

Ocean Interactions with the Antarctic Ice Shelves In East Antarctica

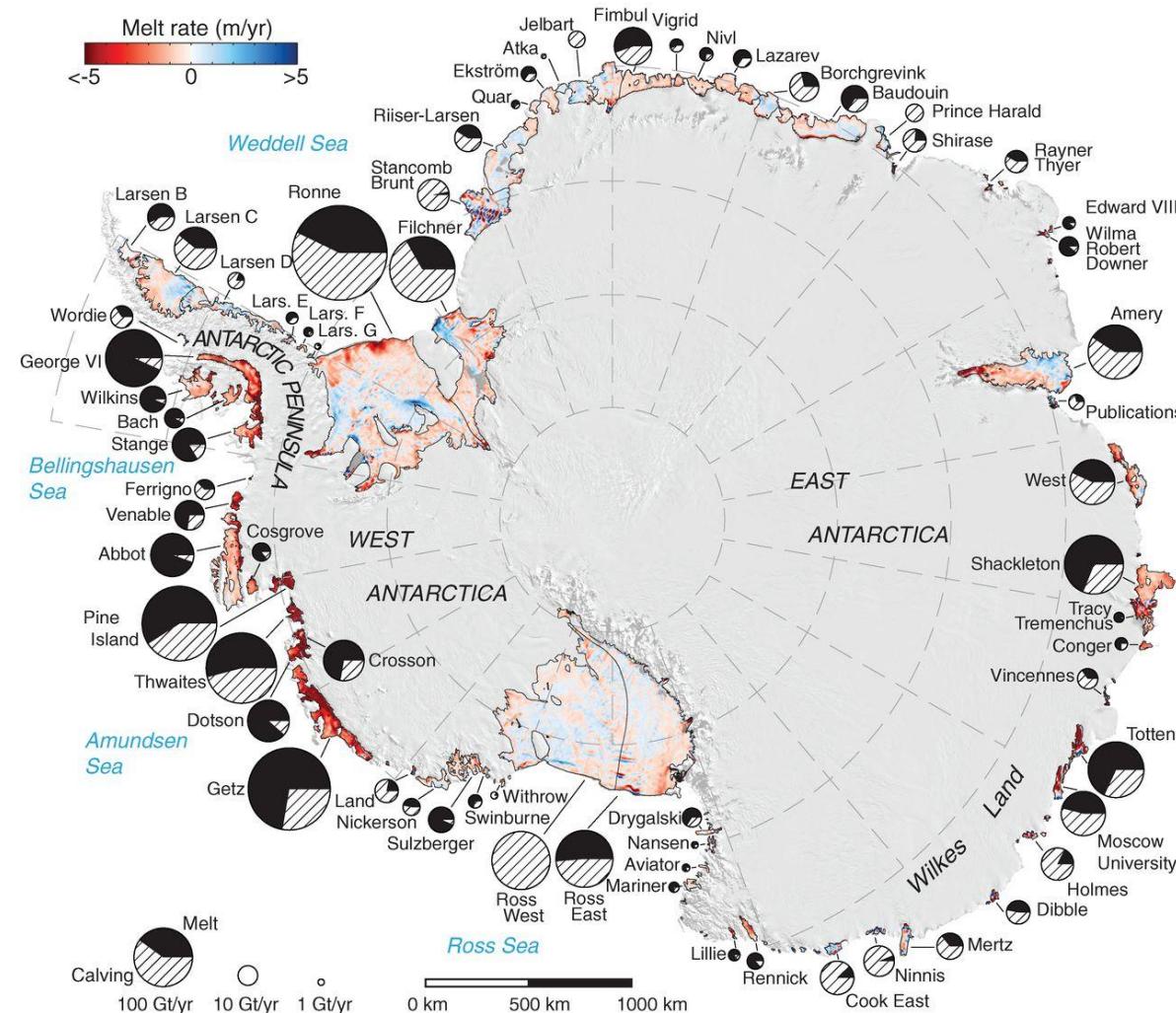
Laura Herraiz-Borreguero, National Oceanographic Centre-Southampton, UK

Peter Schlosser and Robert Newton, LDEO, USA

Steve Rintoul, Beatriz Peña-Molino, ACECRC, Australia

1. East Antarctic ice shelves
2. Amery ice shelf
3. Meltwater tracers: noble gases

Fig. 1 Basal melt rates of Antarctic ice shelves. Color coded from <-5 m/year (freezing) to >+5 m/year (melting) and overlaid on a 2009 Moderate Resolution Imaging Spectroradiometer mosaic of Antarctica.



Antarctic ice–shelf ice–thickness change rate $\Delta T/\Delta t$, 2003–2008

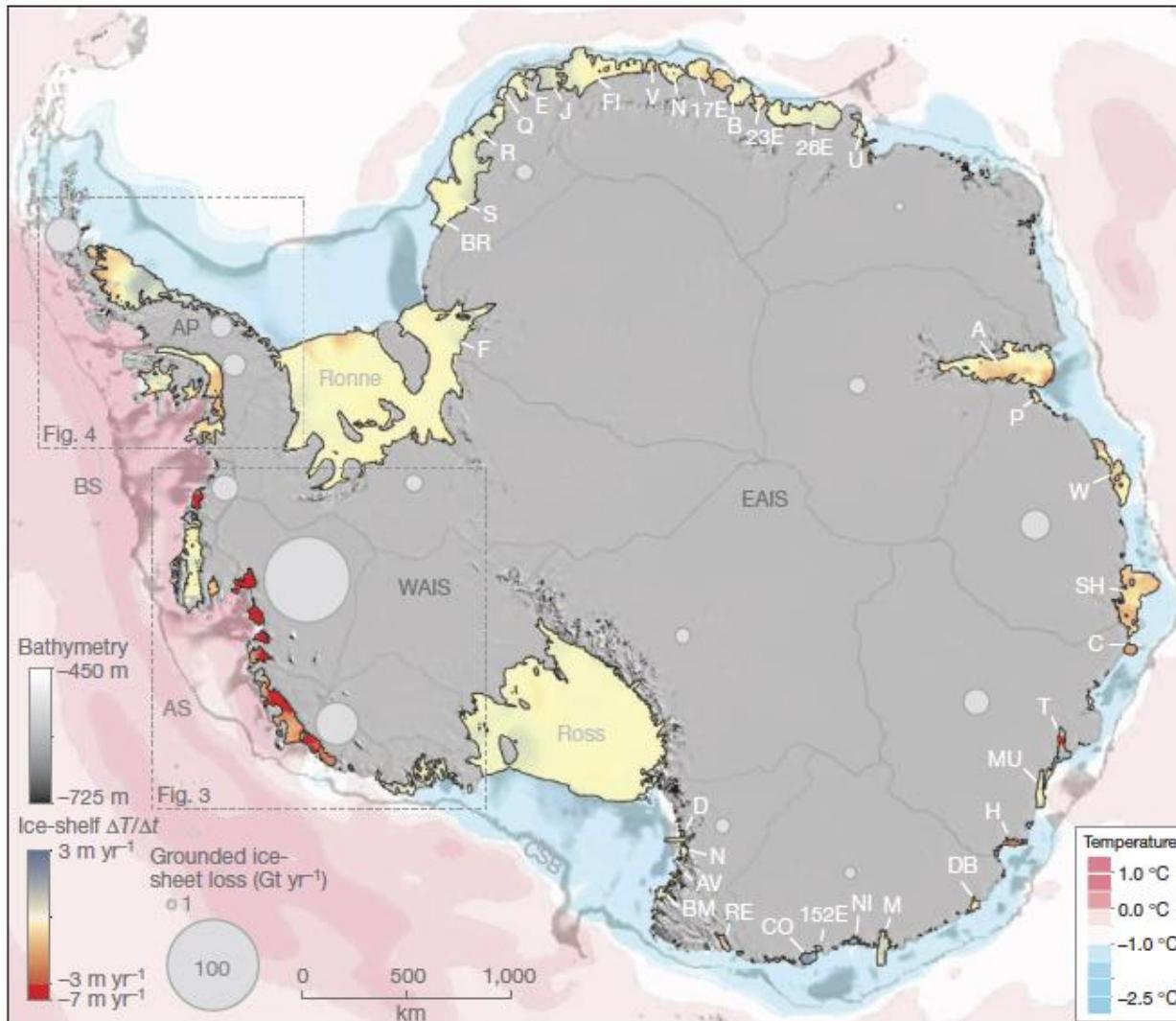
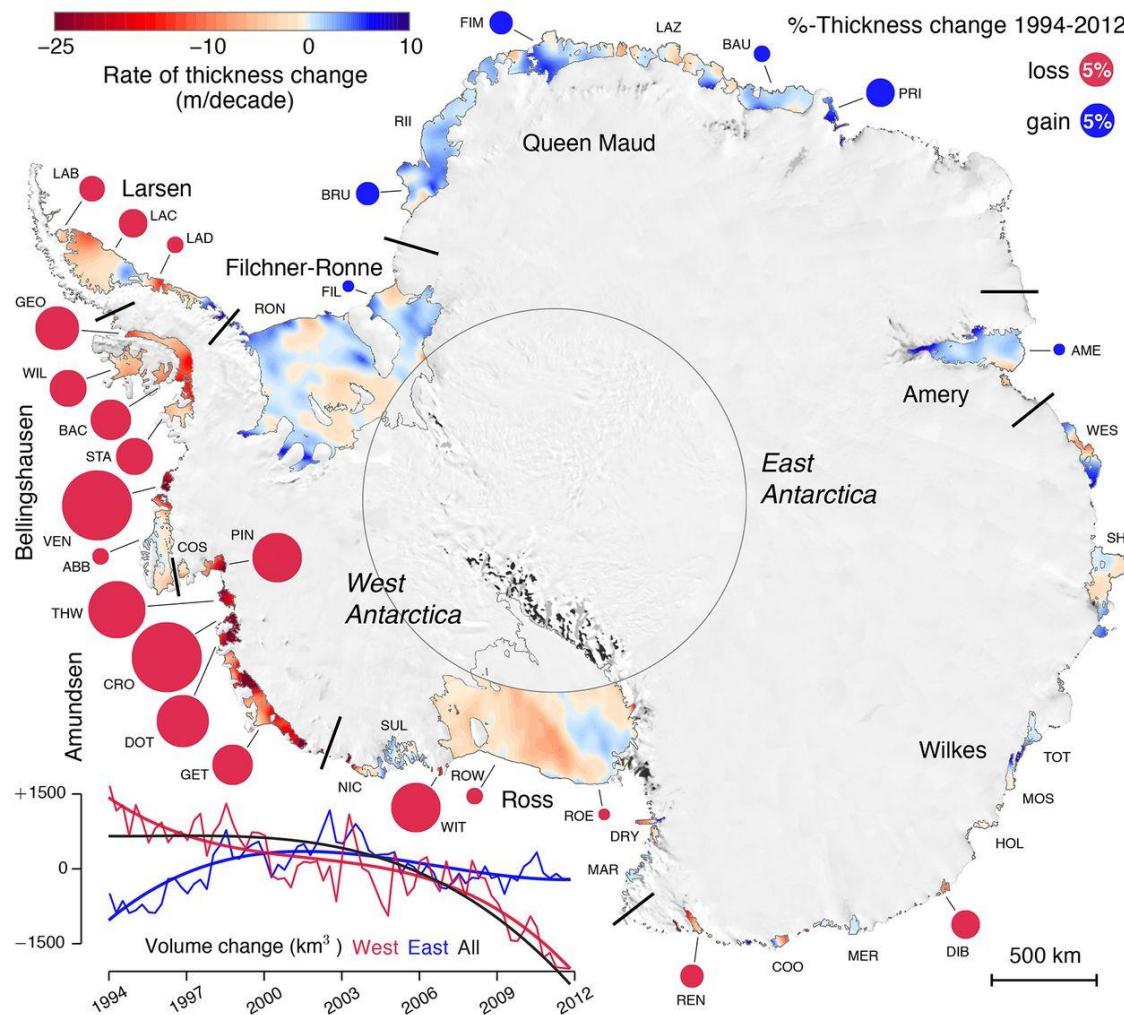
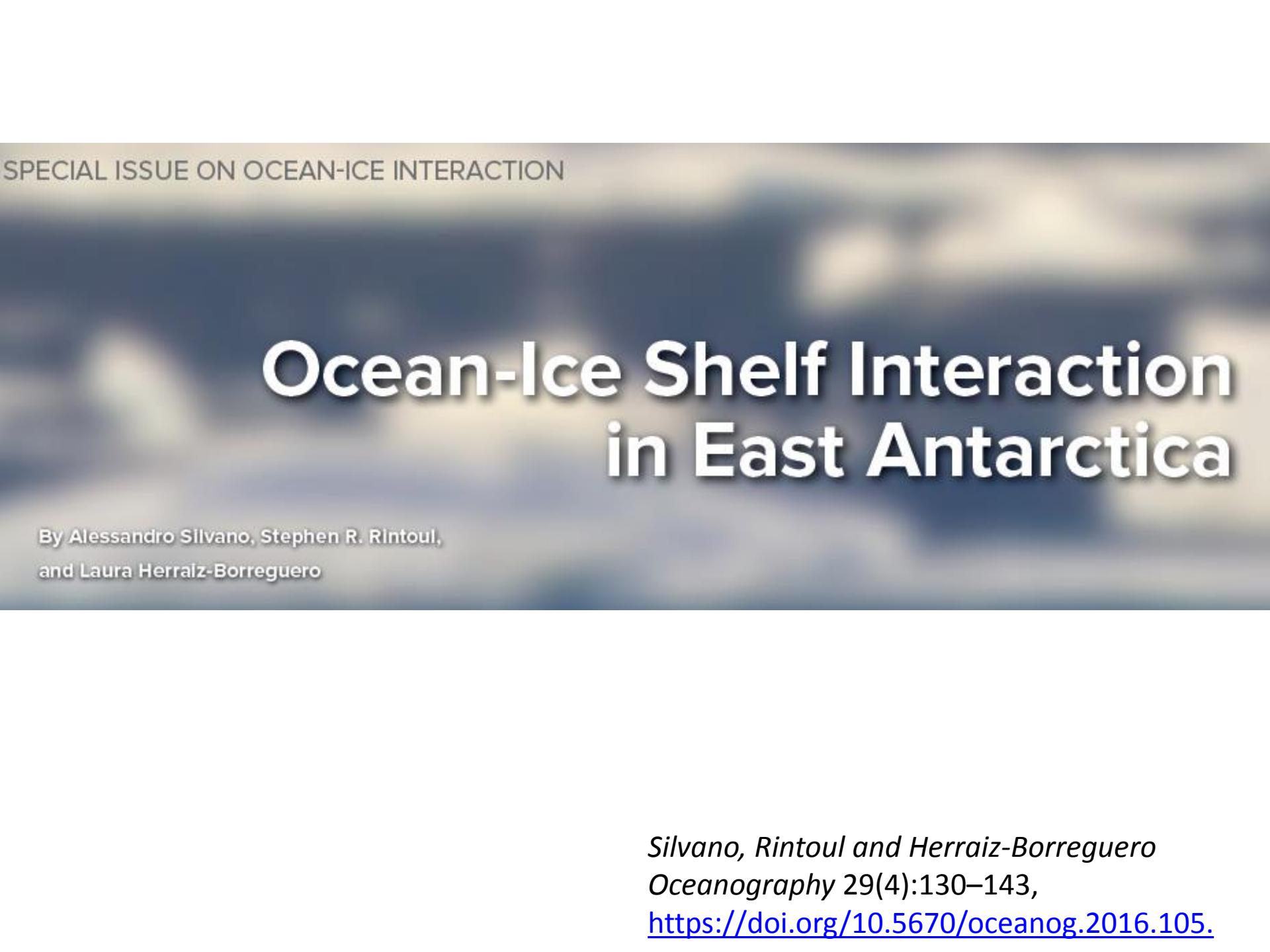


Fig. 1 Eighteen years of change in thickness and volume of Antarctic ice shelves. Rates of thickness change (m/decade) are color-coded from -25 (thinning) to +10 (thickening).





SPECIAL ISSUE ON OCEAN-ICE INTERACTION

Ocean-Ice Shelf Interaction in East Antarctica

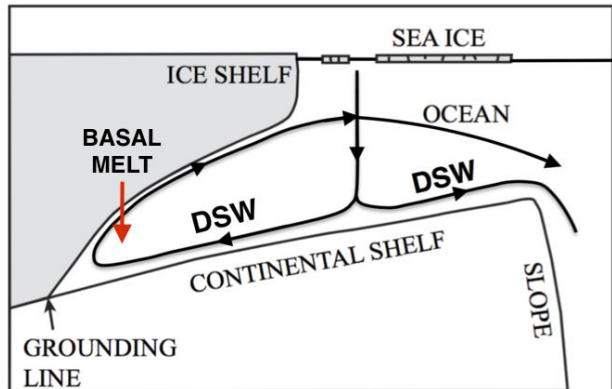
By Alessandro Silvano, Stephen R. Rintoul,
and Laura Herraiz-Borreguero

Silvano, Rintoul and Herraiz-Borreguero
Oceanography 29(4):130–143,
<https://doi.org/10.5670/oceanog.2016.105>.

