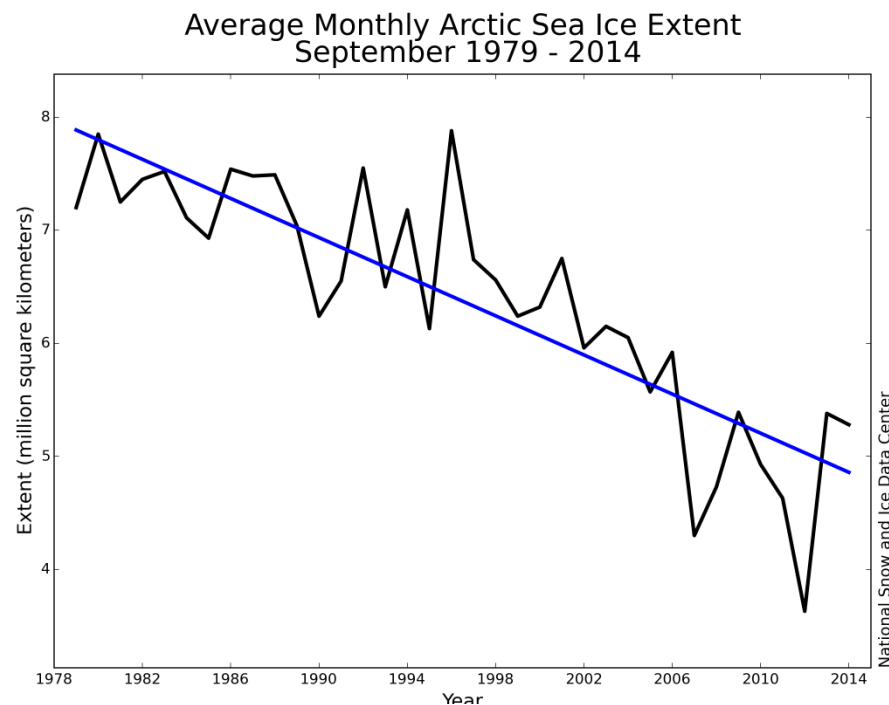
How much, and how fast, will the Totten Glacier contribute to global sea level rise?

Temperature above local pressure freezing point

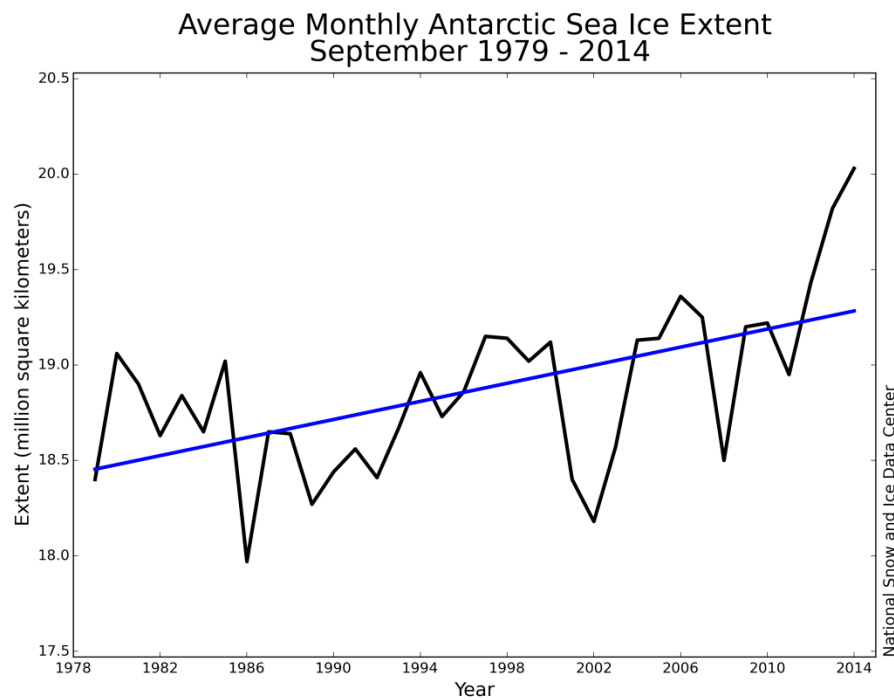




Doug Thost



Arctic: summer minimum sea ice extent is decreasing by 10% per decade



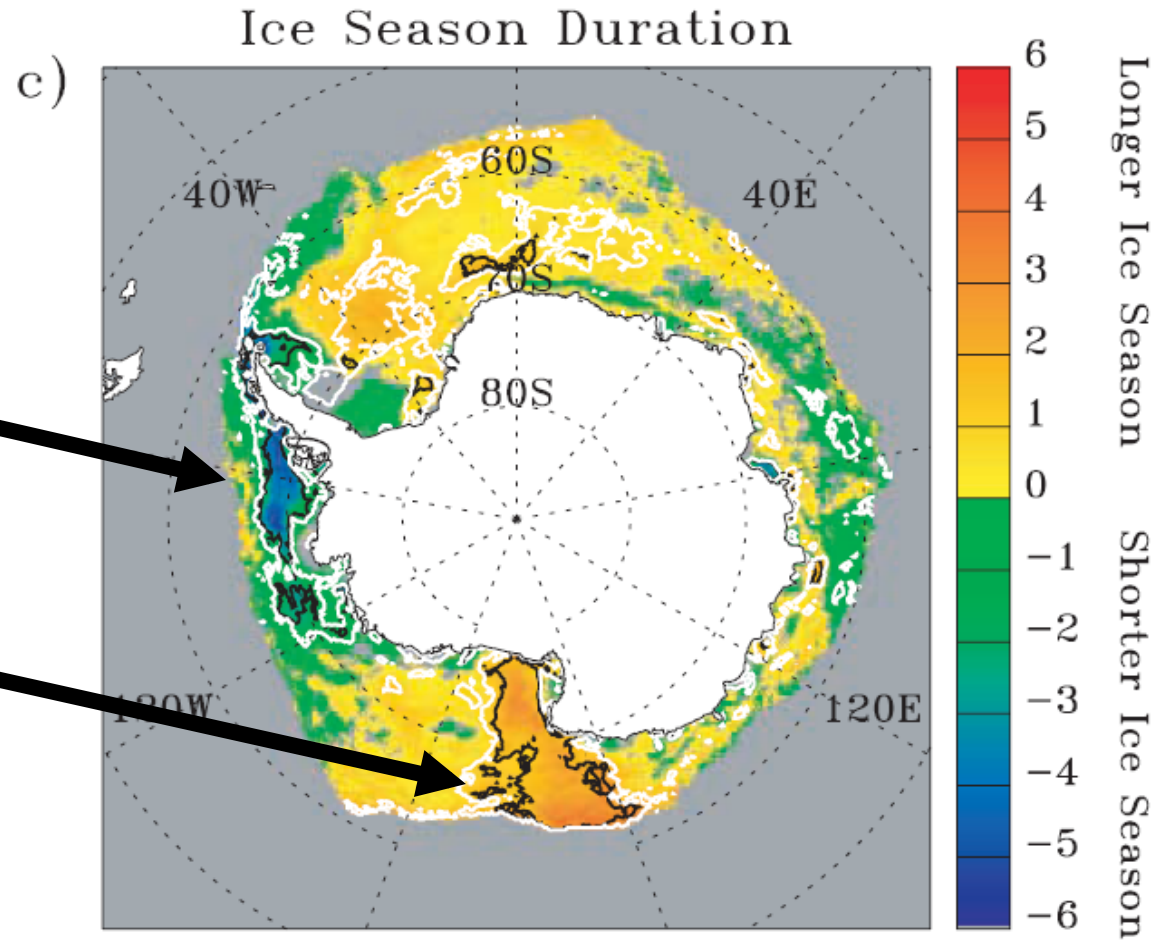
Antarctic: winter maximum sea ice extent is increasing by 1.3% per decade