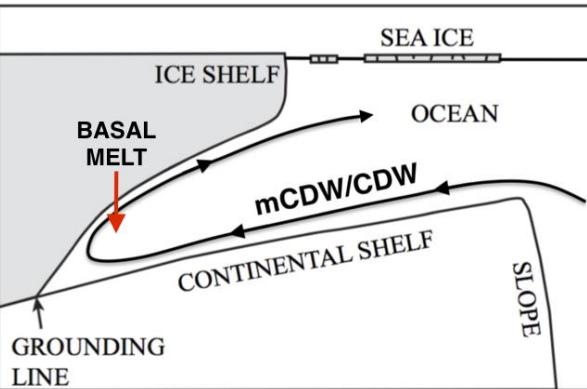
Ocean-Ice Shelf Interaction in East Antarctica

By Alessandro Silvano, Stephen R. Rintoul,
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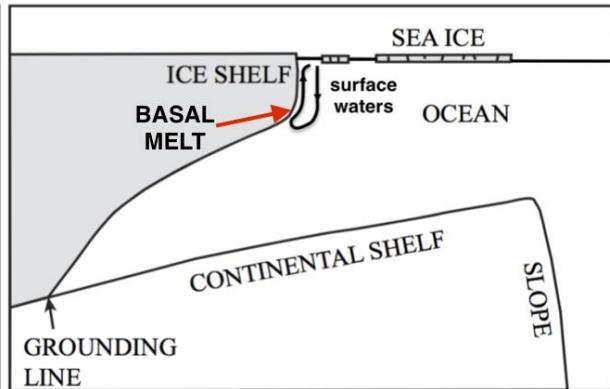
Mode 1



Mode 2

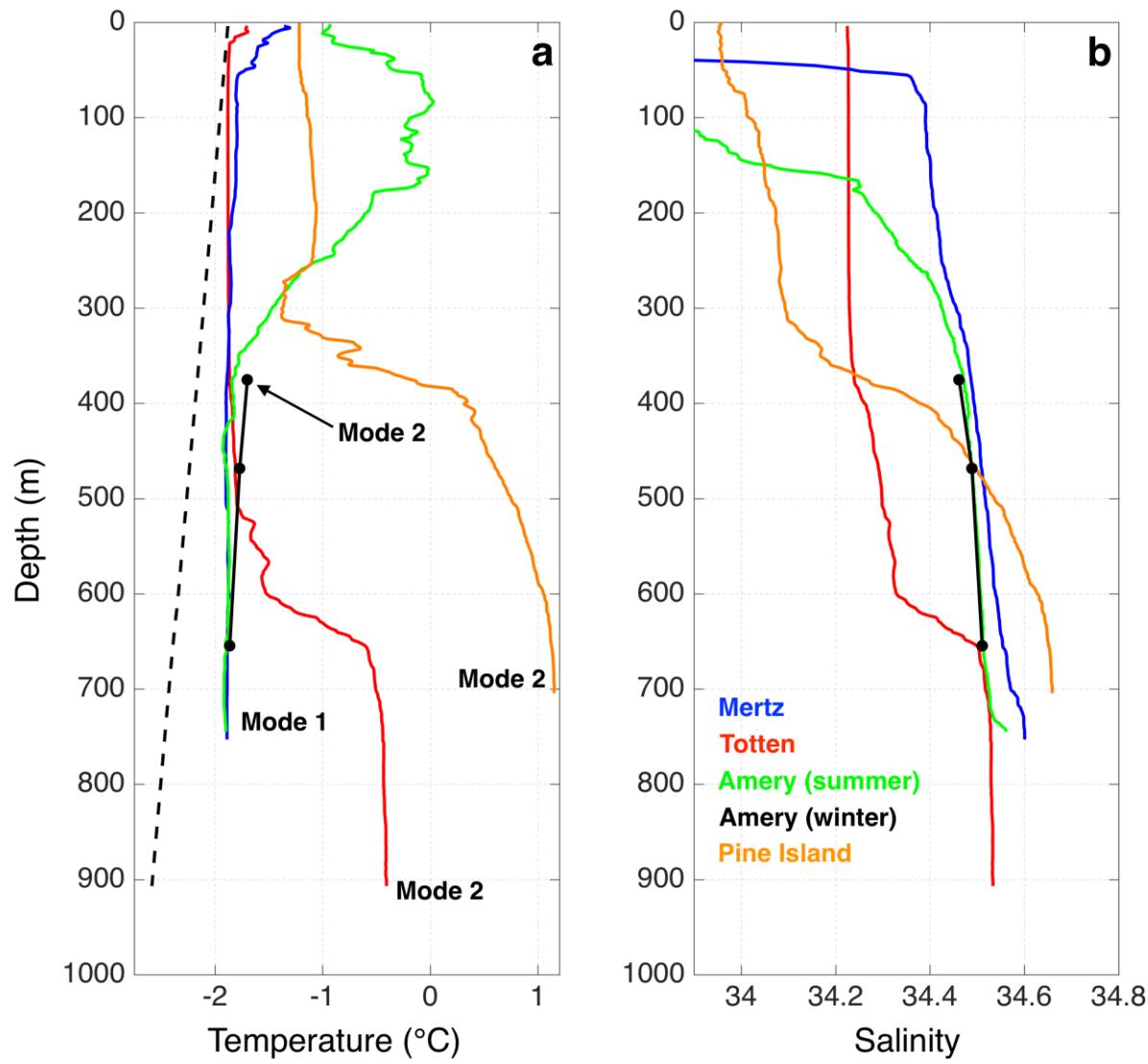


Mode 3

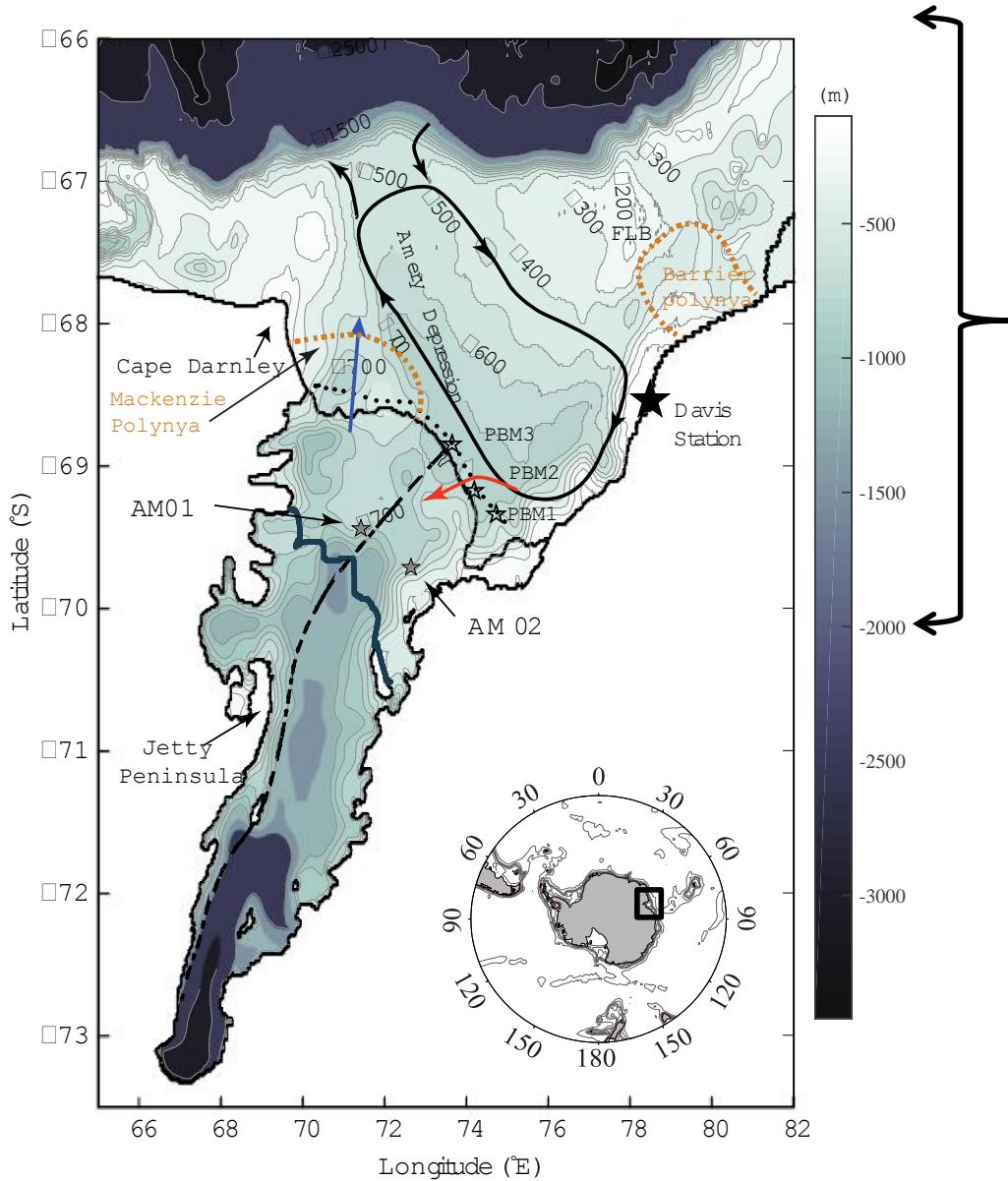


Mertz Glacial Tongue

Water properties of the strongest inflow to ice-shelf cavities in East and West Antarctica.



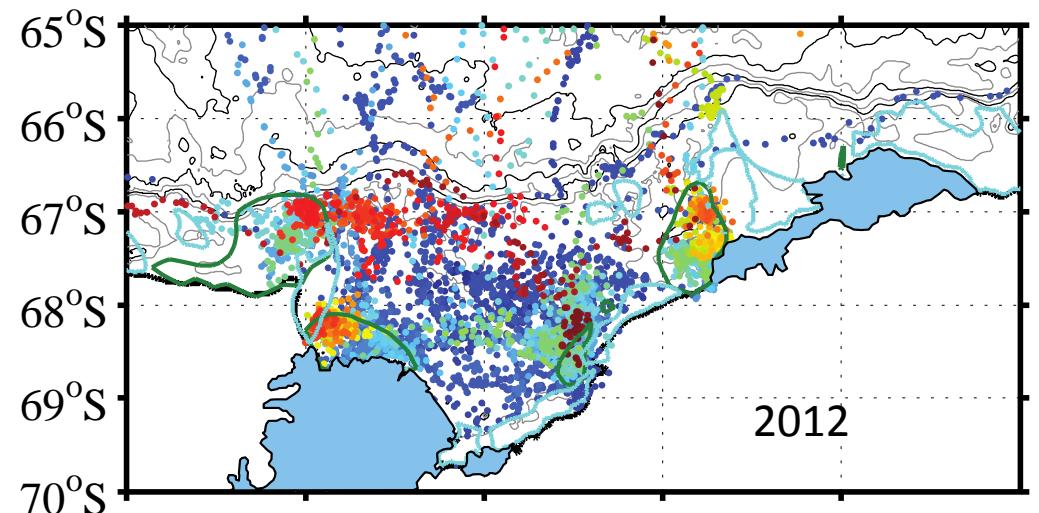
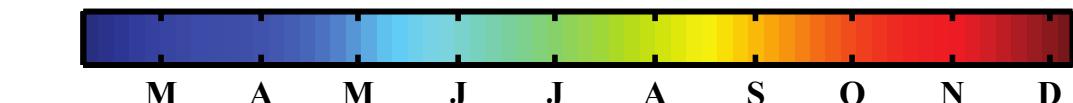
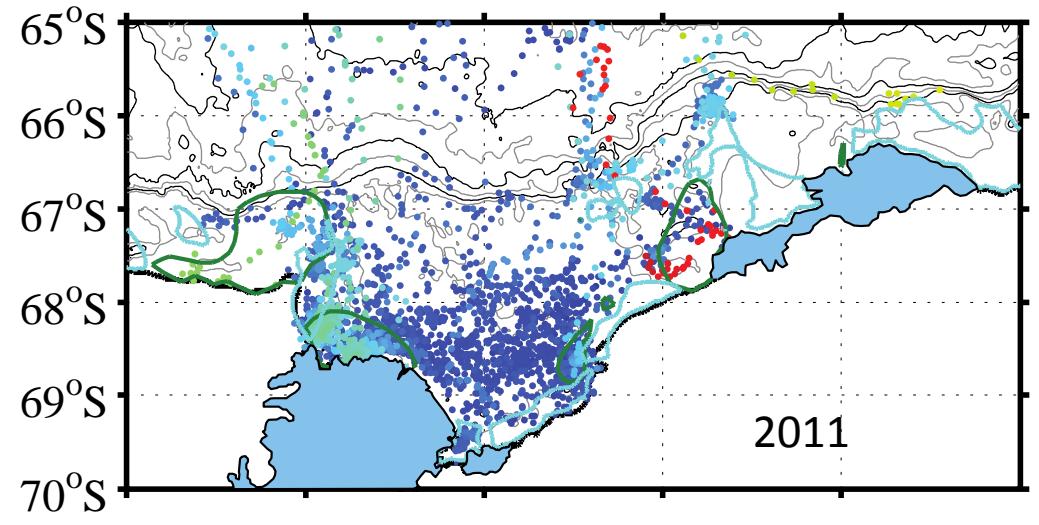
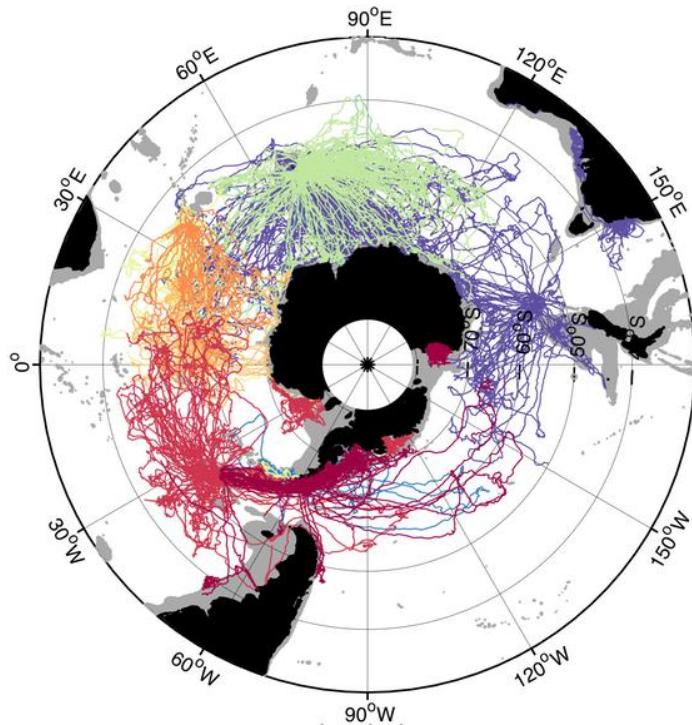
Prydz Bay-Amery ice shelf system: observational platforms



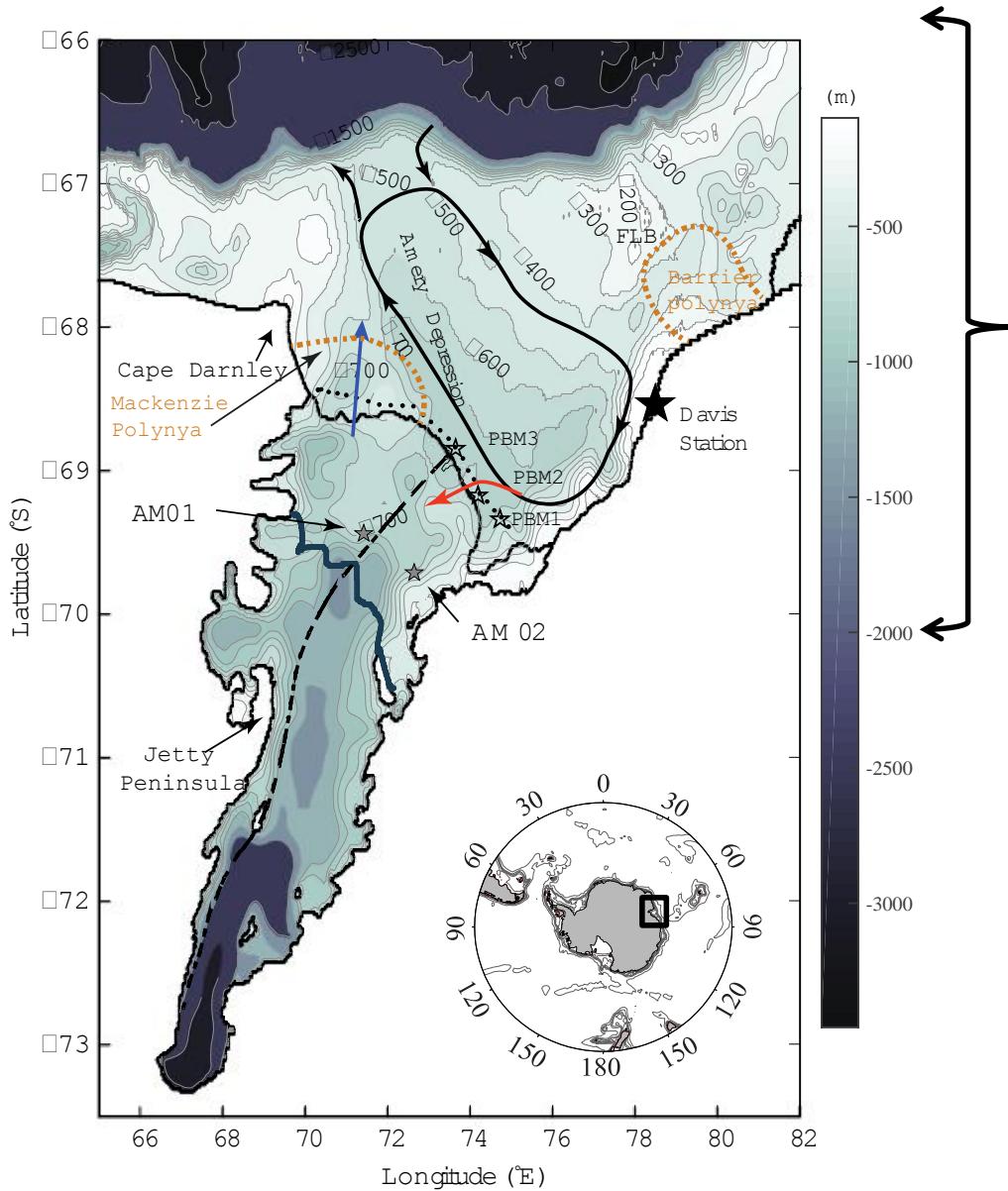
Instrumented marine mammals
(high spatial and temporal coverage)



Prydz Bay-Amery ice shelf system: observational platforms

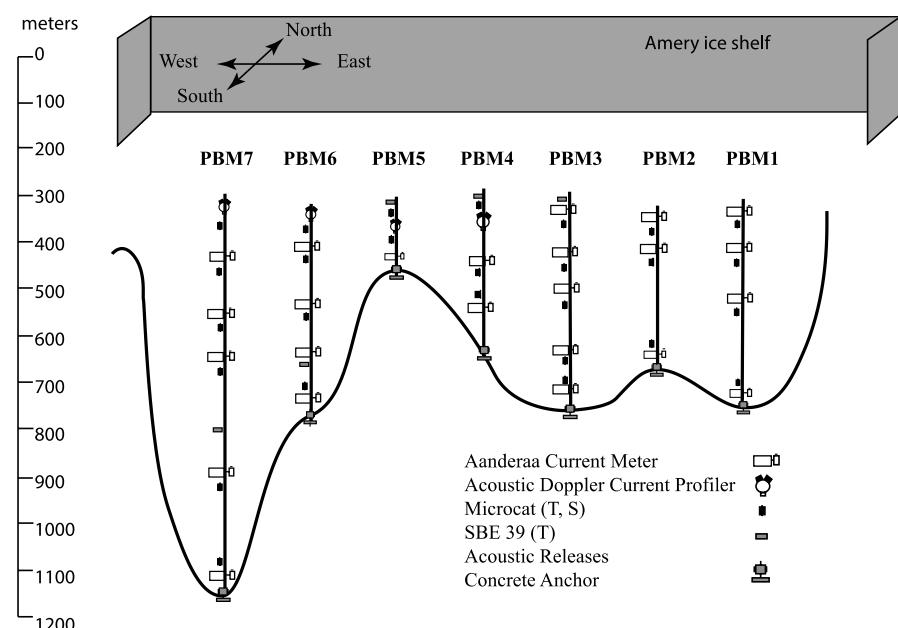


Prydz Bay-Amery ice shelf system: observational platforms

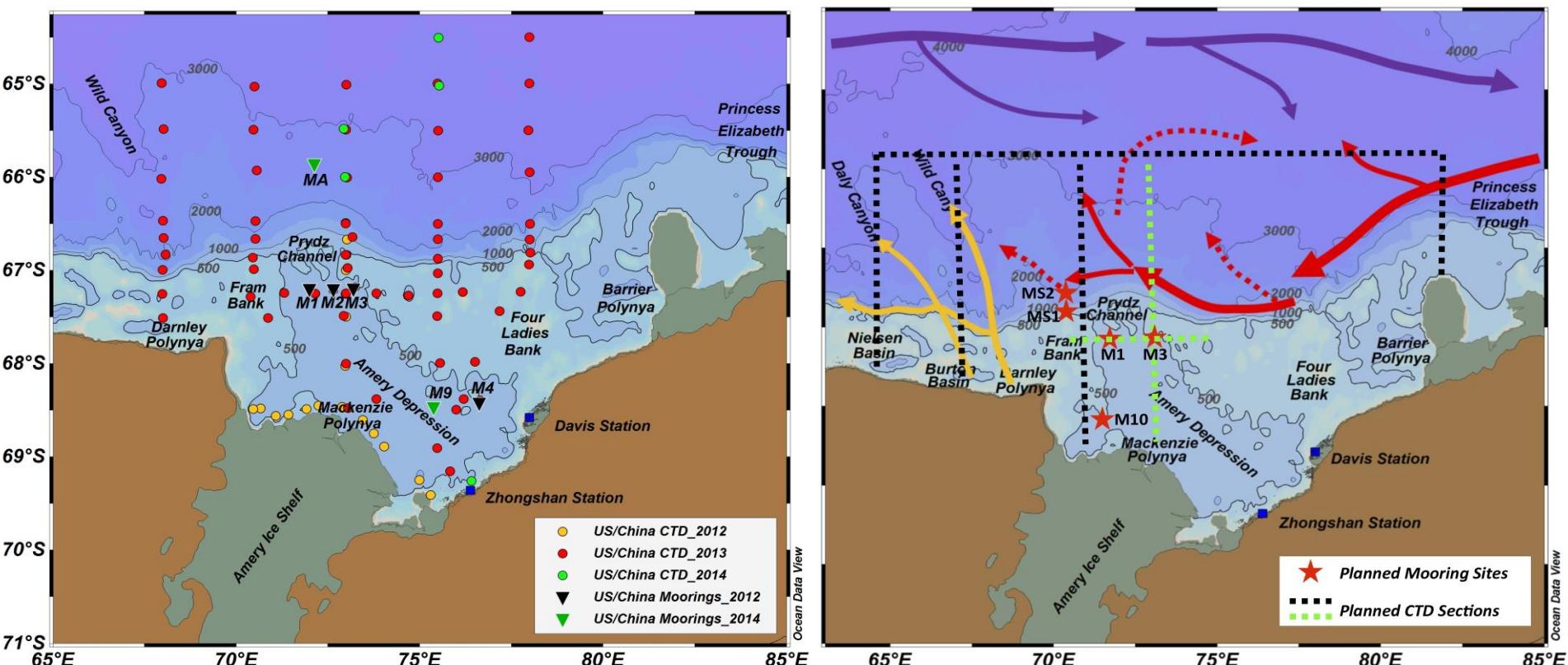


Open ocean instrumented moorings
(temporal resolution)

Ship-based CTD profiles (vertical
resolution)



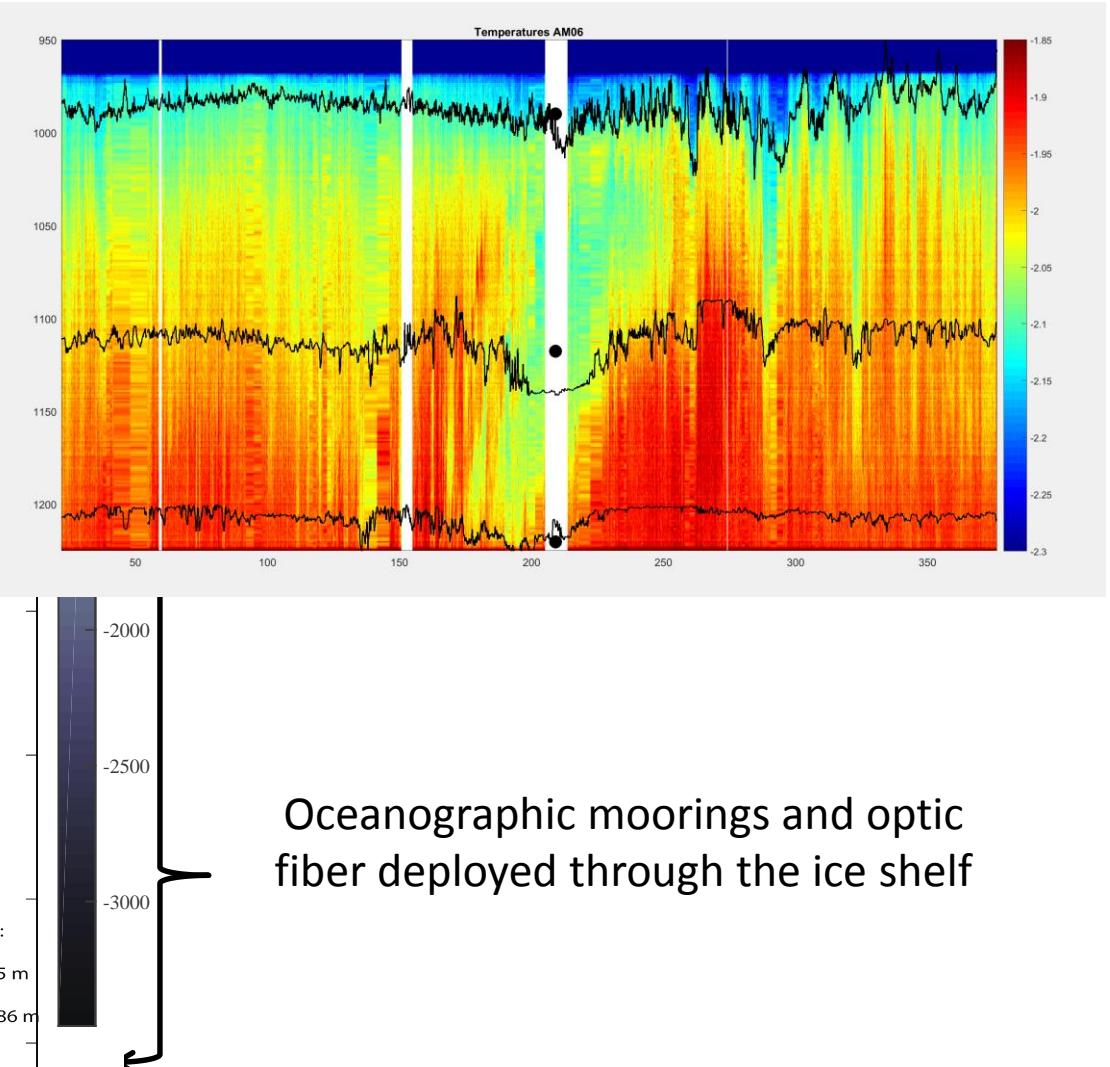
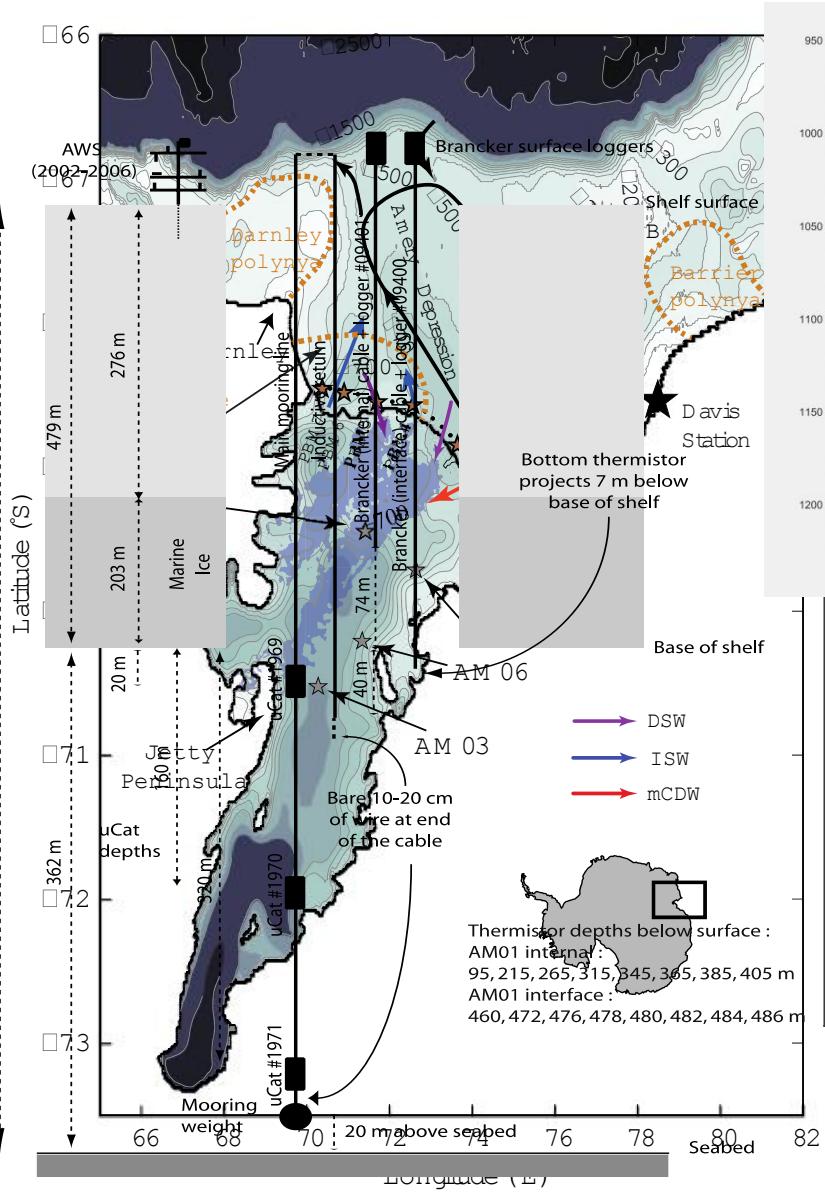
Prydz Bay-Amery ice shelf system: observational platforms



U.S./Chinese Collaborative Study: Investigation of Bottom Water Formation in Prydz Bay, Antarctica

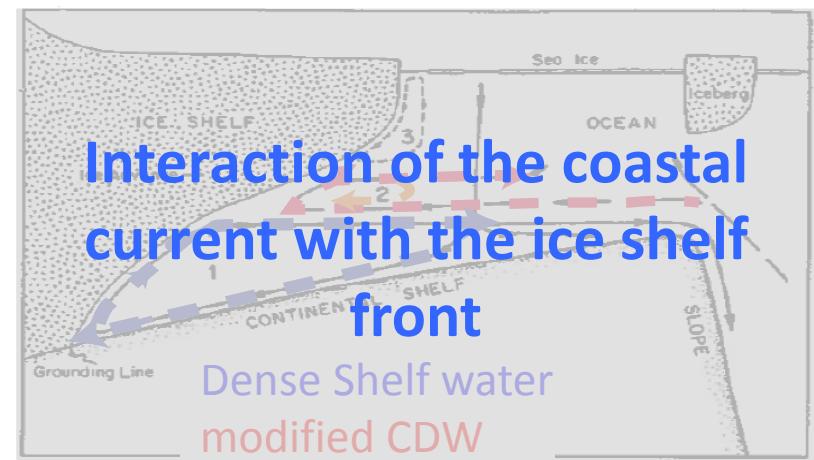
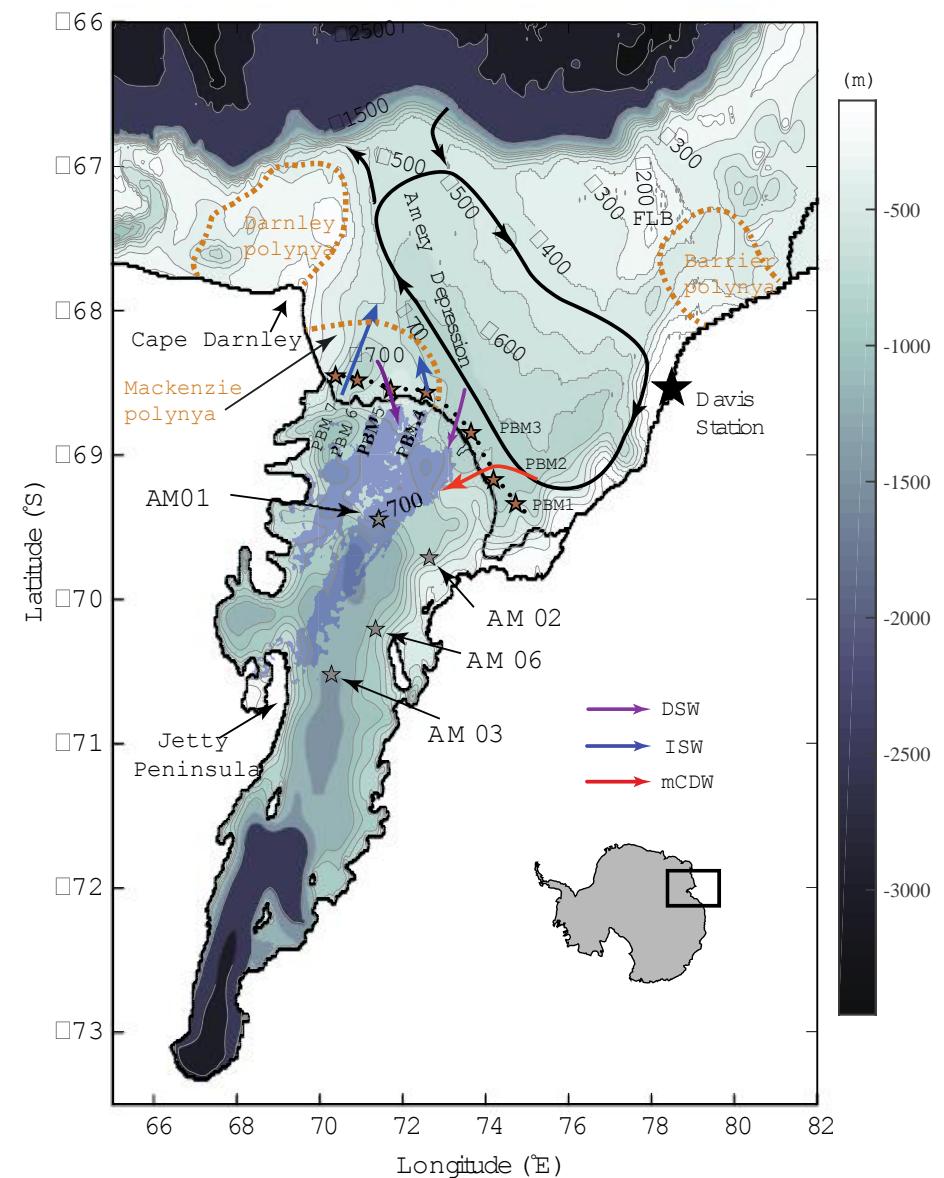
Prydz Bay-Amery ice shelf system: observational platforms ?