Large regional changes in Antarctic sea ice

Changes in sea
ice duration:
1979 – 2006

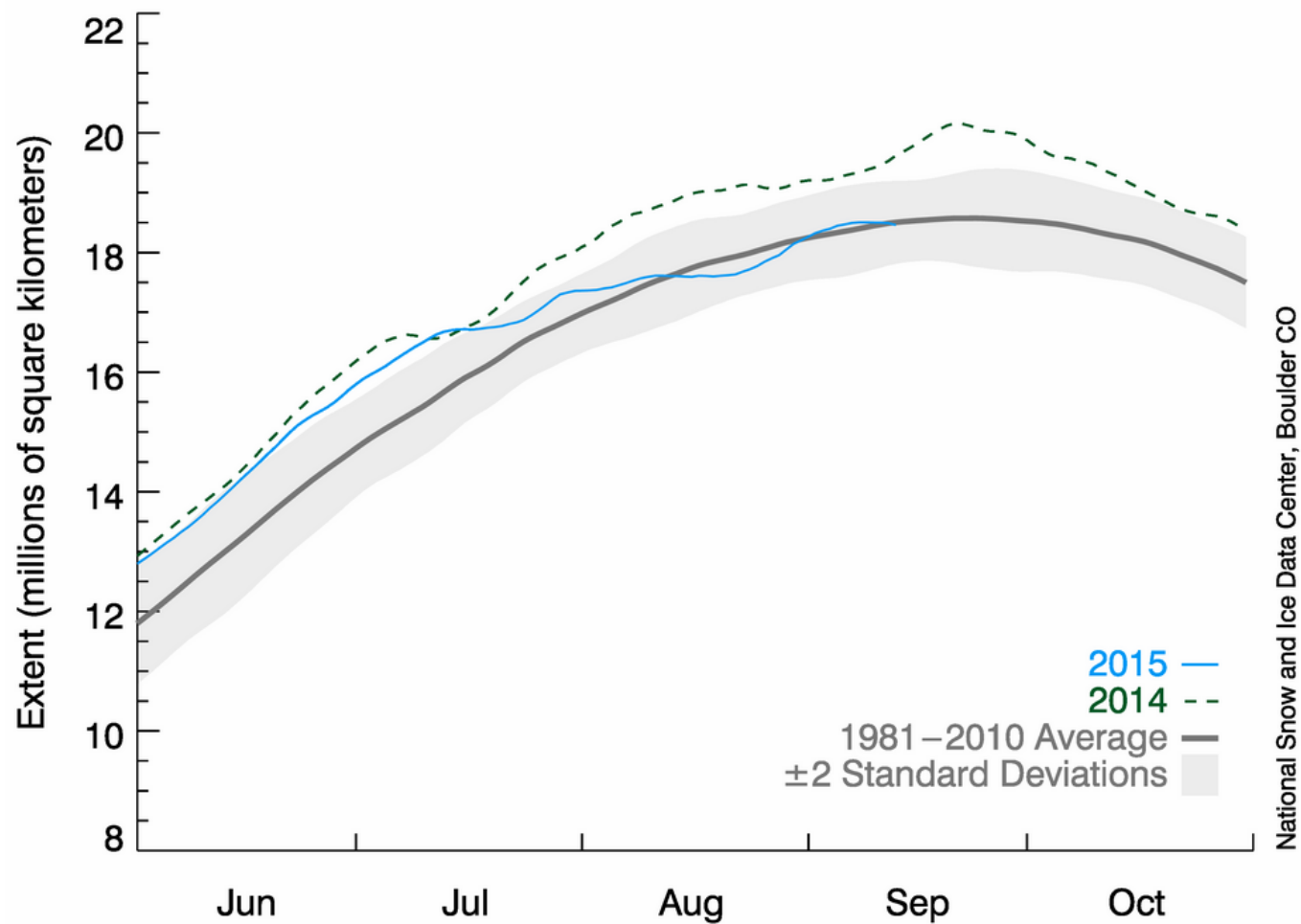
-83 ± 23 days

57 ± 13 days

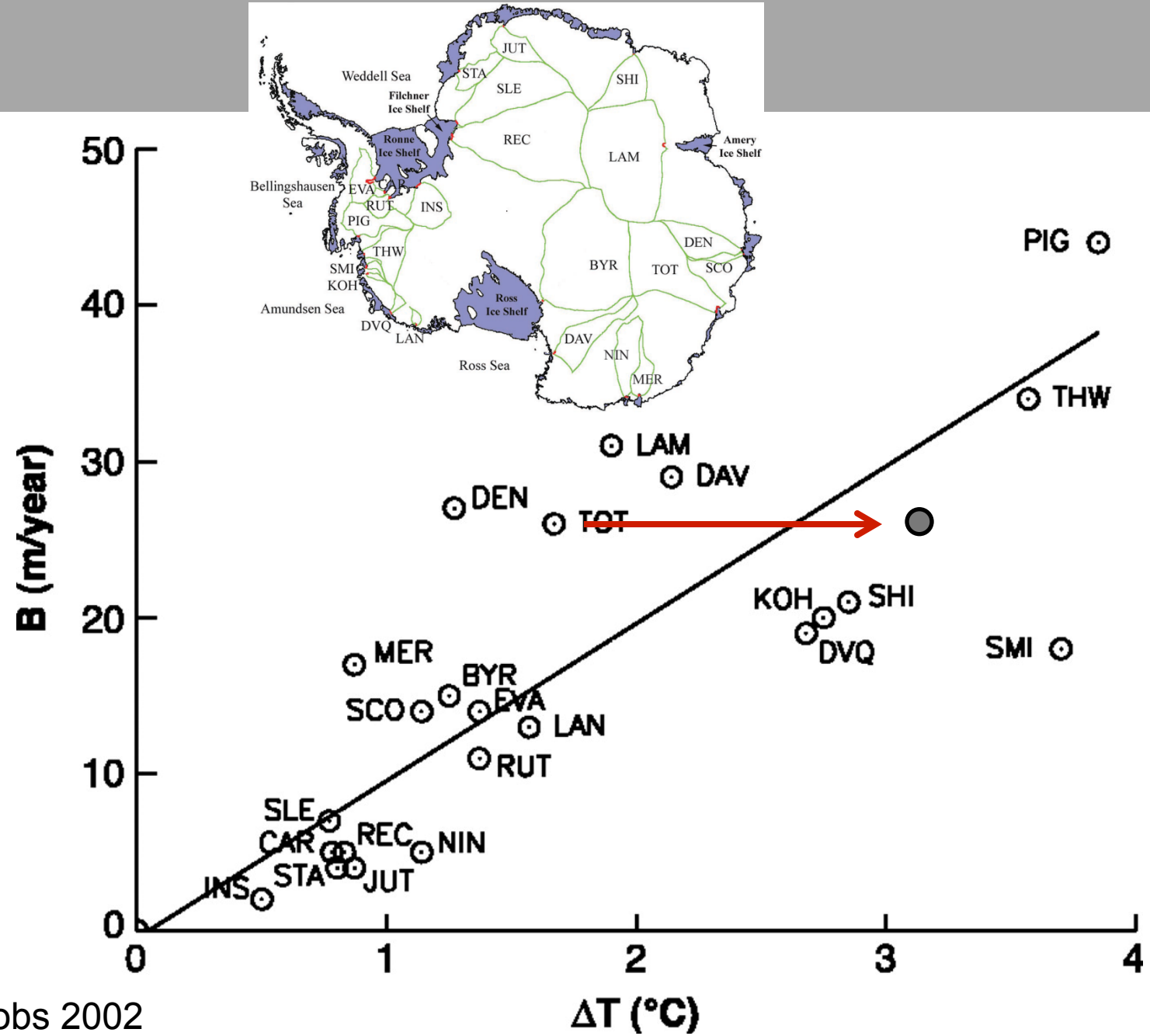


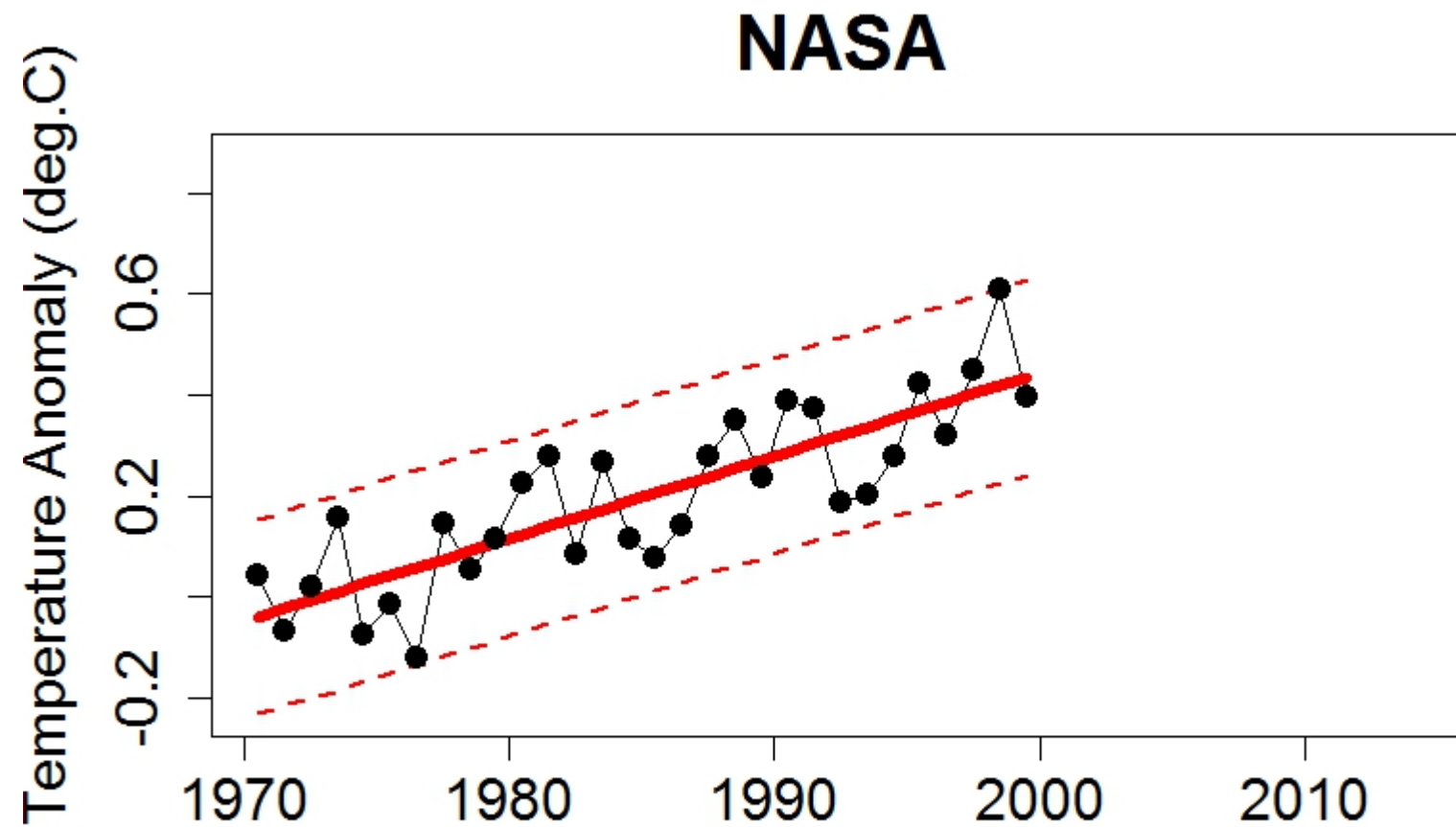
Stammerjohn et al. (2008)