Warner et al., IGS 2016



Prydz Bay-Amery ice shelf system: what do we know ?

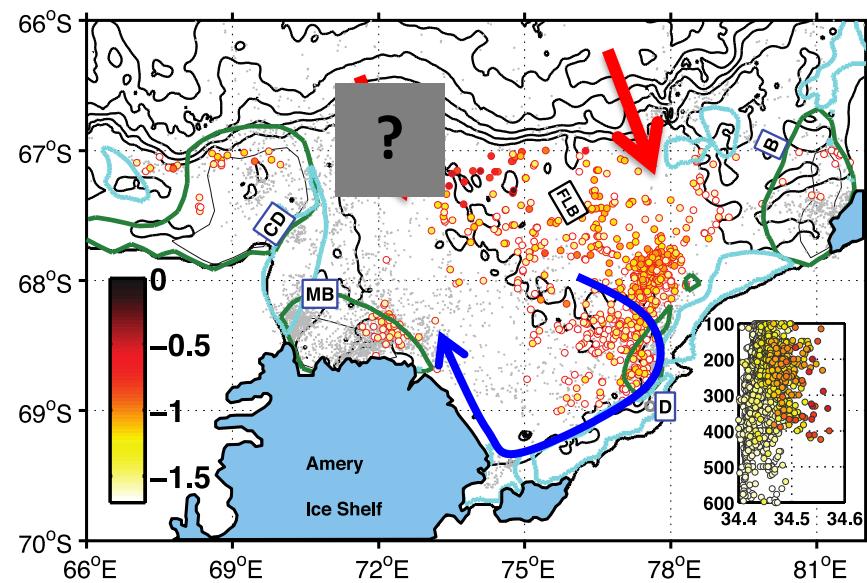
Eastern Side: Mode 1 and 2 overlap during the autumn-winter months



Western side: Mode 1 dominates the circulation beneath the AIS



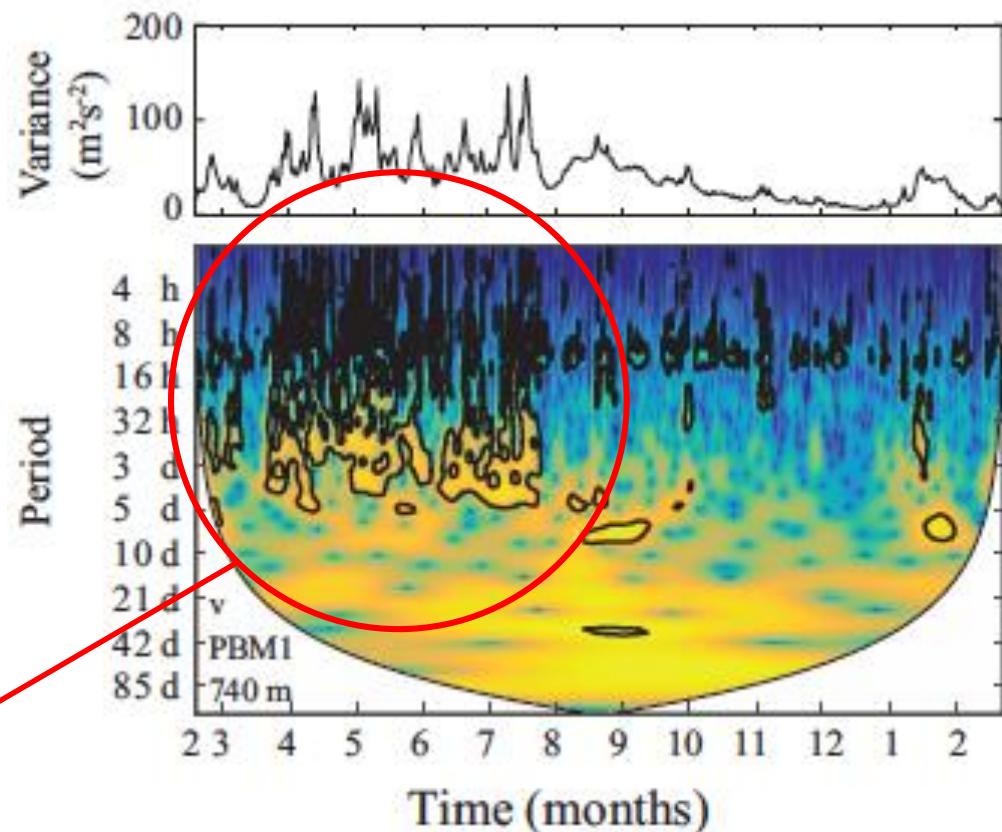
Prydz Bay-Amery ice shelf system: mCDW inflow and coastal current



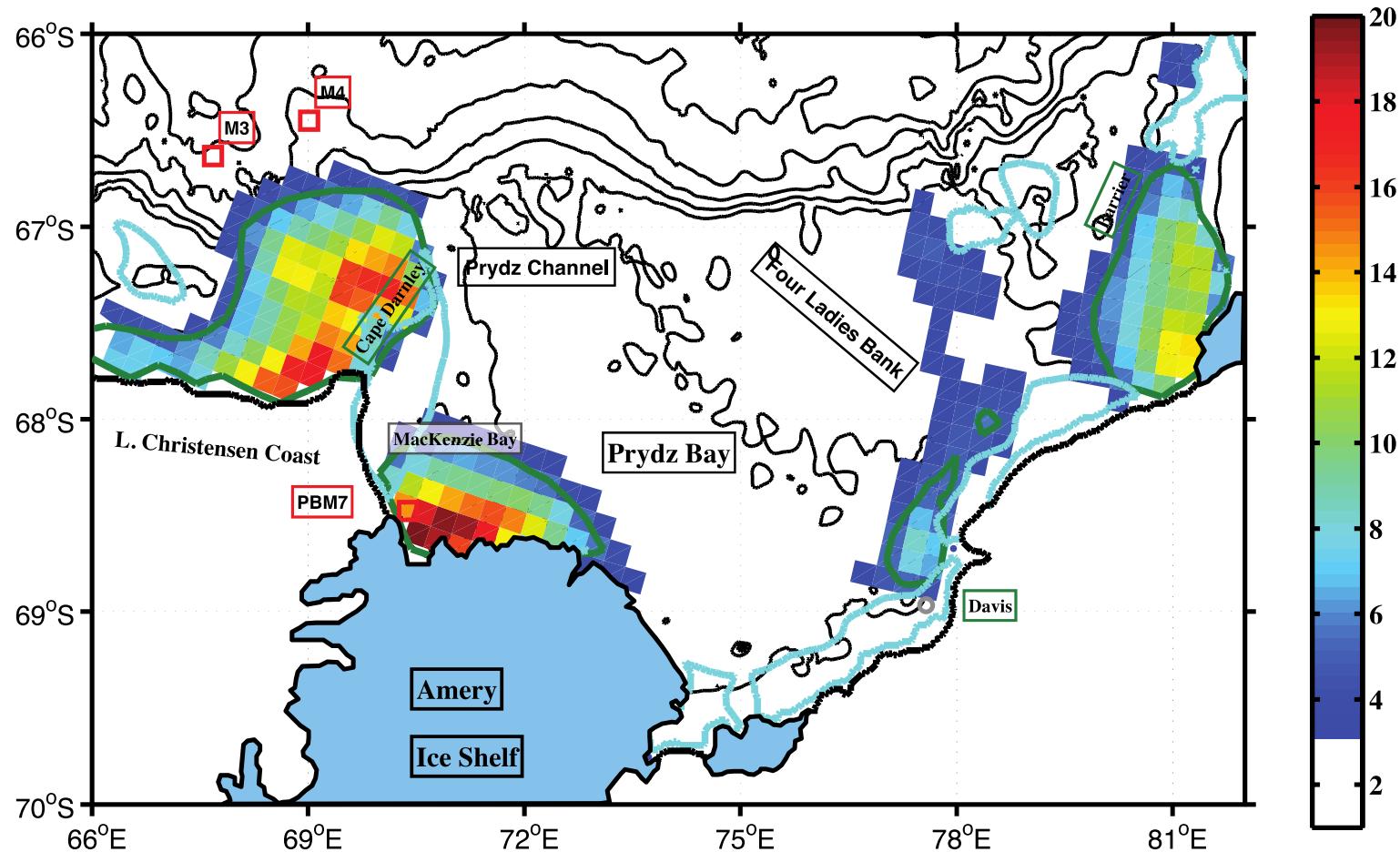
In regional scales, conservation of potential vorticity causes the flow to follow contours of constant depth. Thus, the ice shelf constitutes a barrier for the flow to enter the ice shelf cavity.

Eddies and coastally-trapped waves

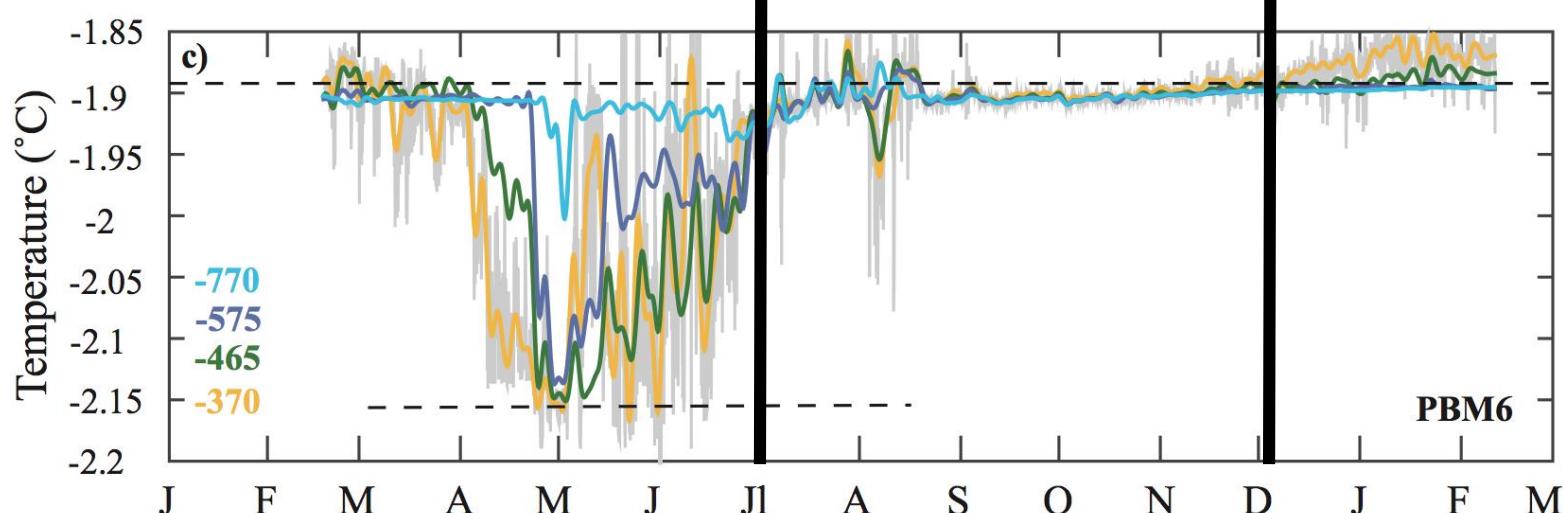
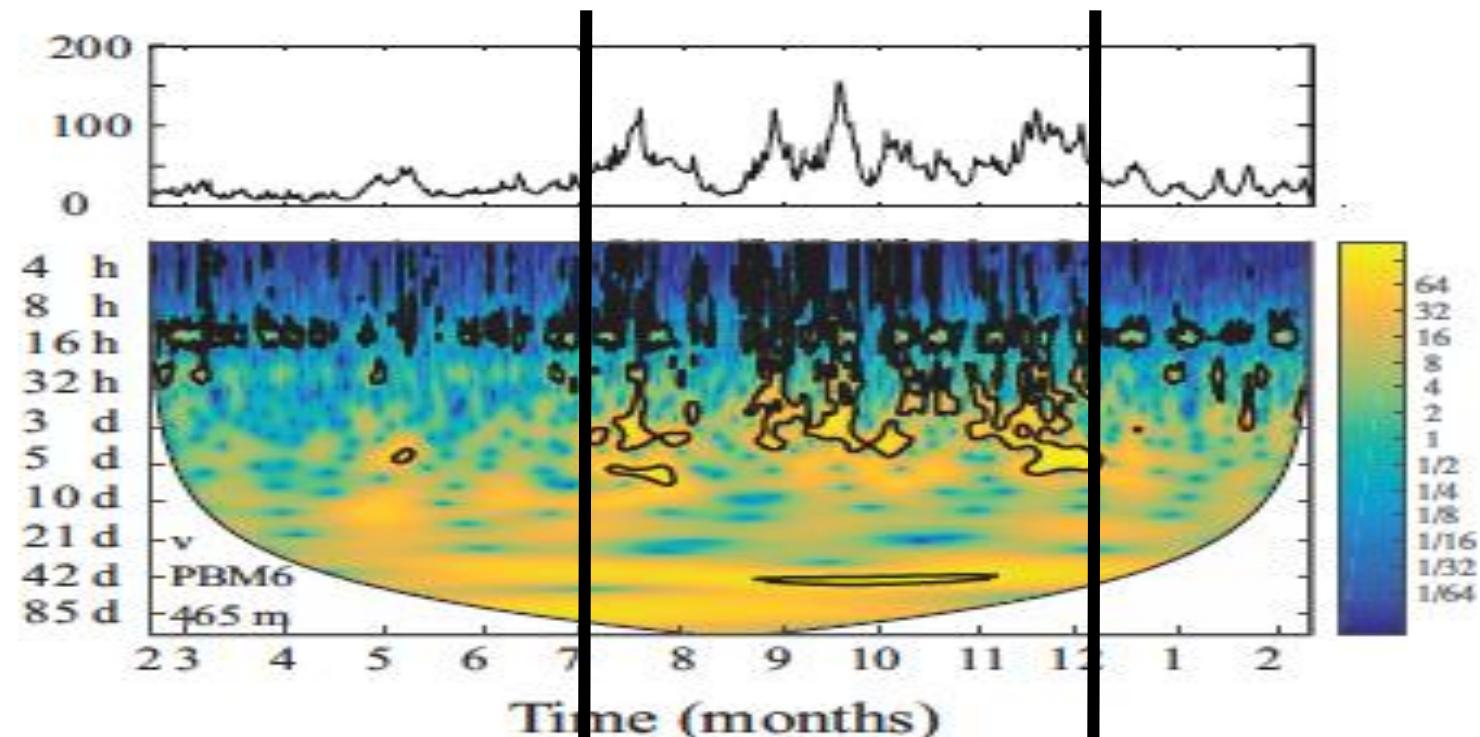
Wavelet analyses of PBM1, at 740 m