Antarctic Sea Ice Extent (Area of ocean with at least 15% sea ice)



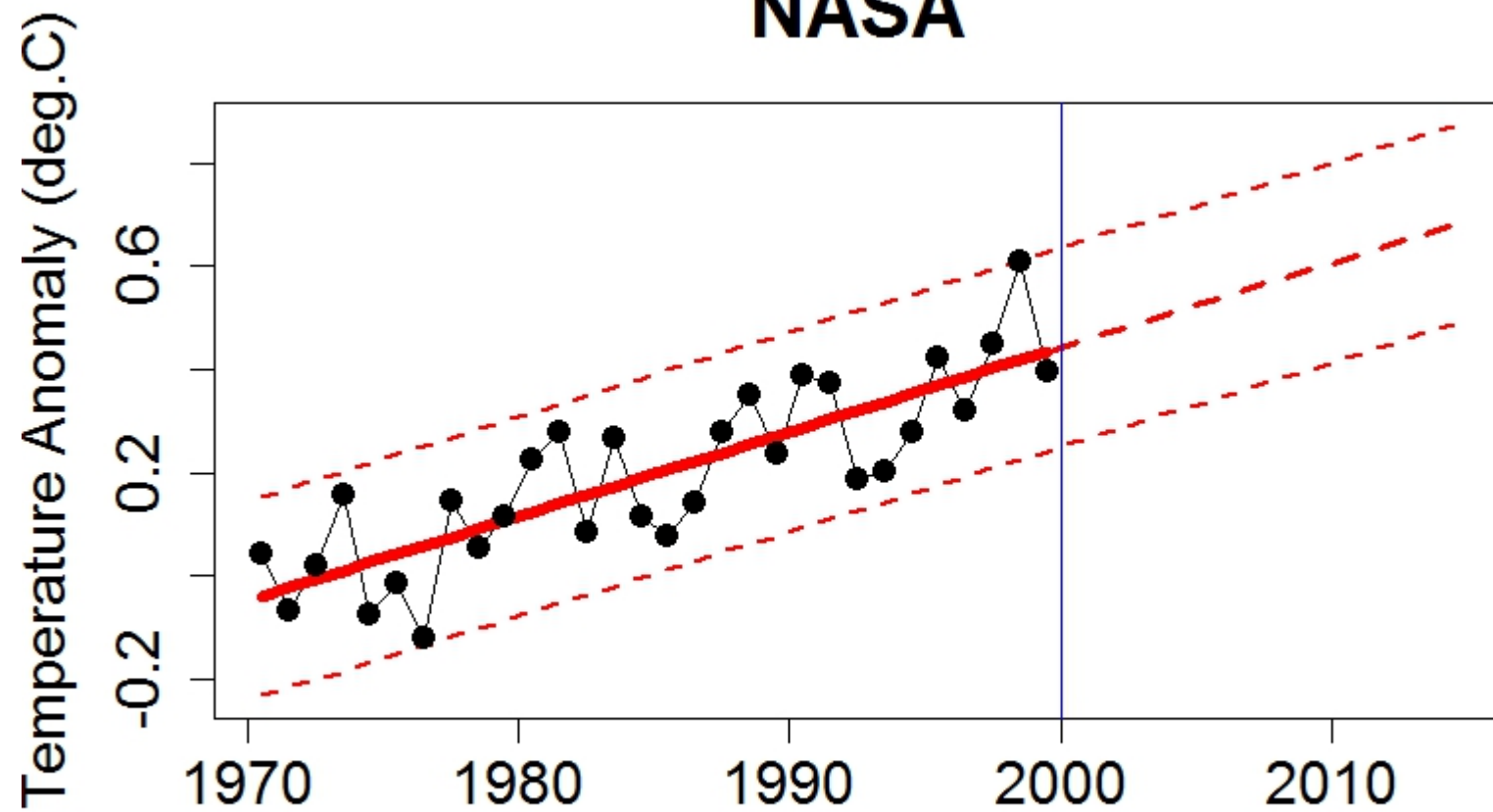
National Snow and Ice Data Center, Boulder CO



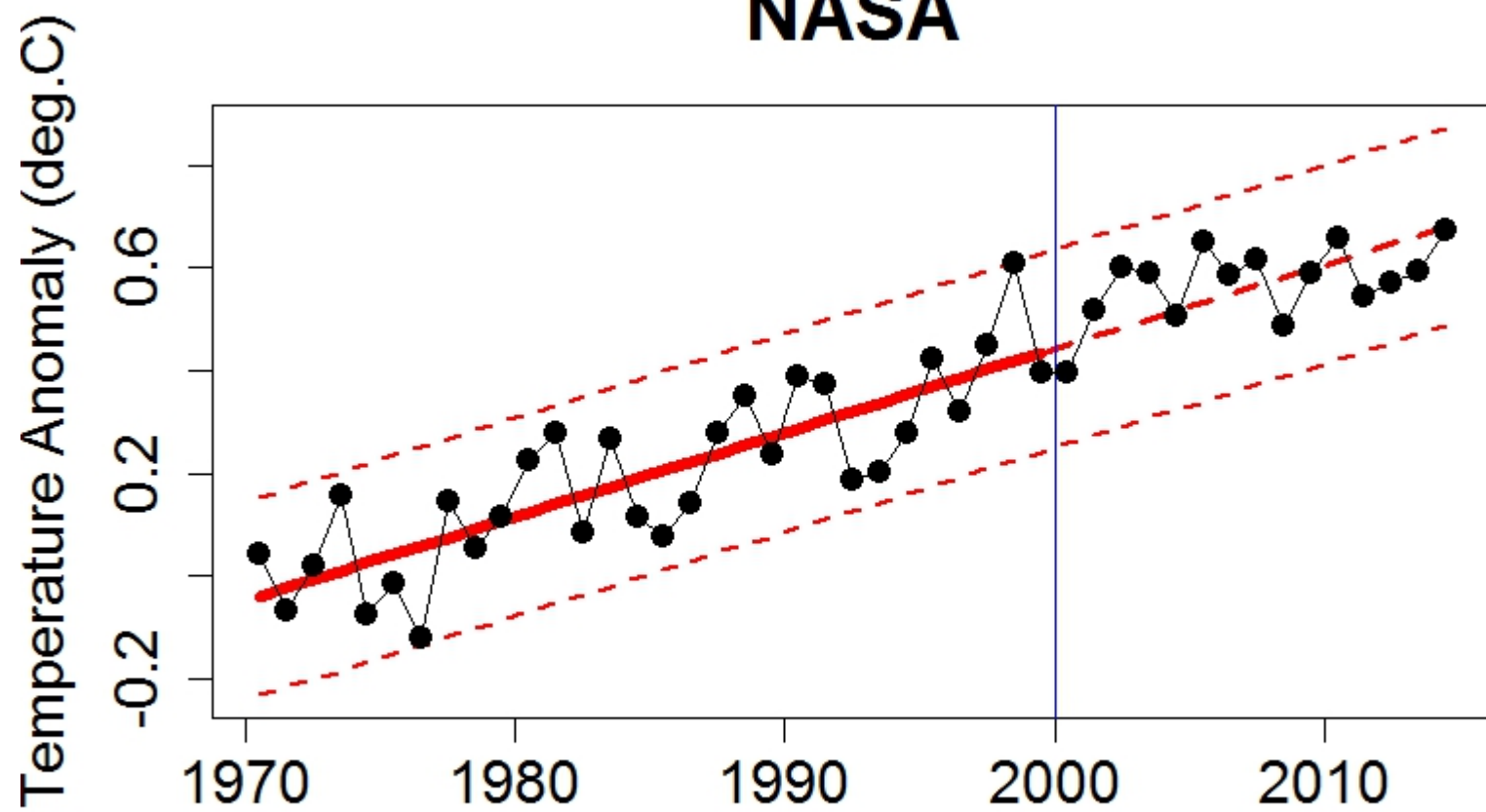


<https://tamino.wordpress.com/2015/01/20/its-the-trend-stupid-3/>

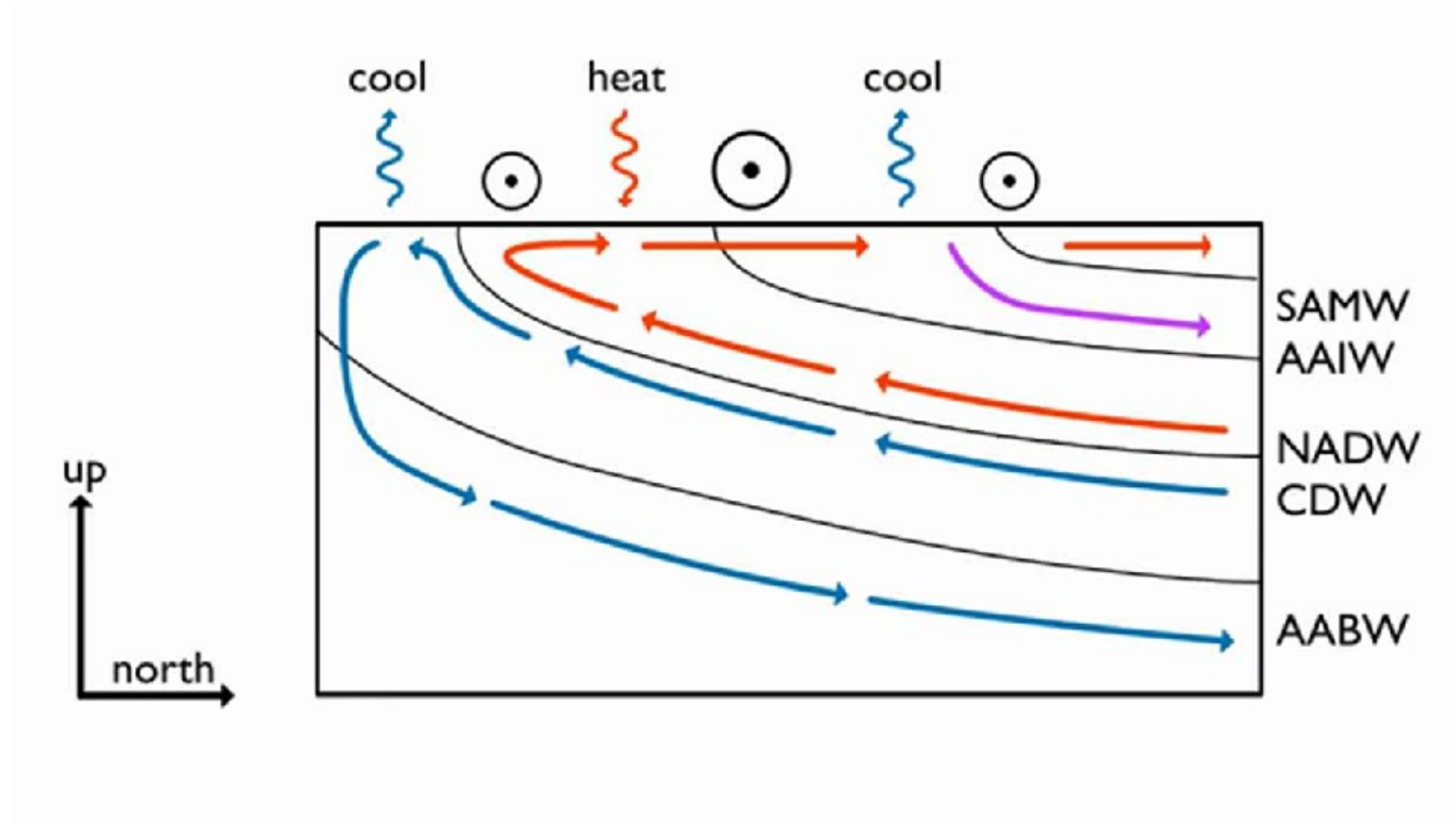
NASA



NASA



The overturning circulation and the Antarctic Circumpolar Current are linked

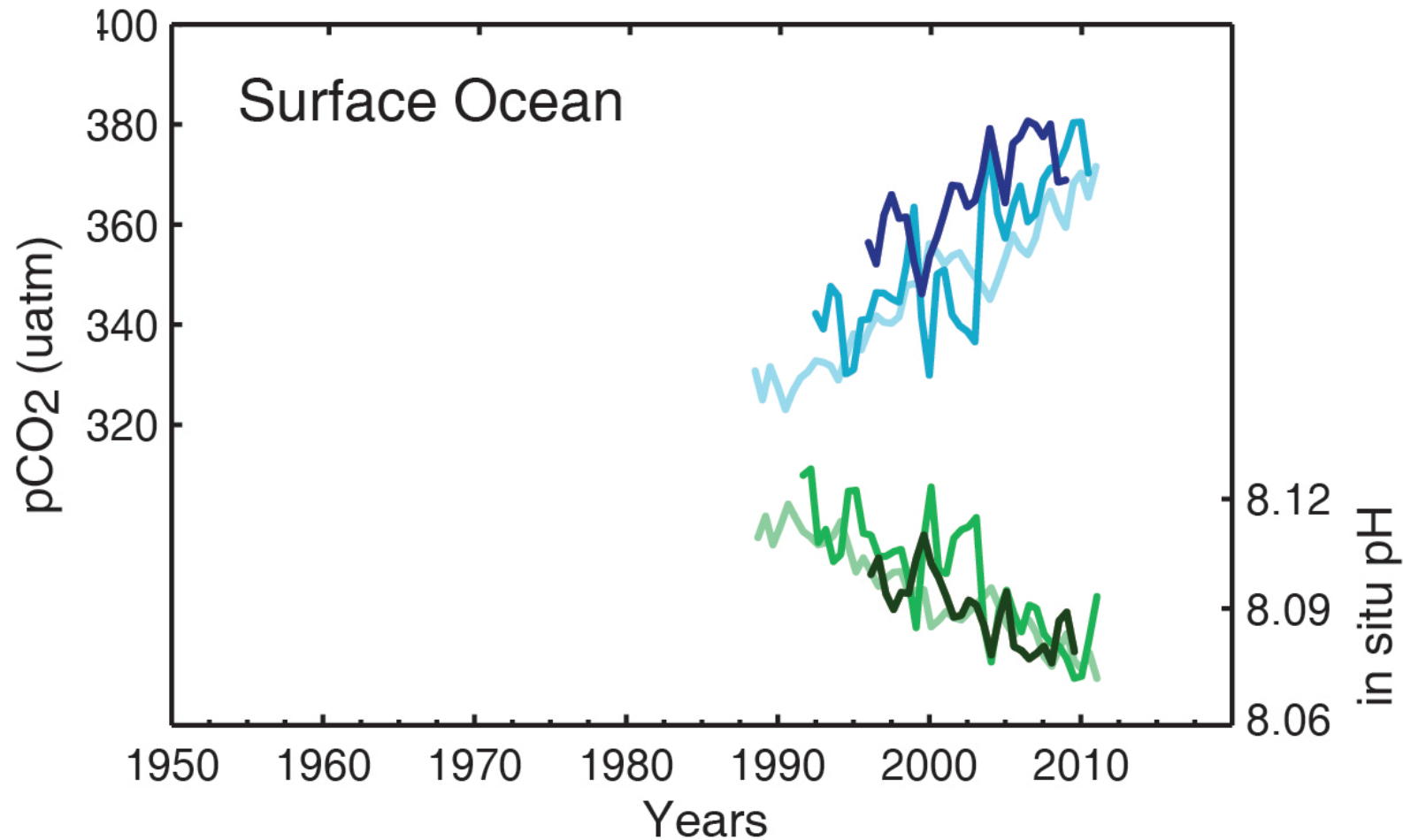


Abernathey et al., 2011; <https://vimeo.com/17822148>

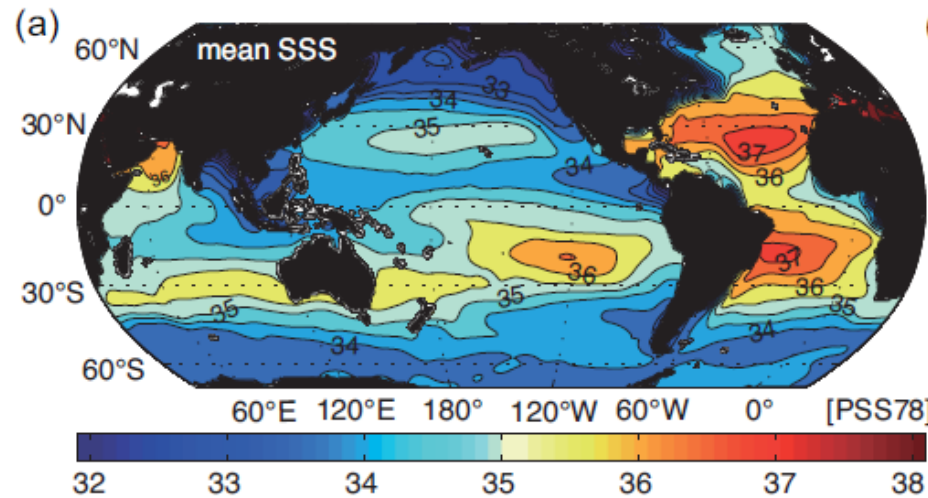
Headlines of IPCC 5th Assessment Report

- Warming is unequivocal. Many observed changes are unprecedented on timescales of decades or millenia.
- Human influence on climate is clear.
- Climate will continue to change in future.
- Limiting climate change will require substantial and sustained reductions in emissions.

Ocean uptake of carbon dioxide is acidifying the ocean



Changes in ocean salinity: the ocean “rain gauge”





1 Circumpolar Current =
454 billion stubbies/second

West Antarctic Ice Sheet

- Contains ice volume equivalent to 5 m of global sea level rise (3.3 m in marine-based ice sheet)
- Marine ice sheet – potentially unstable
- Has collapsed and re-formed many times in the past
- Thinning rapidly, at accelerating rate – marine ice sheet instability already underway? (Joughin et al., 2014; Rignot et al., 2014; Favier et al., 2014)
- Thinning result of basal melt of the floating ice shelves
- West Antarctica believed to be more vulnerable to basal melt because it is closer to warm waters of the ACC – the “soft underbelly of the Antarctic Ice Sheet” (Hughes, 1980)