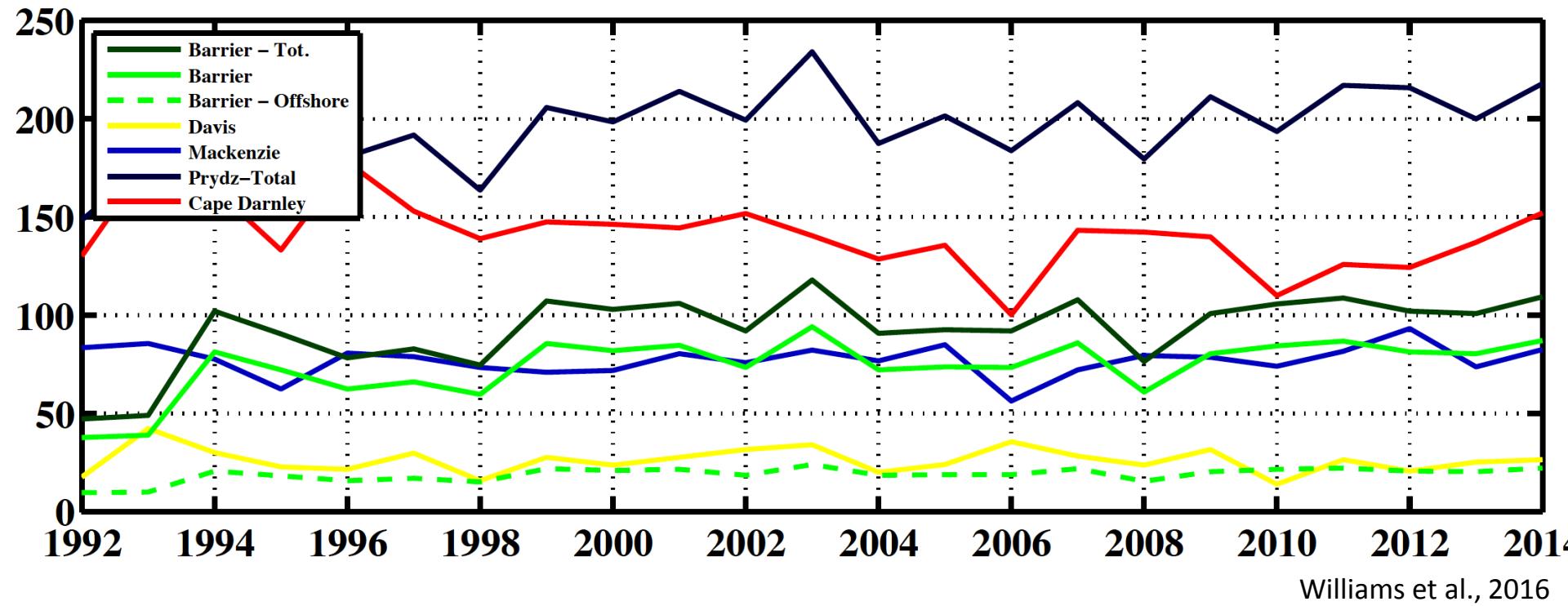
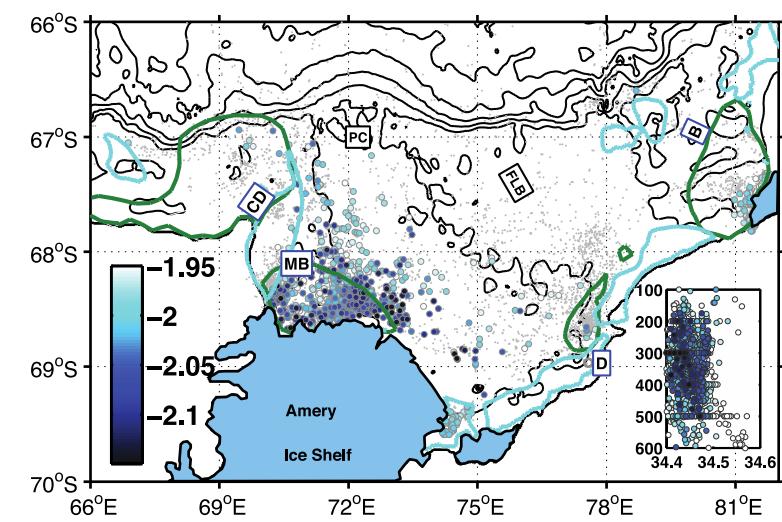
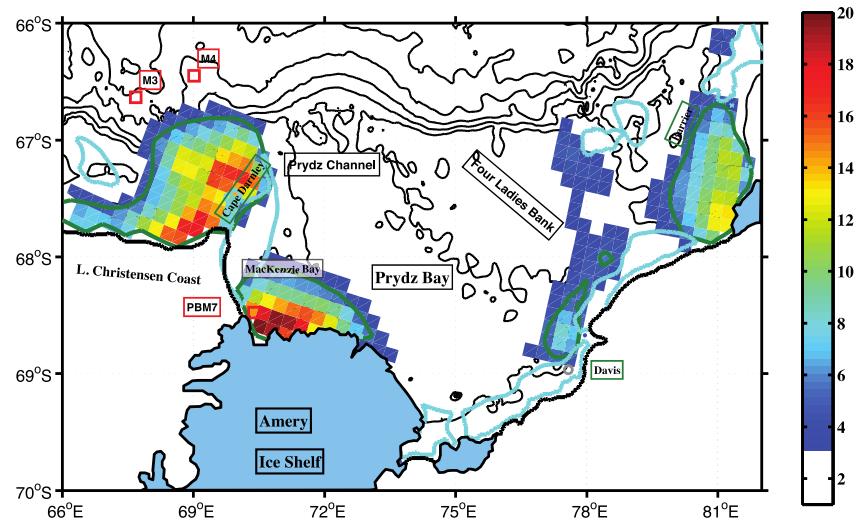
Prydz Bay-Amery ice shelf system: Mackensie Polynya and ocean circulation variability



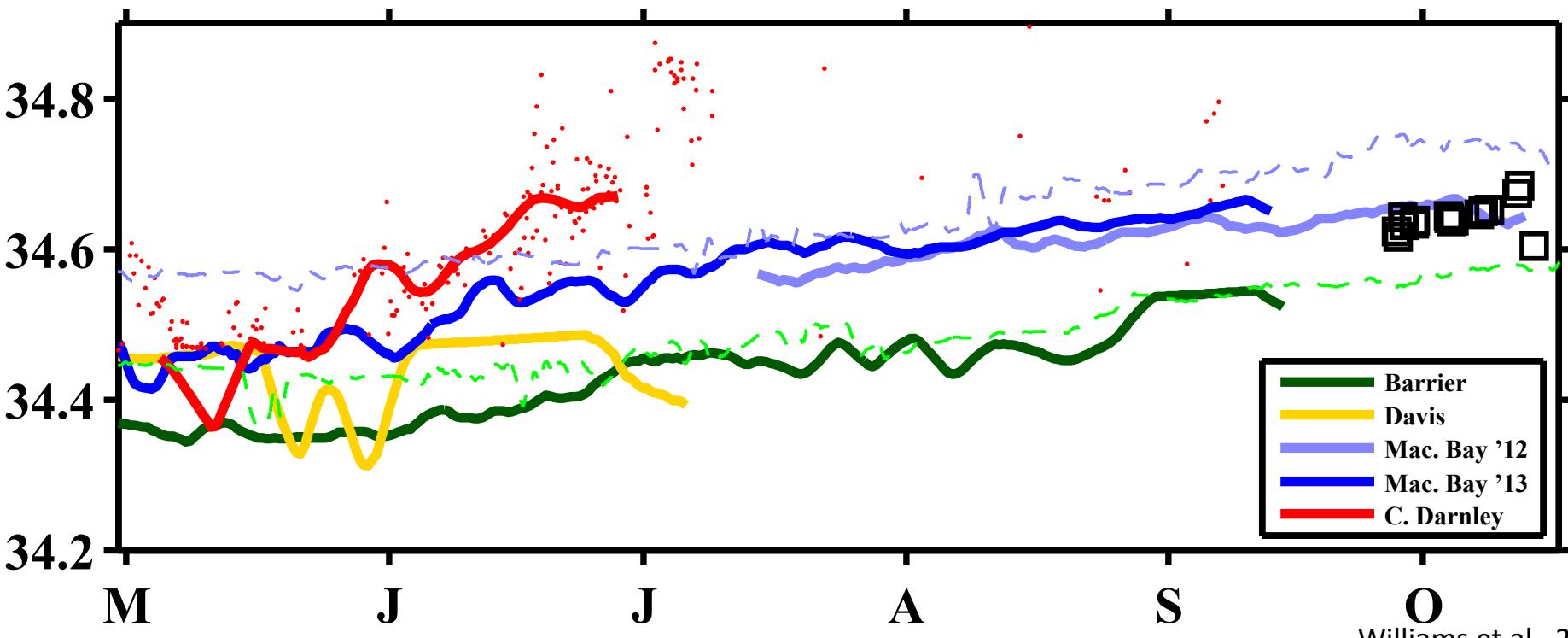
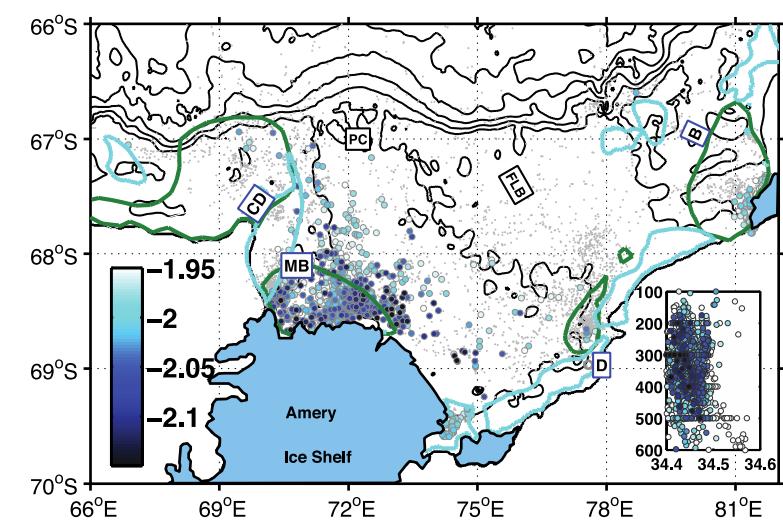
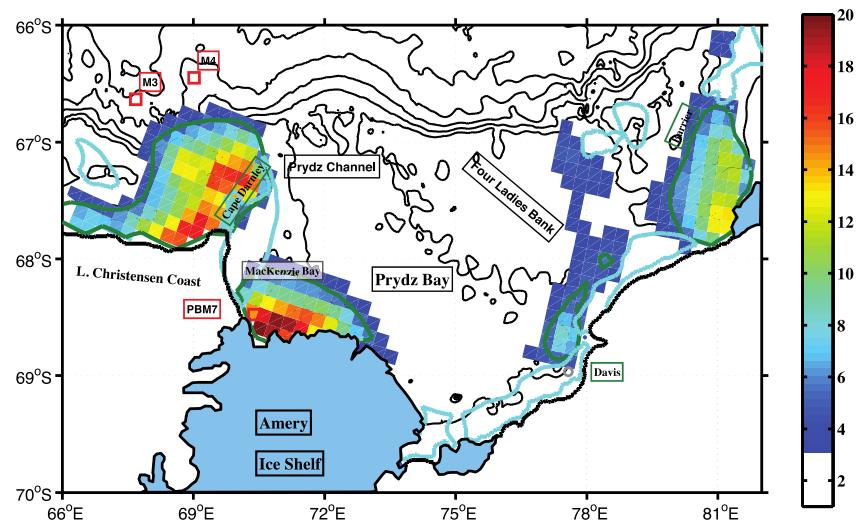
Prydz Bay-Amery ice shelf system: Mackensie Polynya and ocean circulation variability



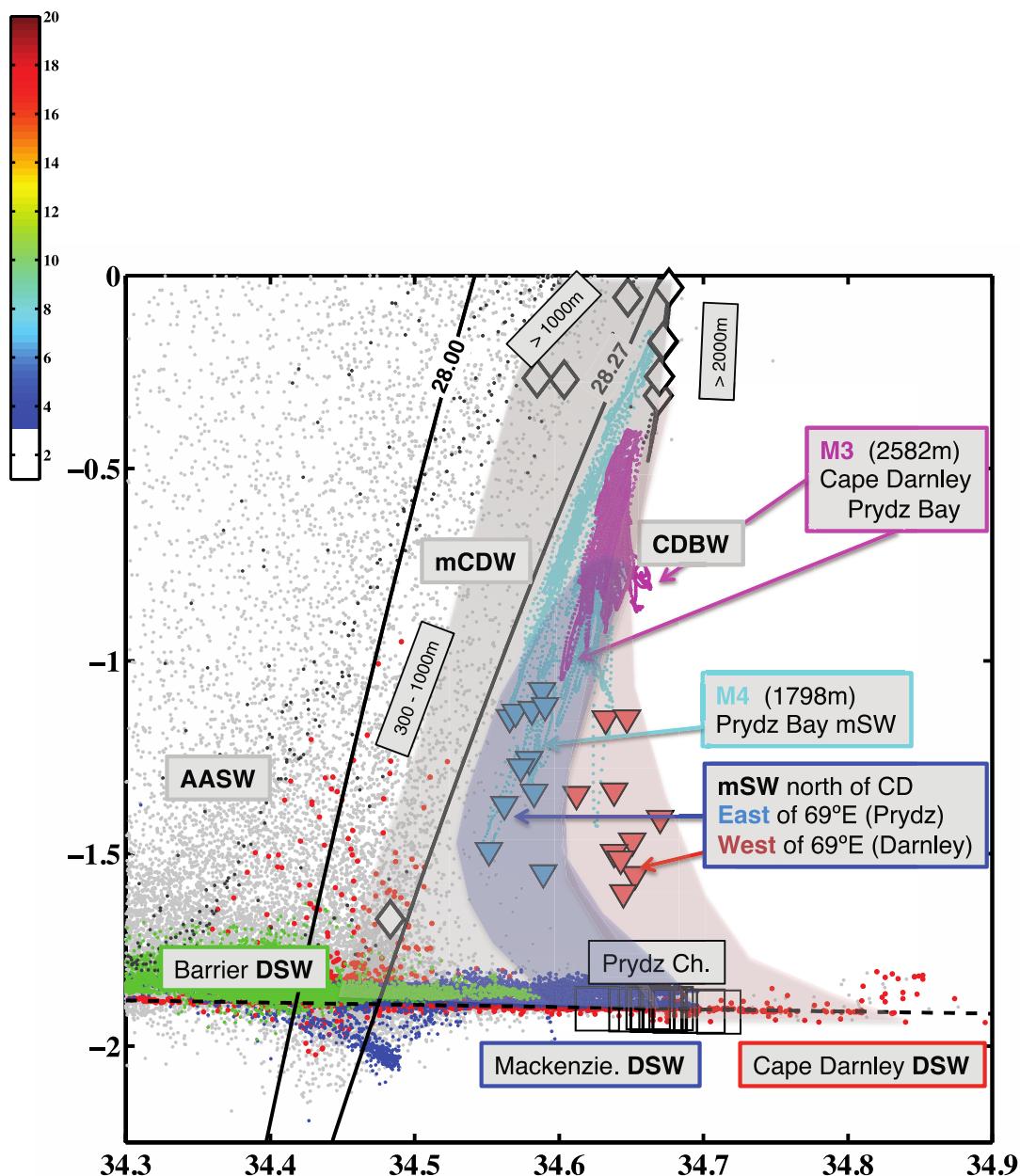
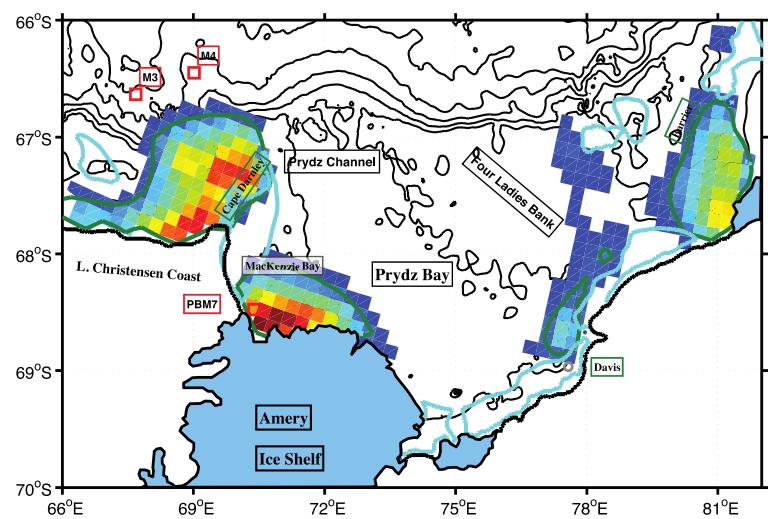
Prydz Bay-Amery ice shelf system: Mackensie Polynya and ocean circulation variability



Prydz Bay-Amery ice shelf system: Mackenzie Polynya and ocean circulation variability

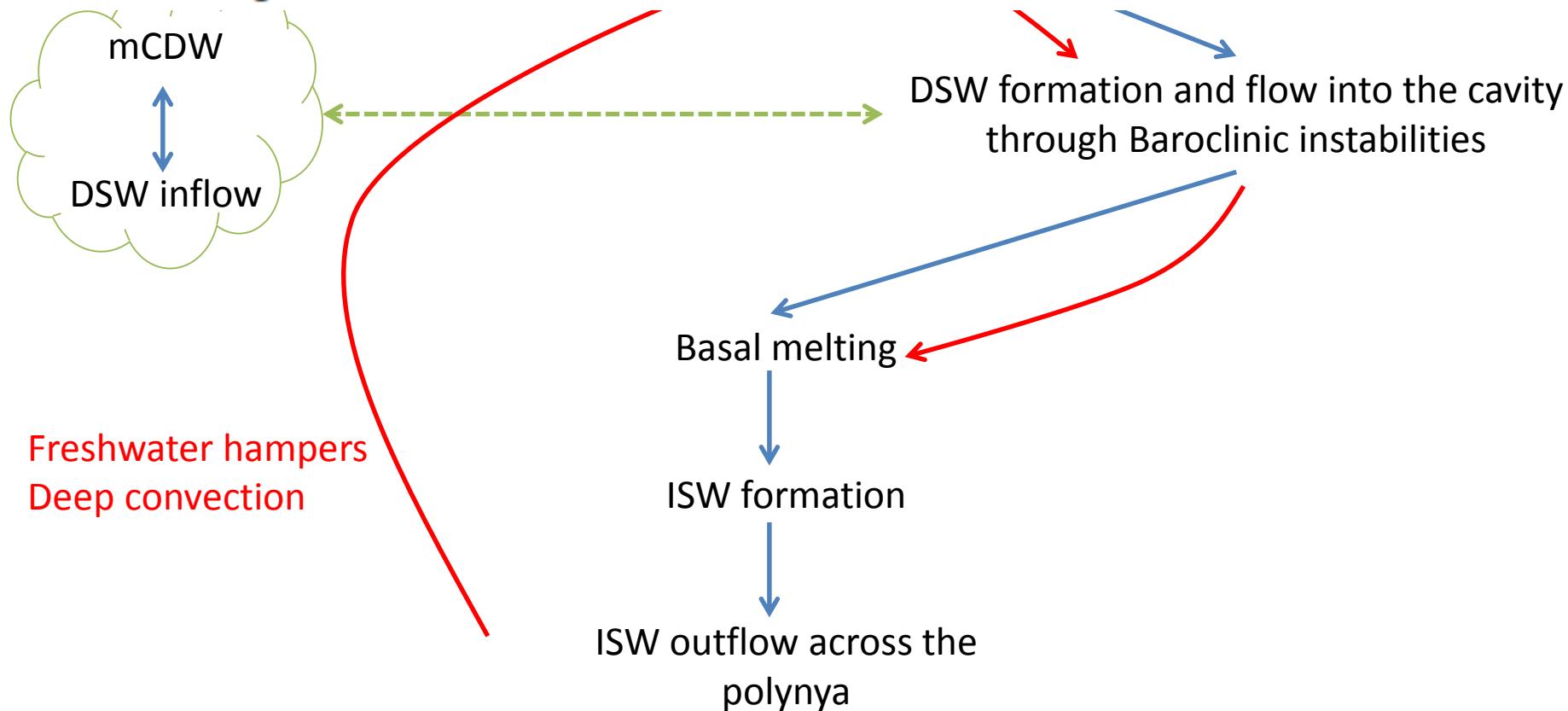


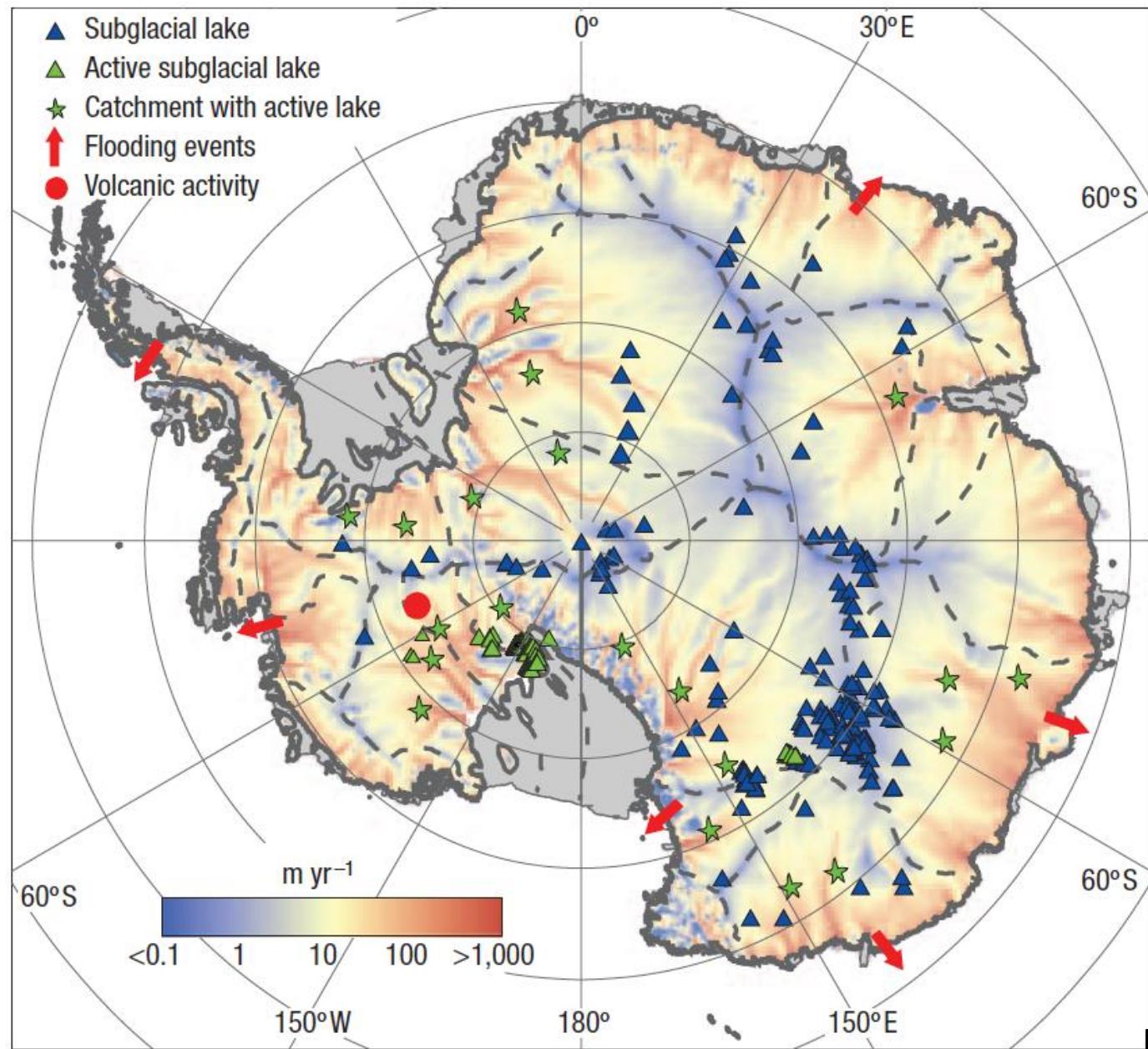
Prydz Bay-Amery ice shelf system: Mackensie Polynya and ocean circulation variability



Large flux of iron from the Amery Ice Shelf marine ice to Prydz Bay, East Antarctica

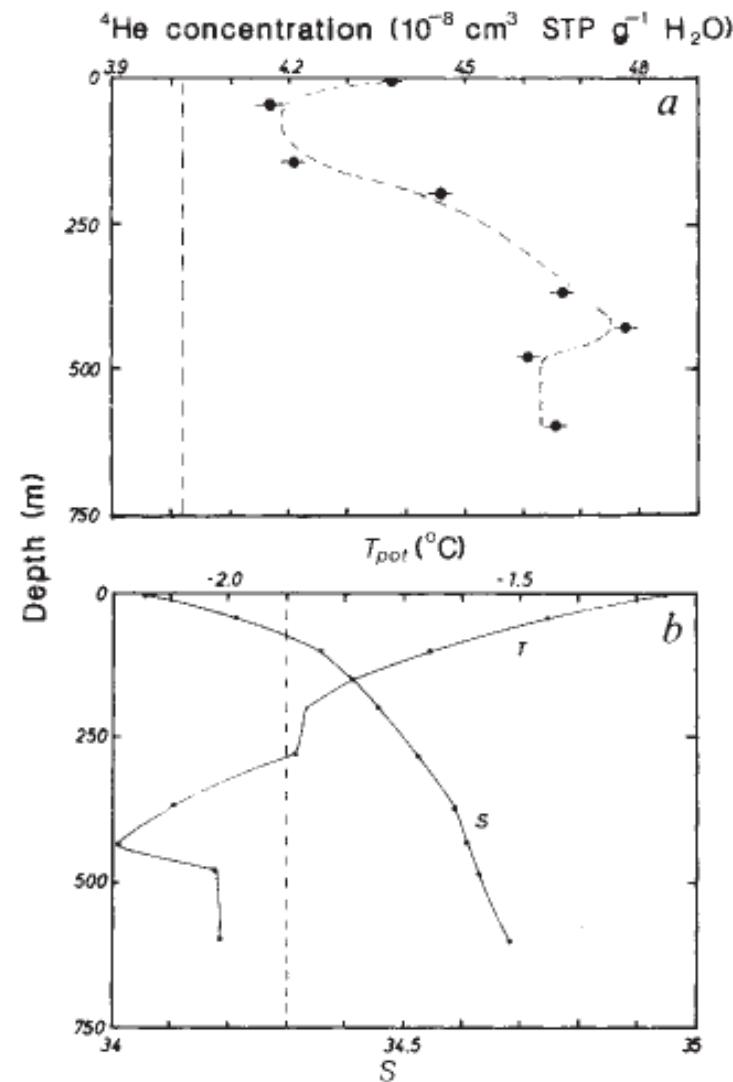
L. Herraiz-Borreguero^{1,2}, D. Lannuzel^{3,4}, P. van der Merwe³, A. Treverrow³, and J. B. Pedro¹



a

Bell, 2008

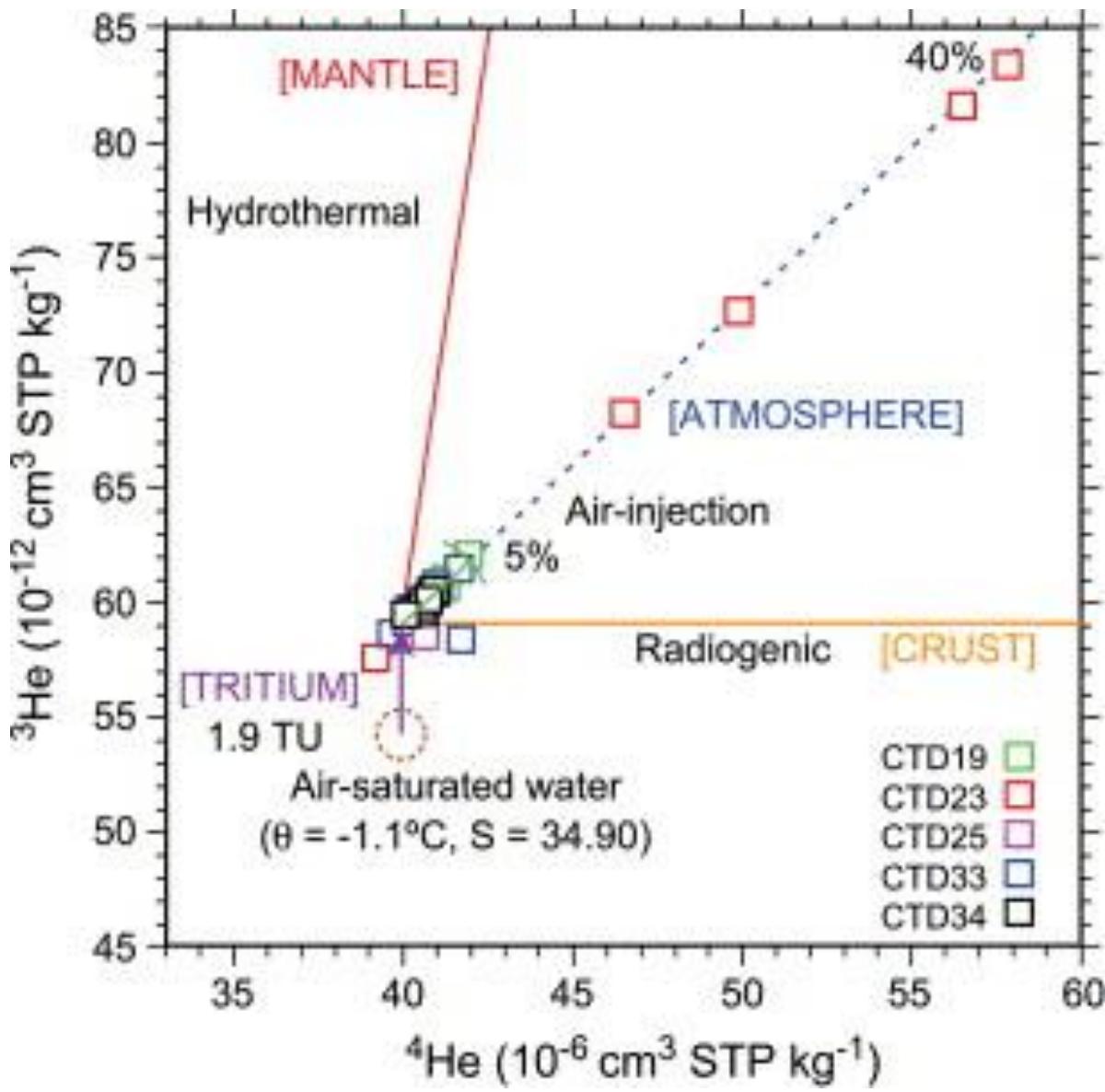
Helium ($^{3,4}\text{He}$), a good tracer of glacial melt water

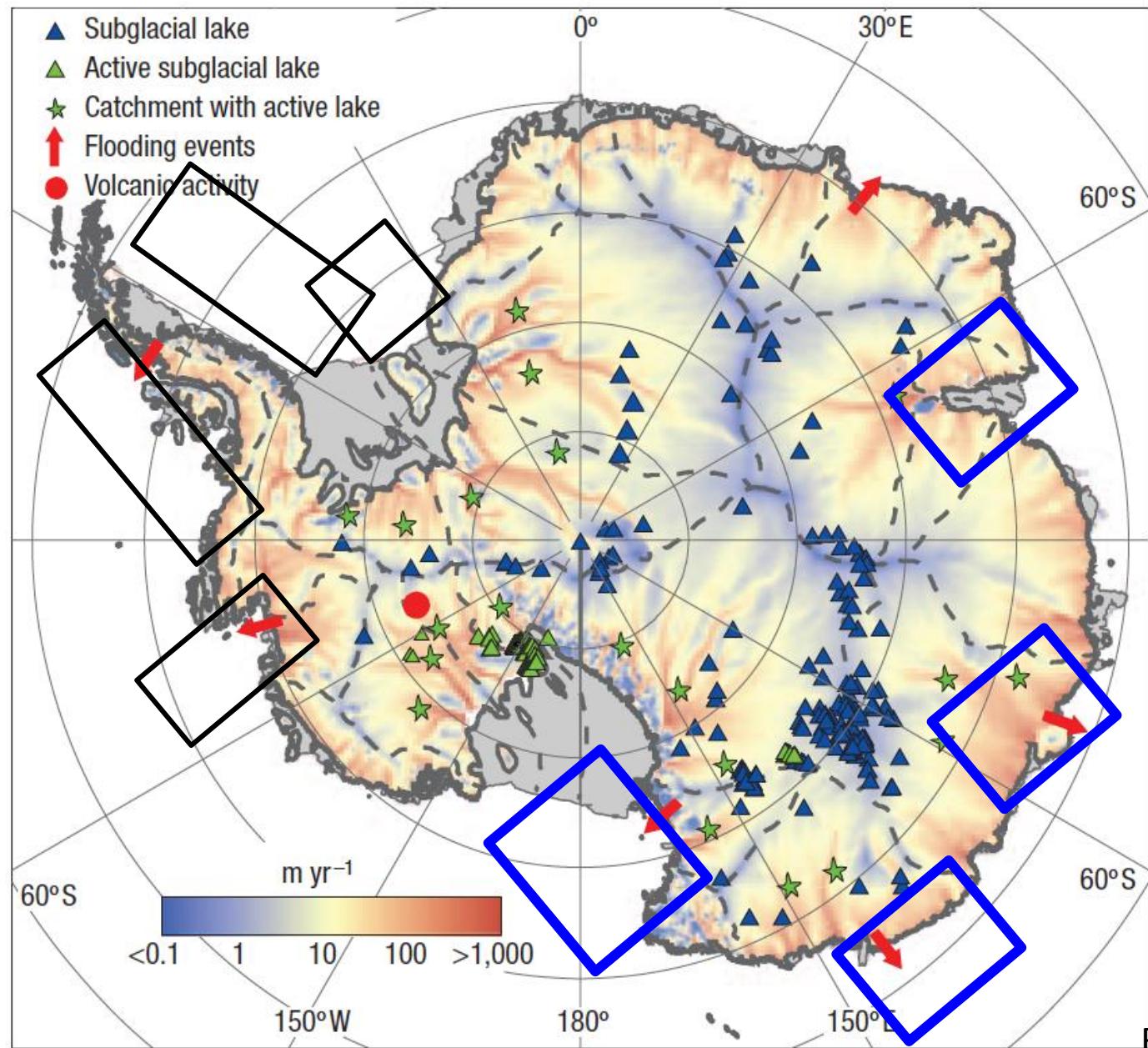


The low solubility of helium and neon in seawater results in **concentrations well above solubility equilibrium with the atmosphere** (the primary source of helium and neon in the ocean), producing a glacial melt signal which can be traced from an ice front, across the continental shelf and into the abyssal ocean.