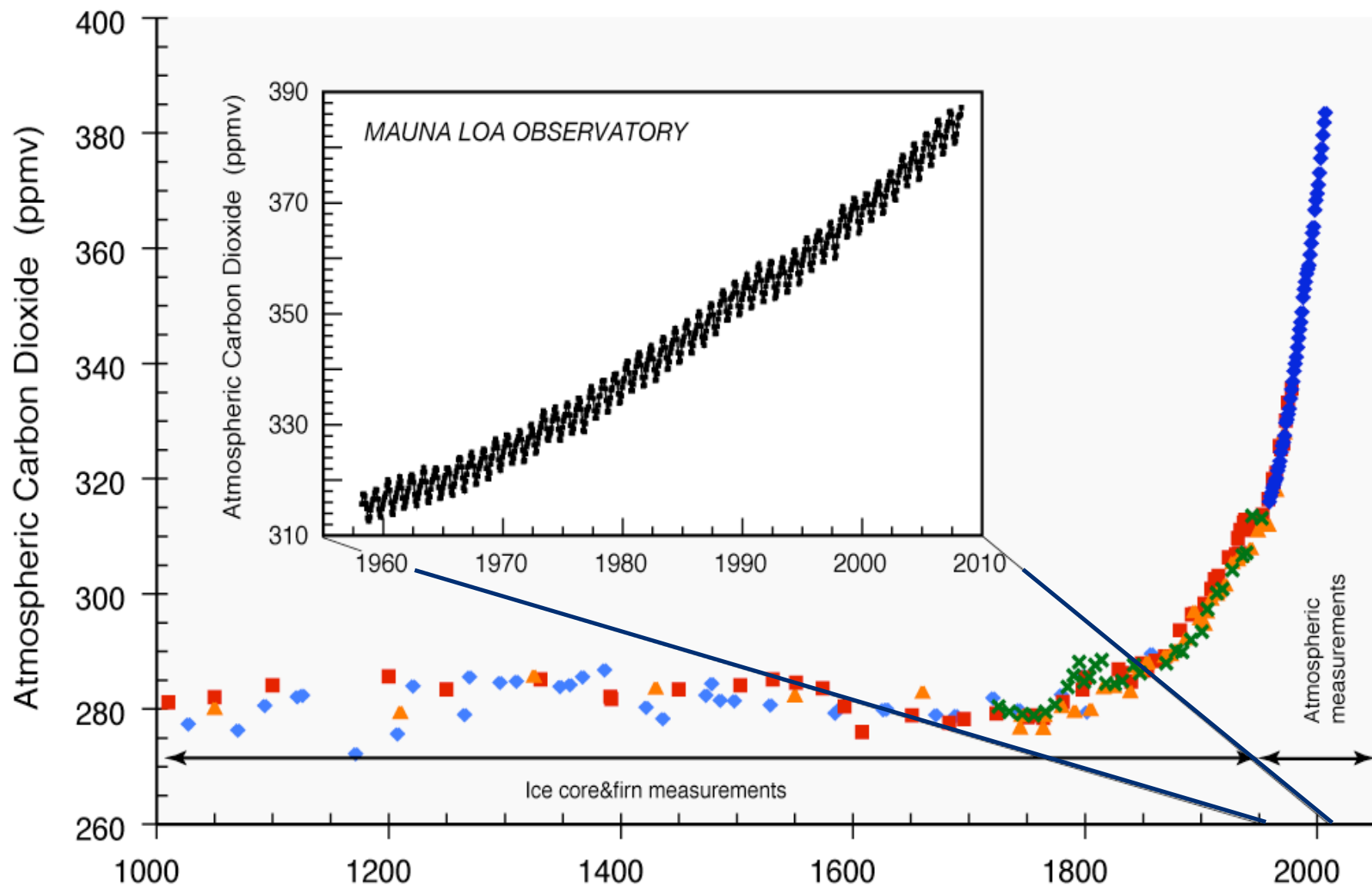
Climate Change 2013: The Physical Science Basis

Working Group I contribution to the IPCC Fifth Assessment Report

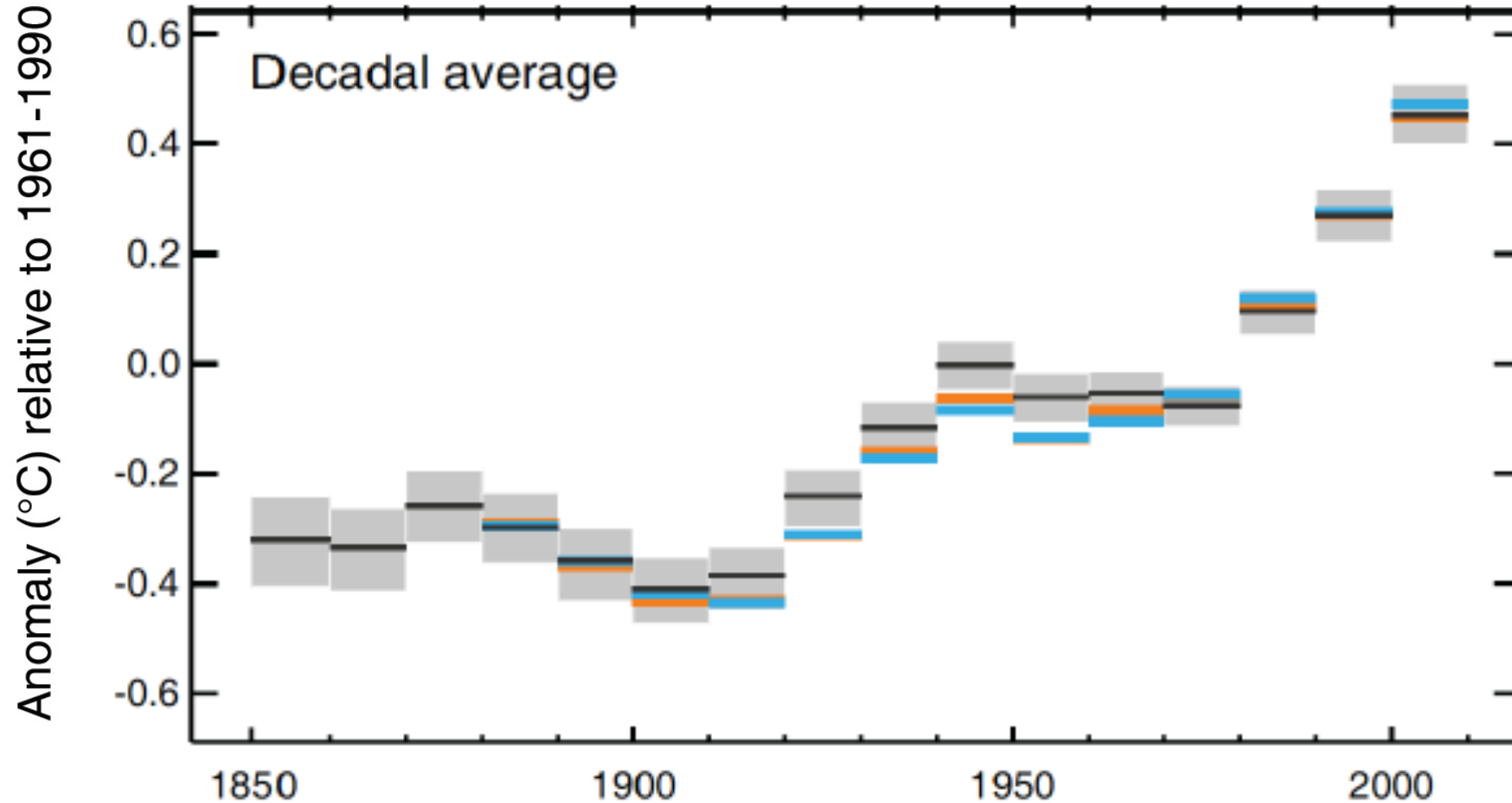


© Yann Arthus-Bertrand / Altitude

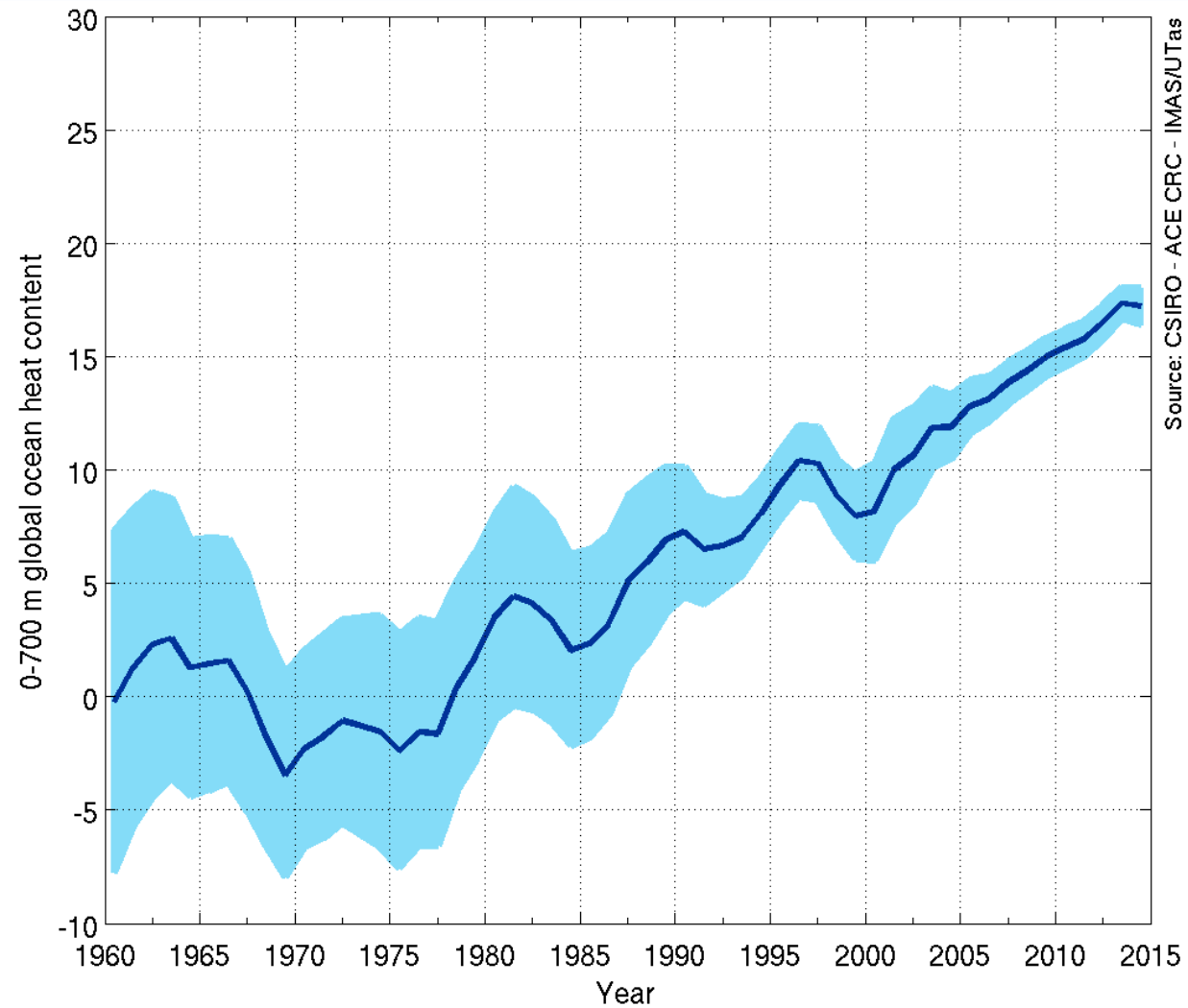
Atmospheric carbon dioxide has increased by 40%



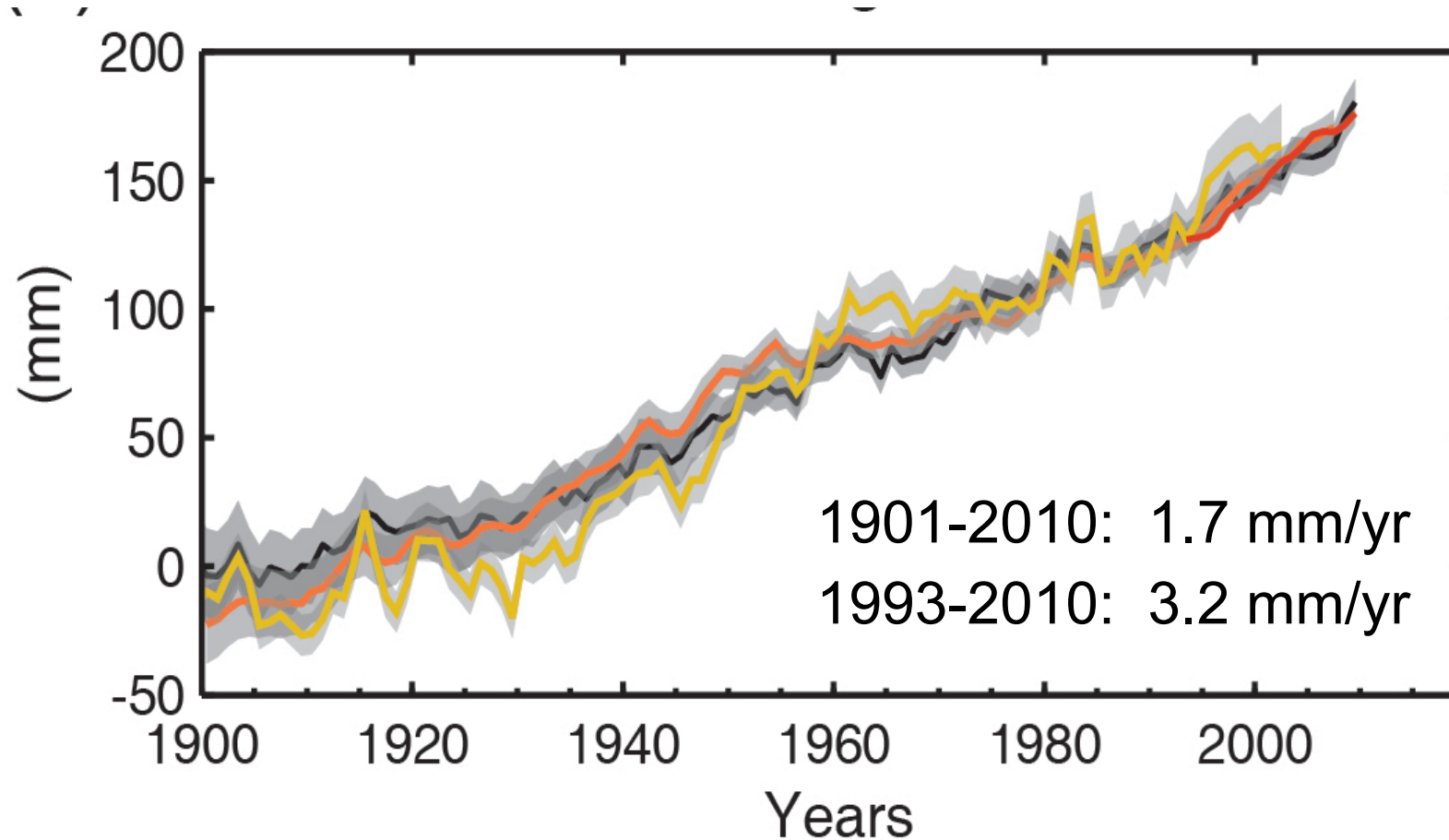
Each of the last three decades has been warmer than any previous decade since 1850.