^{4}He : Nuclear reactions in rocks and sediments
From the decay of U and Th, geothermal
Activity, volcanic eruptions (e.g. **high $[^{4}\text{He}]$**
are common in deep ground water discharge)

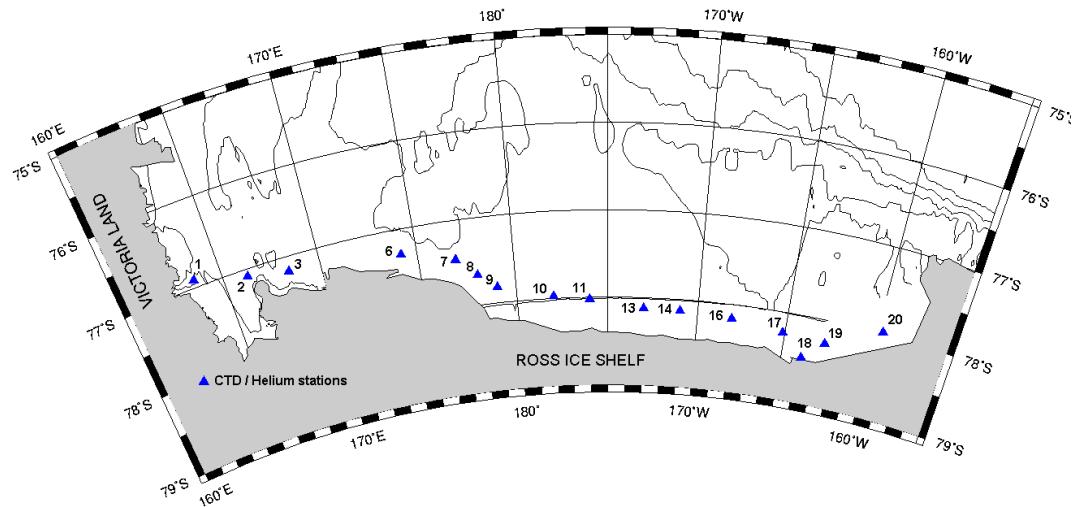
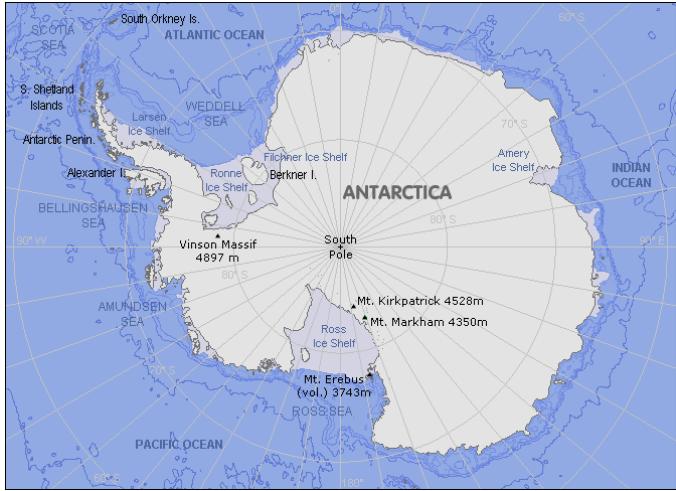
^{3}He : Radioactive decay of atmospheric tritium



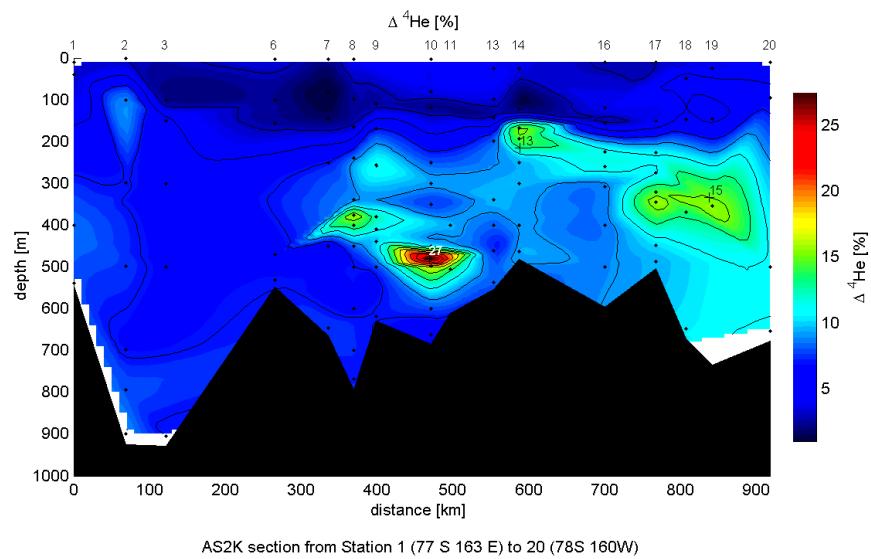
a

Bell, 2008

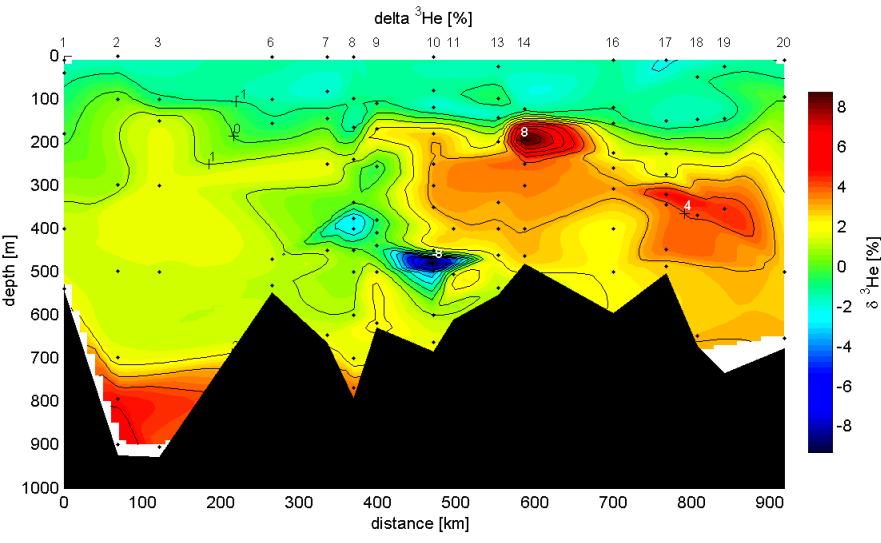
RIS: Palmer 2000 section



Helium isotope sections

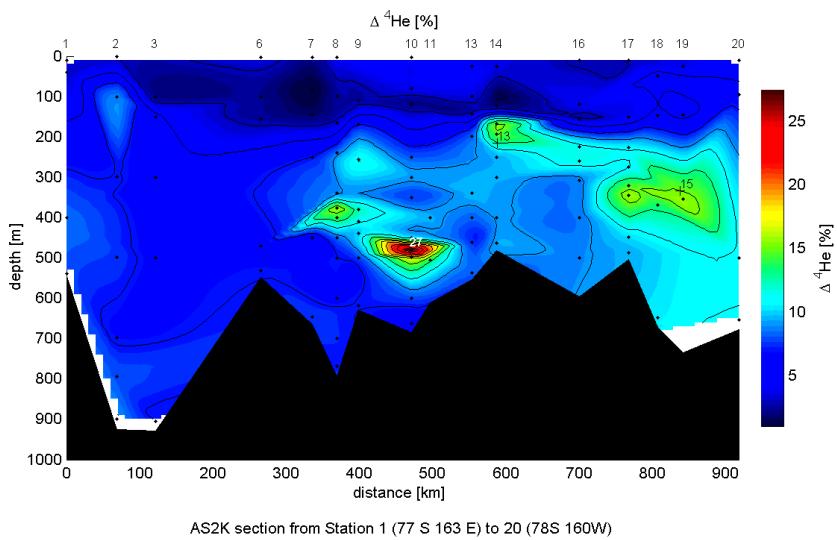


**Distinct ${}^4\text{He}$ excesses
in ISW cores; larger
than Ne excesses**

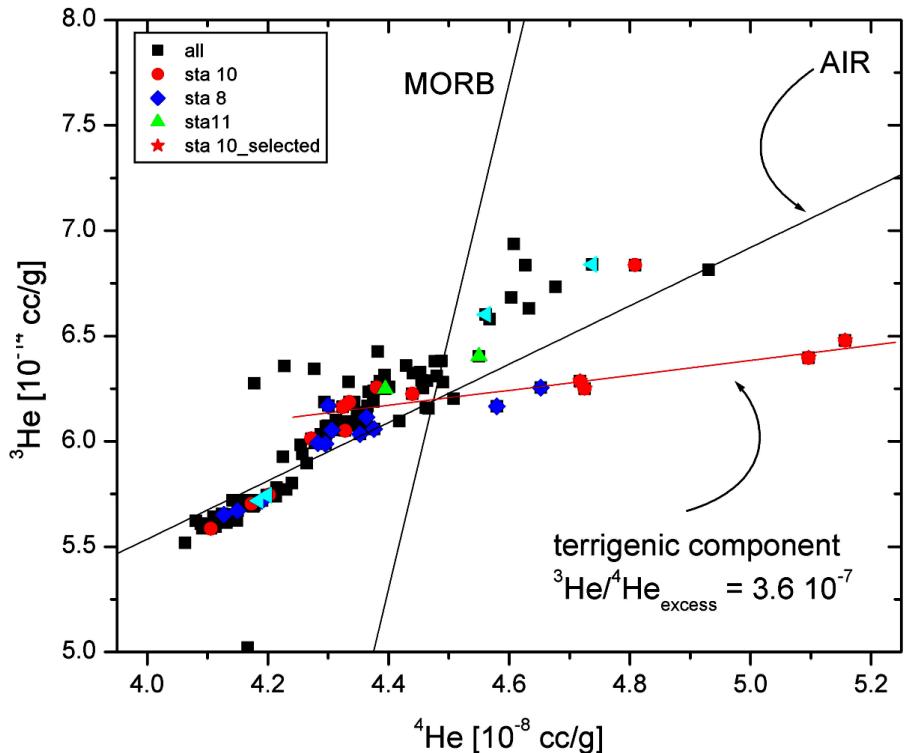


**$\delta^3\text{He}$ signals related to ${}^4\text{He}$ excesses
in ISW cores indicate addition of
terrigenic helium of crustal origin**

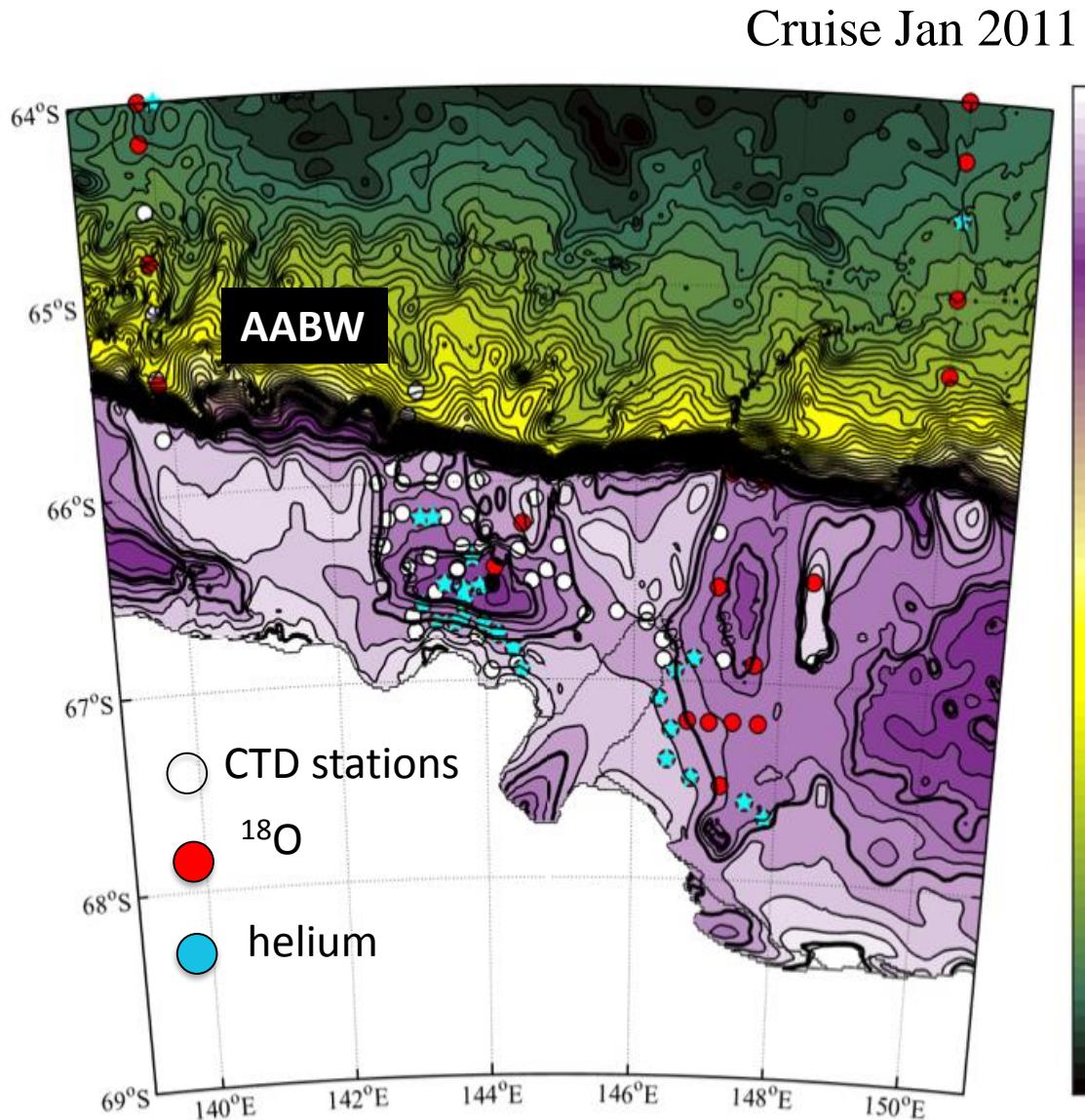
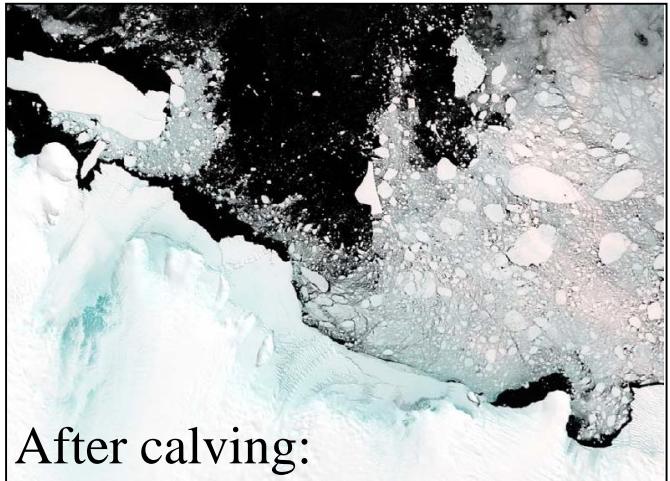
Terrigenic helium signal



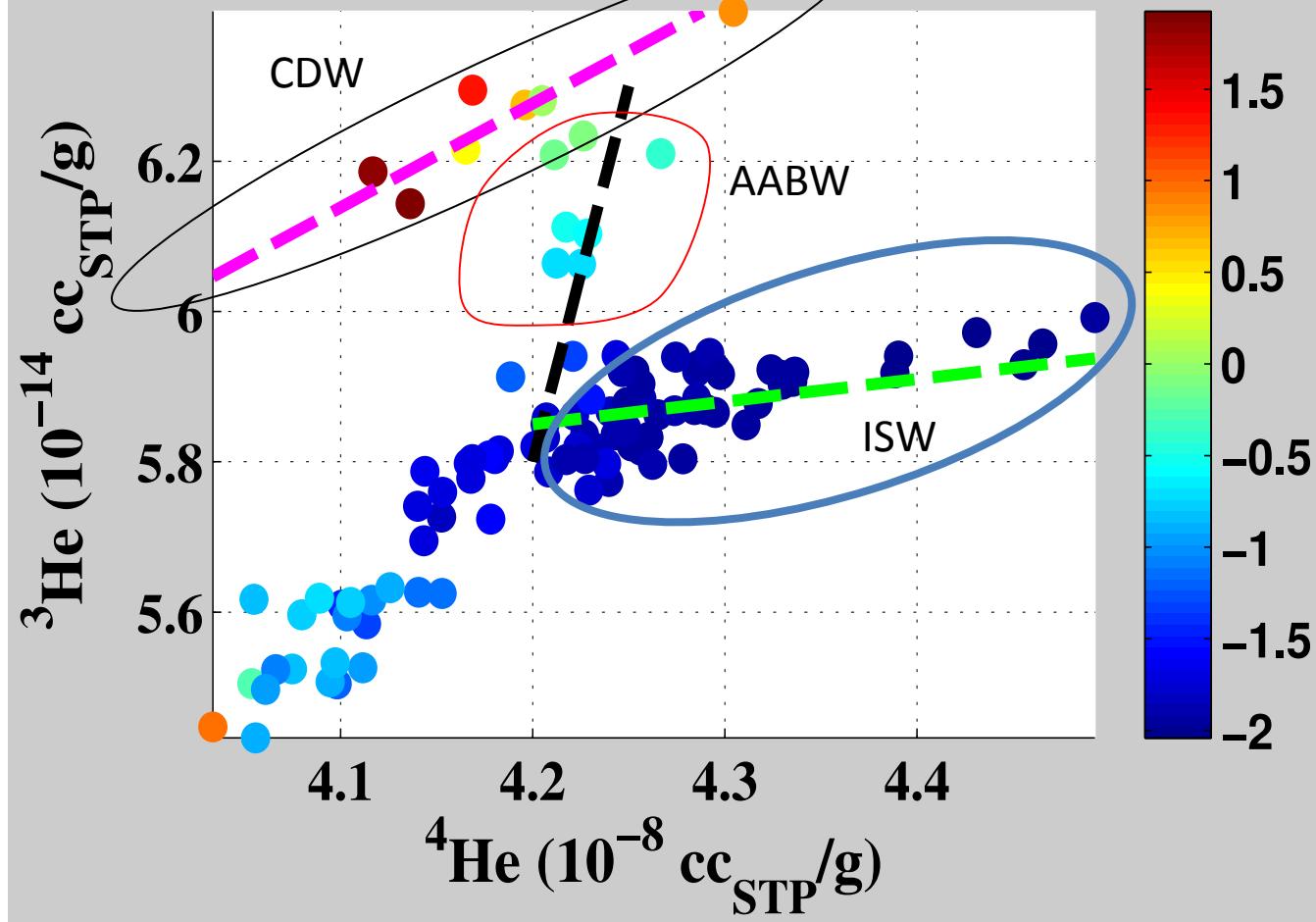
We can use this data to quantify the contribution of the subglacial freshwater contribution to the observed meltwater plume exiting the ice shelves cavities