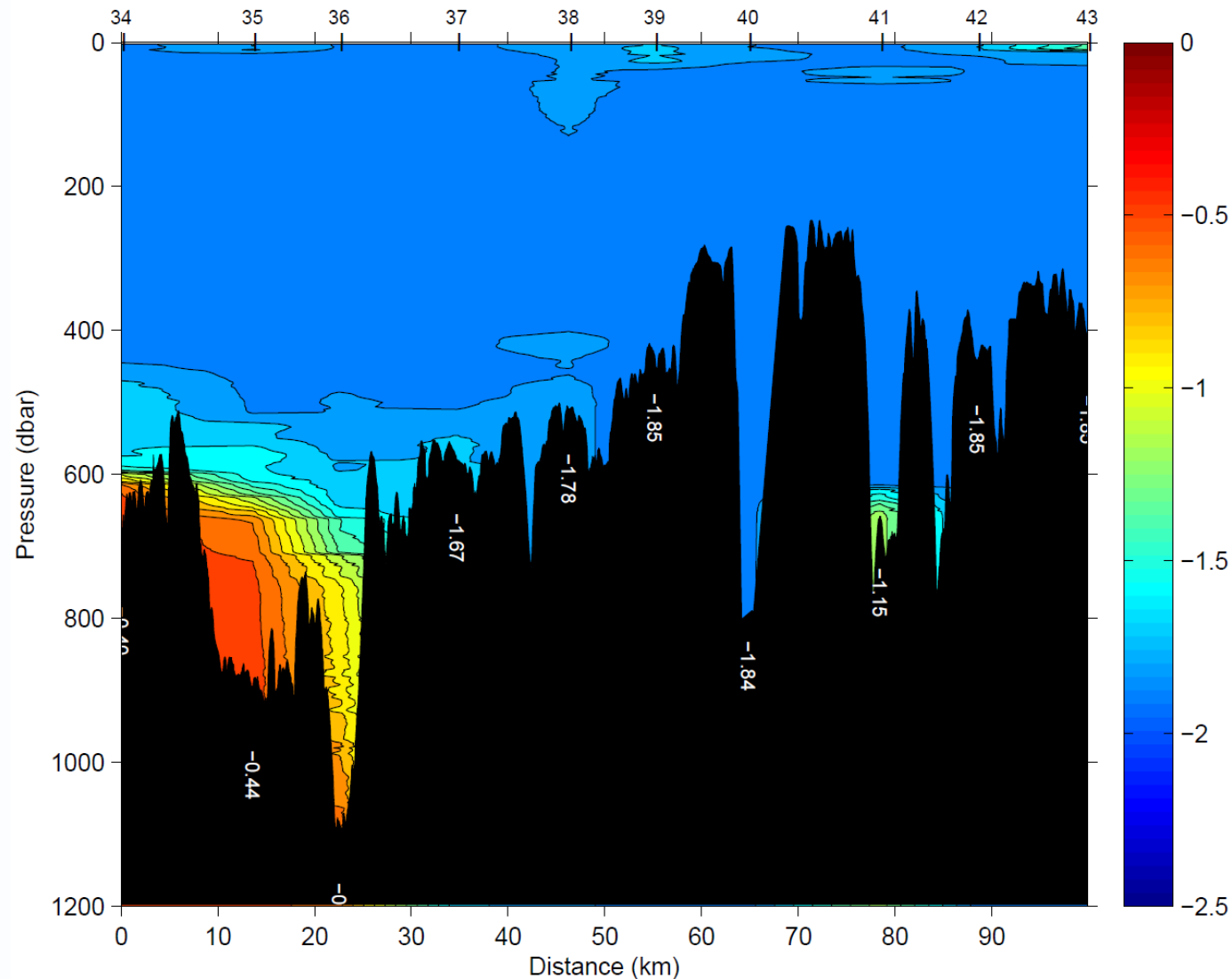
The ocean has warmed since 1970



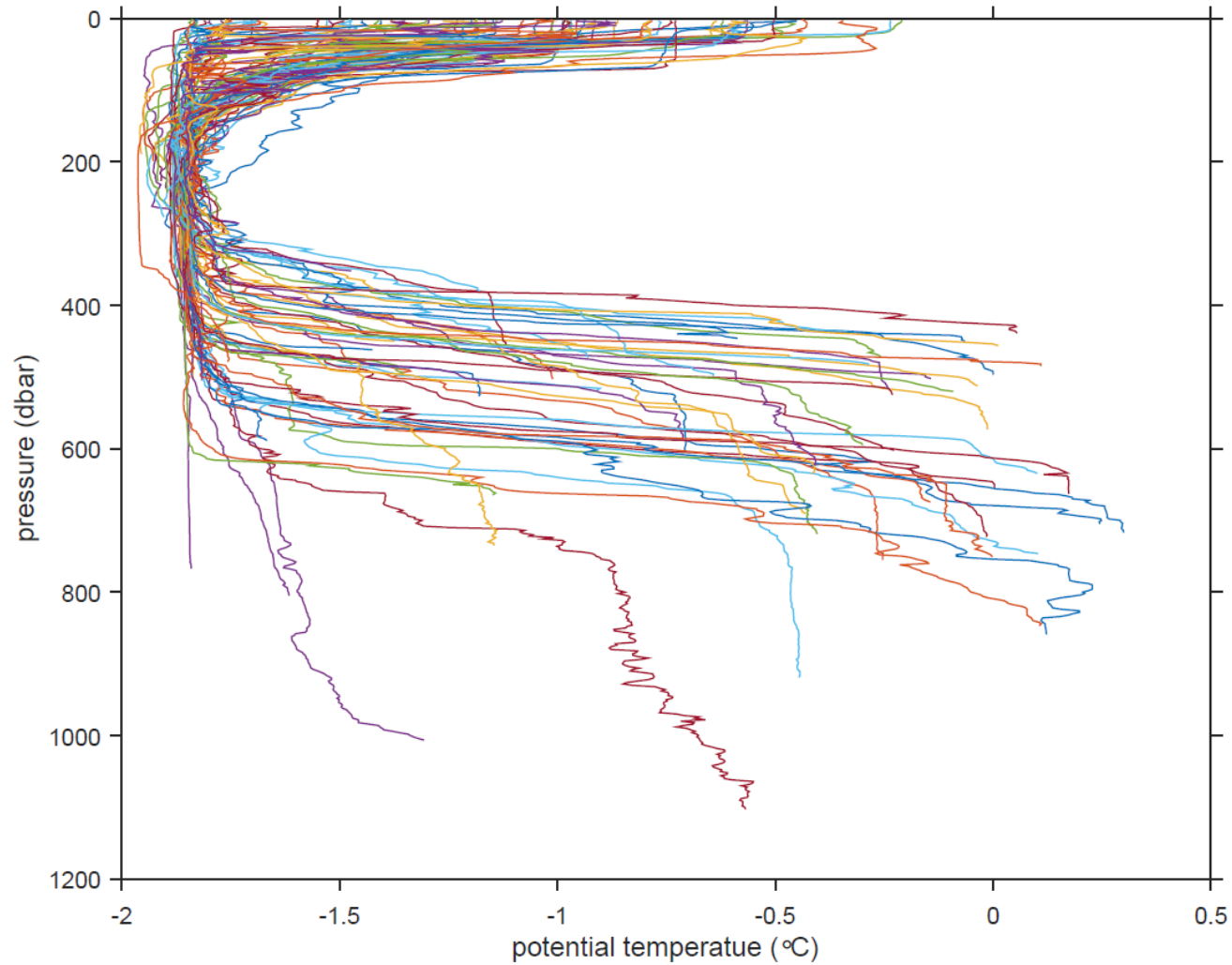
Global mean sea level increased by 0.19 m between 1901 and 2010



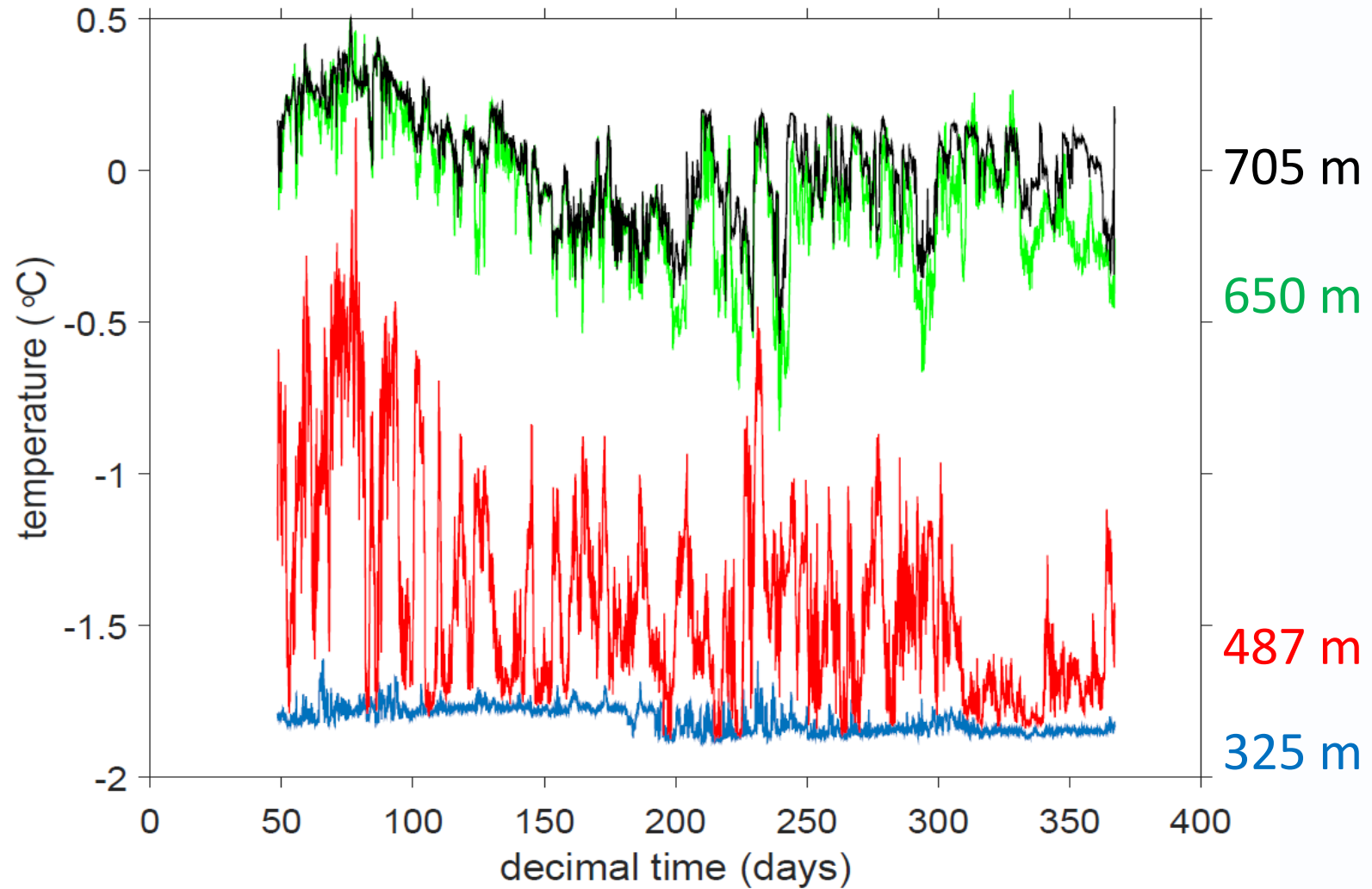
A deep channel at the Totten ice front allows warm water to enter ice-shelf cavity



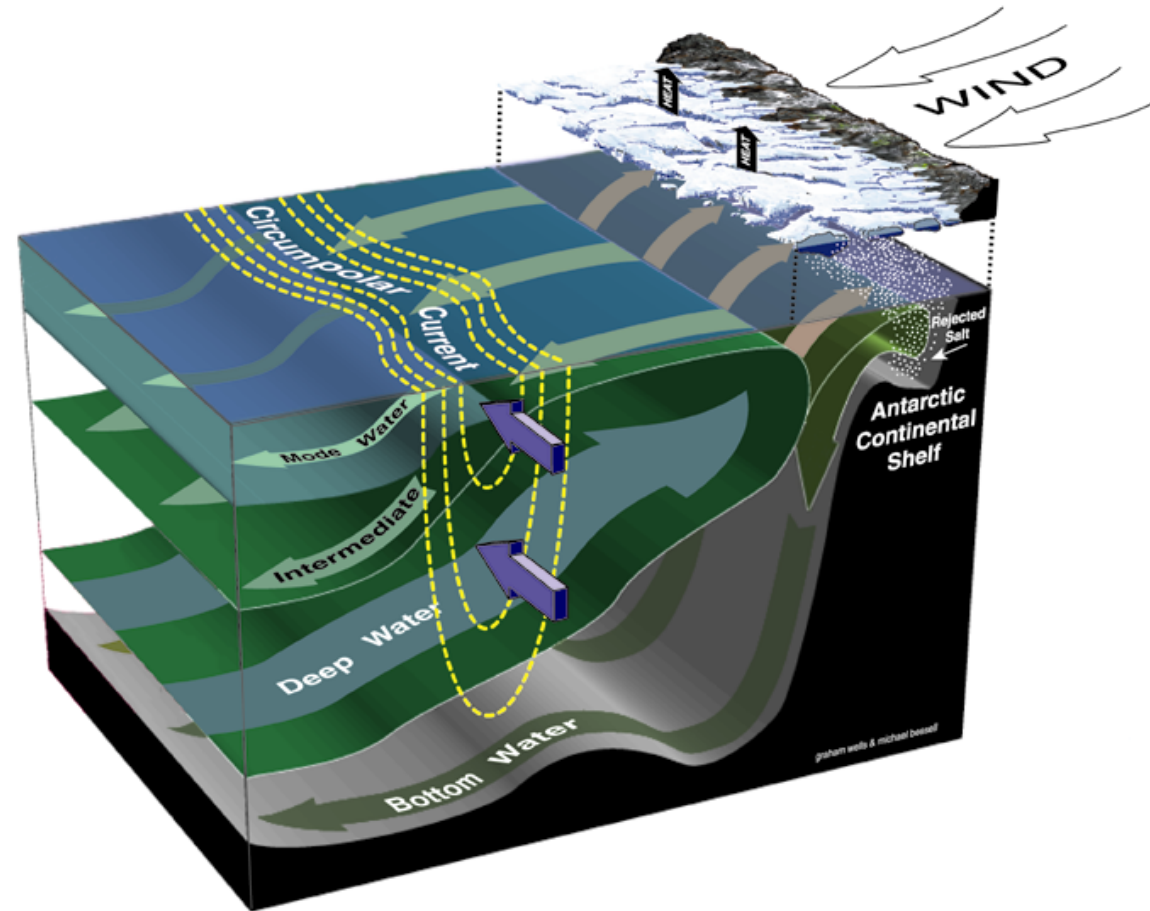
Lots of warm water on the Totten shelf



Warm water present at depth year-round



How will warming, freshening and changes in winds influence the overturning circulation?



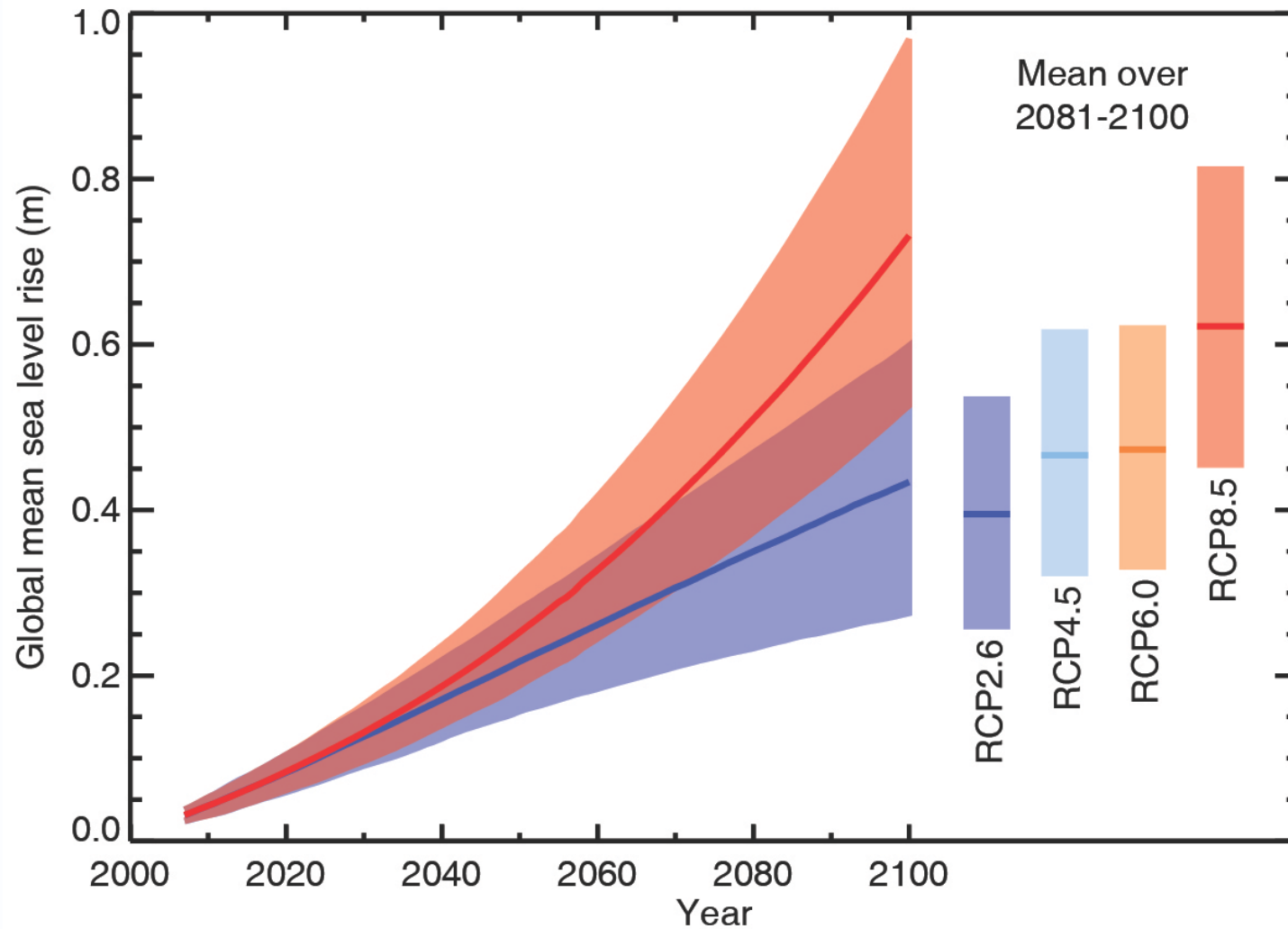
How will ocean acidification (and other stressors) affect marine ecosystems?



How will warming and changes in circulation affect ice shelves and sea level rise?



Future sea level rise

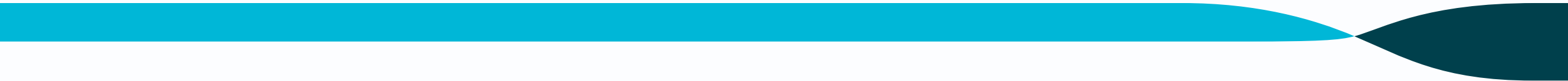


Sea level rise beyond 2100: IPCC

It is *virtually certain* that global mean sea level rise will continue for many centuries beyond 2100.

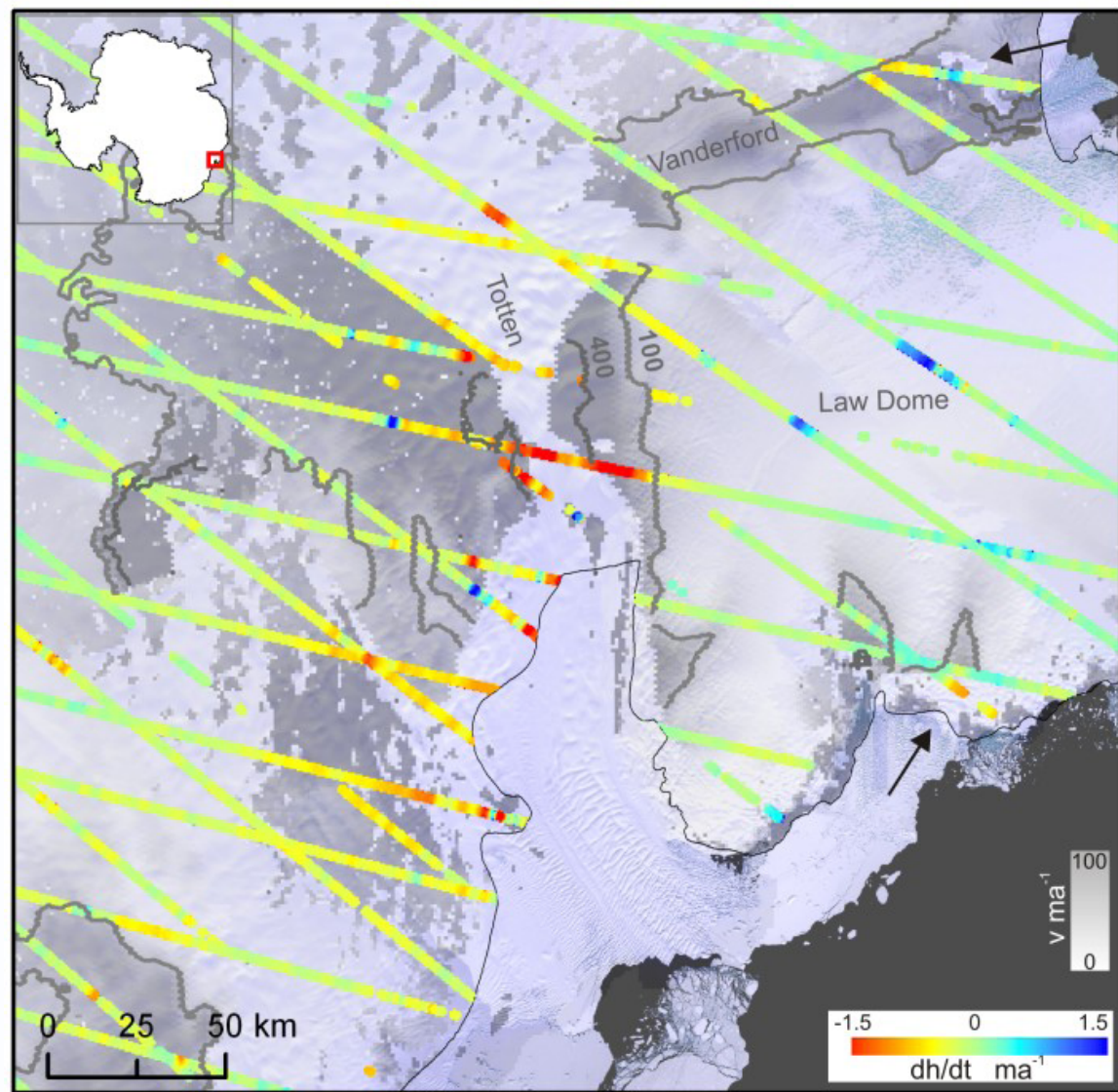
By 2300, there is *medium confidence* that GMSL rise will be less than 1 m if CO₂ concentrations stay below 500 ppm, but 1 to more than 3 m for 700–1500 ppm.

Larger sea level rise could result from sustained mass loss by ice sheets, and some part of the mass loss might be irreversible.



What does the future hold?





Potential temperature at Mertz Glacier