Mertz Glacial Tongue



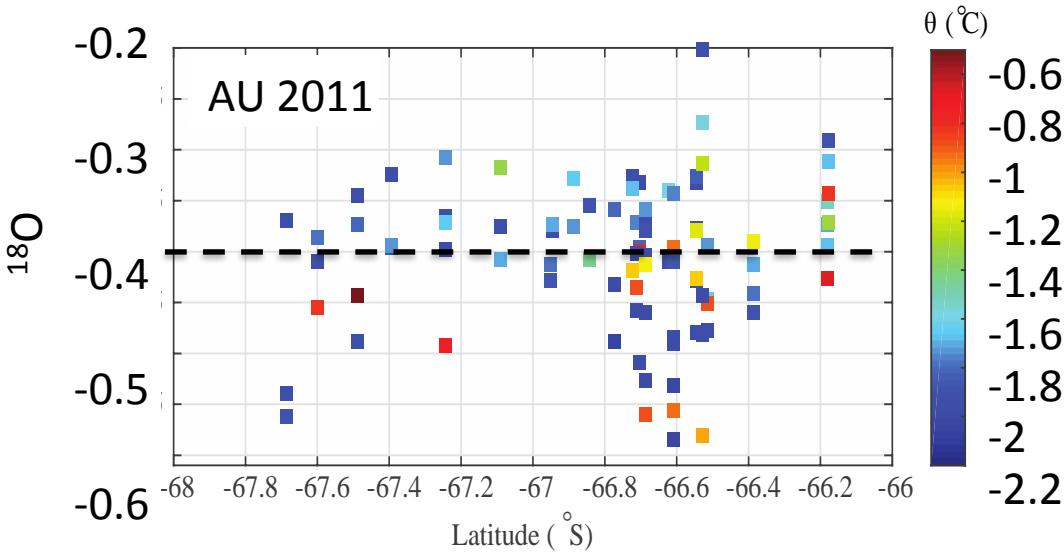
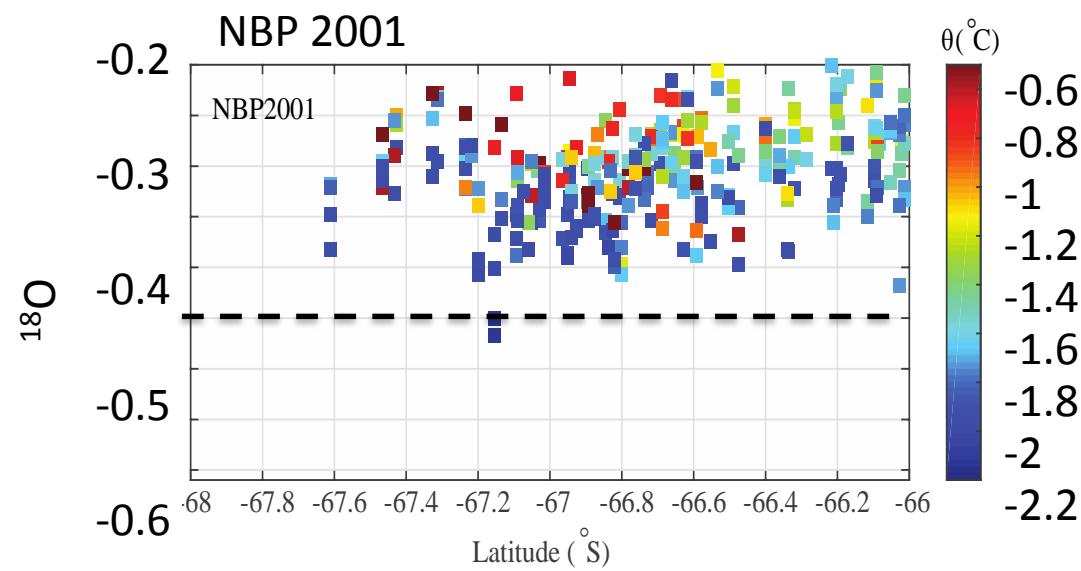
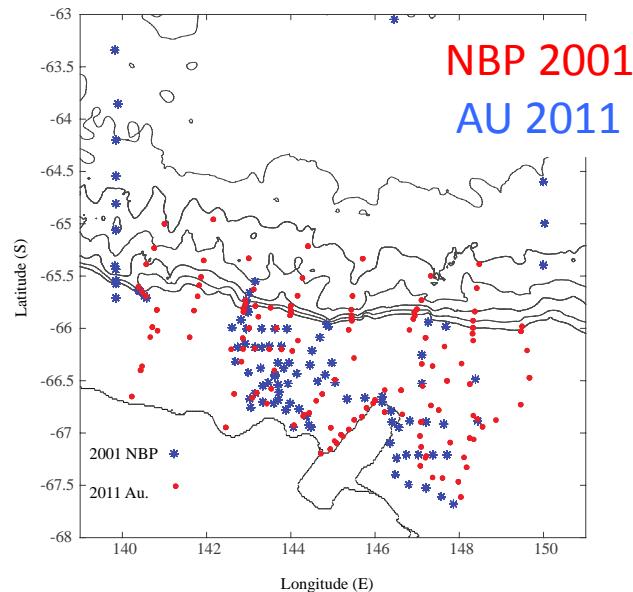
AU1121, plotted with mixing lines: Air (m), Mantle (k), Terrig (g) He θ



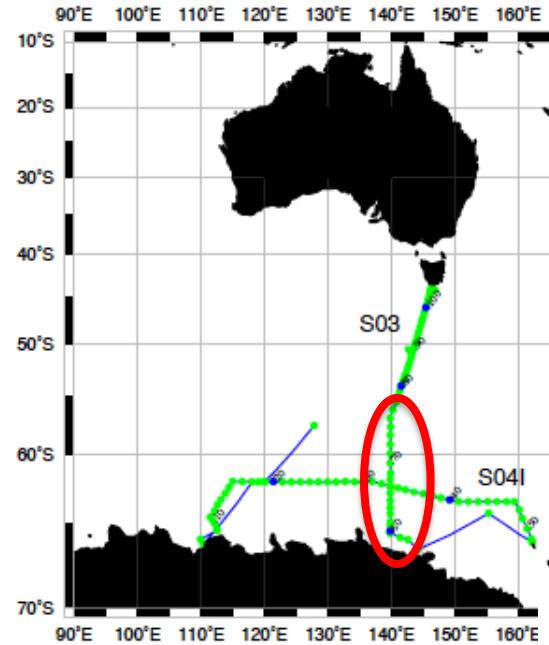
Circumpolar Deep Water (CDW)

Antarctic Bottom Water (AABW)

Ice Shelf Water (ISW)



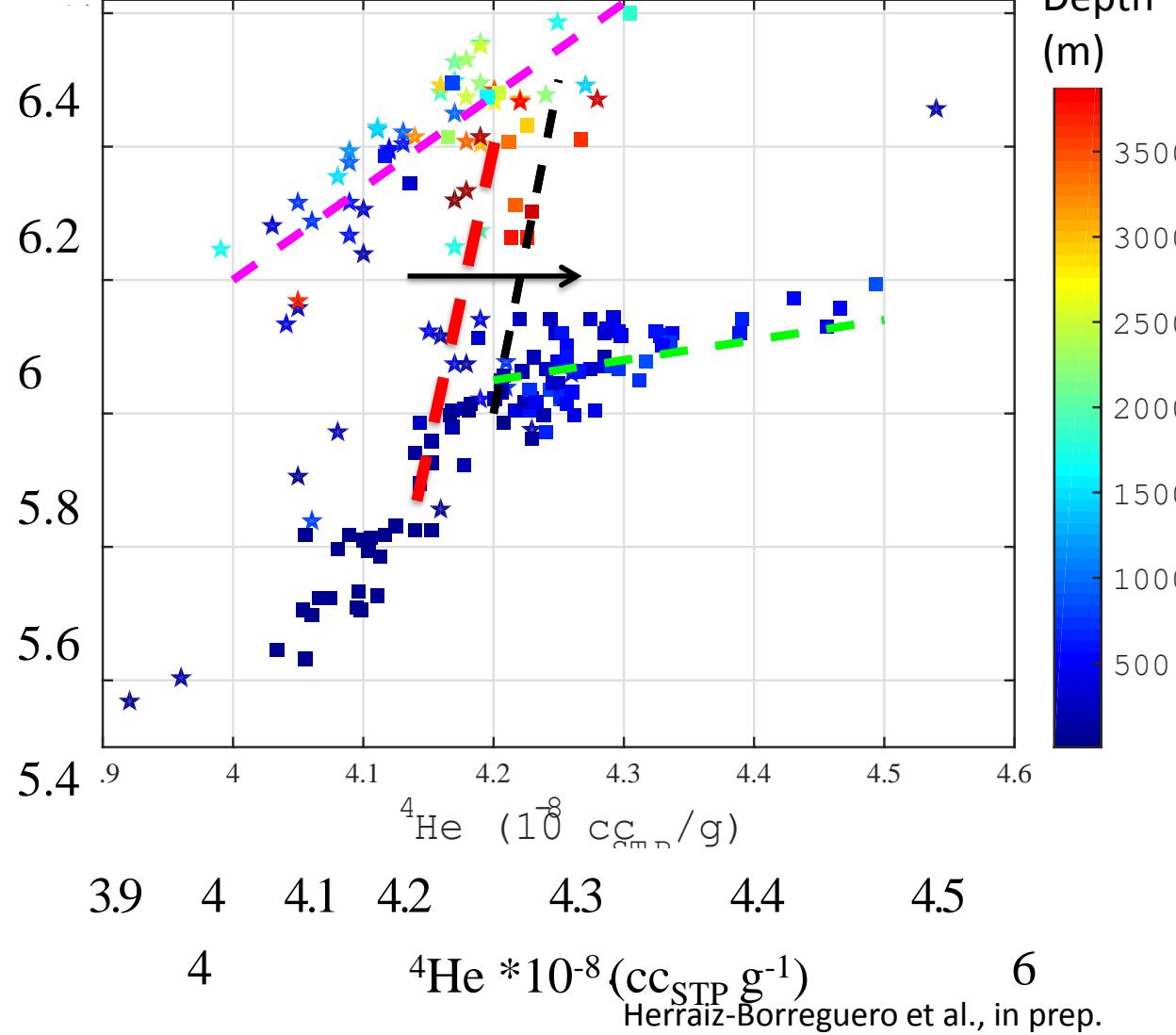
Parameter	Mean f_{MGW} ($\pm 1\text{std}$) per mil	Max f_{MGW} Per mil	AABW f_{MGW} Per mil
2011 MC, θ , S, ^{18}O	1.88 ± 1.73	6.57	1.56
2001 MC, θ , S, ^{18}O	0.88 ± 0.72	4.00	0.6



★ 1994

□ 2011

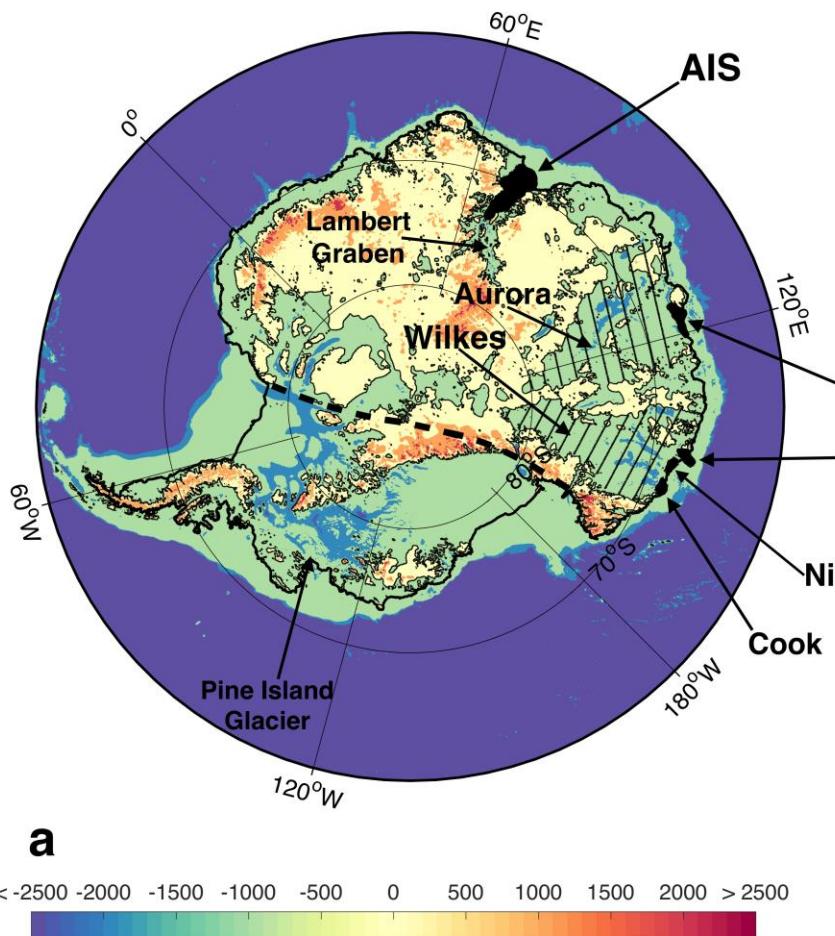
${}^3\text{He} * 10^{-14} (\text{cc}_{\text{STP}} \text{ g}^{-1})$



Conclusions

1. Need more work on how currents interact with ice shelves fronts and what it means for the inflow of waters
2. Polynyas:
 1. ocean stratification matters
 2. Freshwater can actually hamper the formation of dense shelf waters
3. Noble gases gives us information on the source of the glacial freshwater and how this freshwater is exported
4. It can inform models

Bed Topography (m)

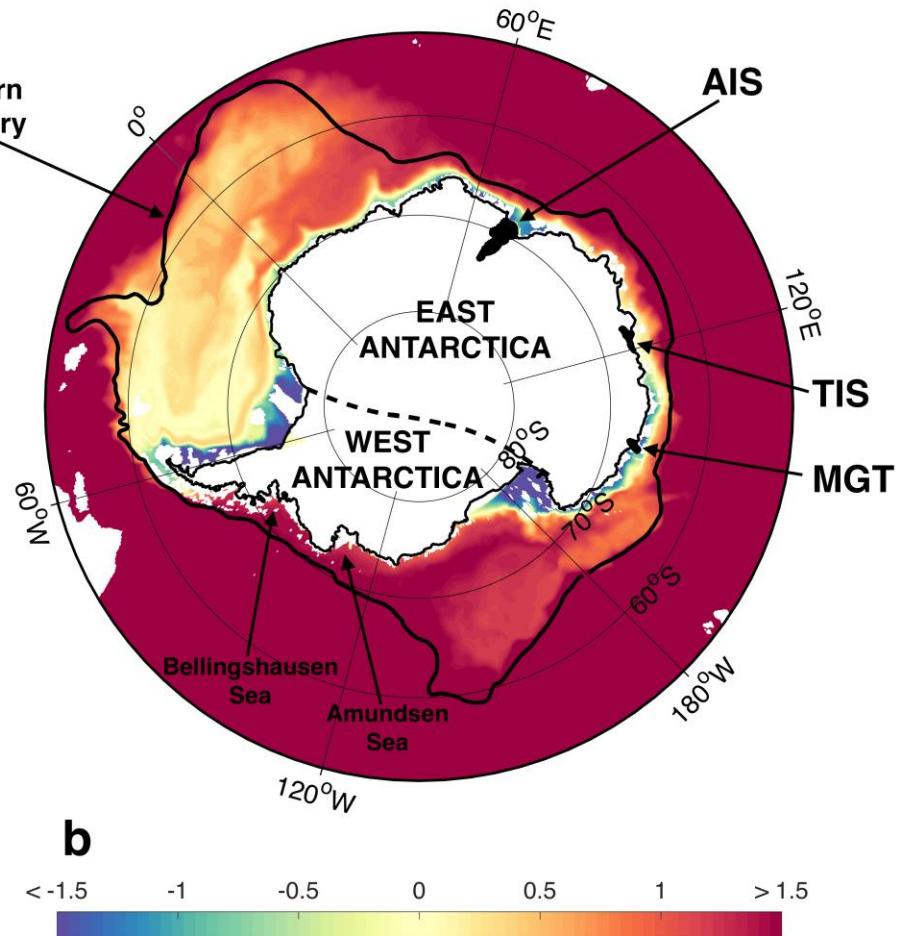


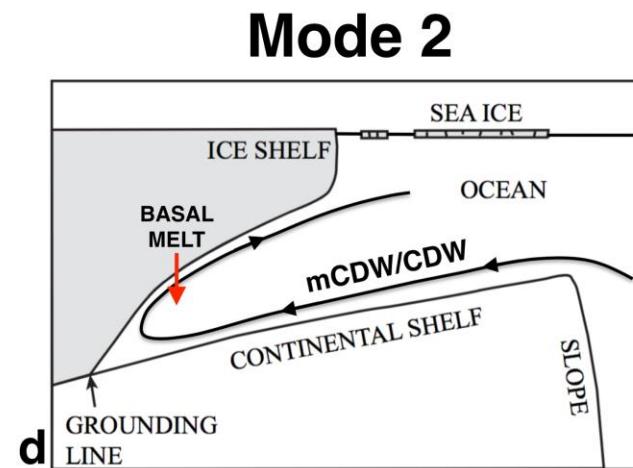
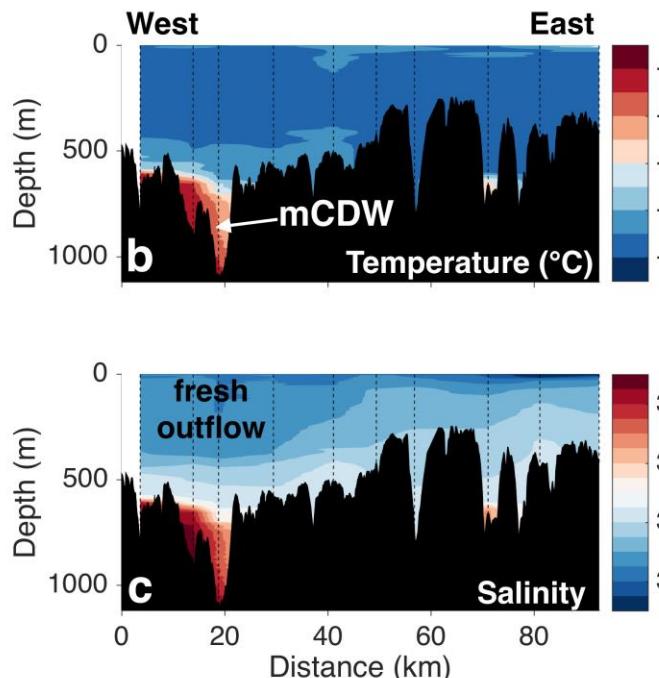
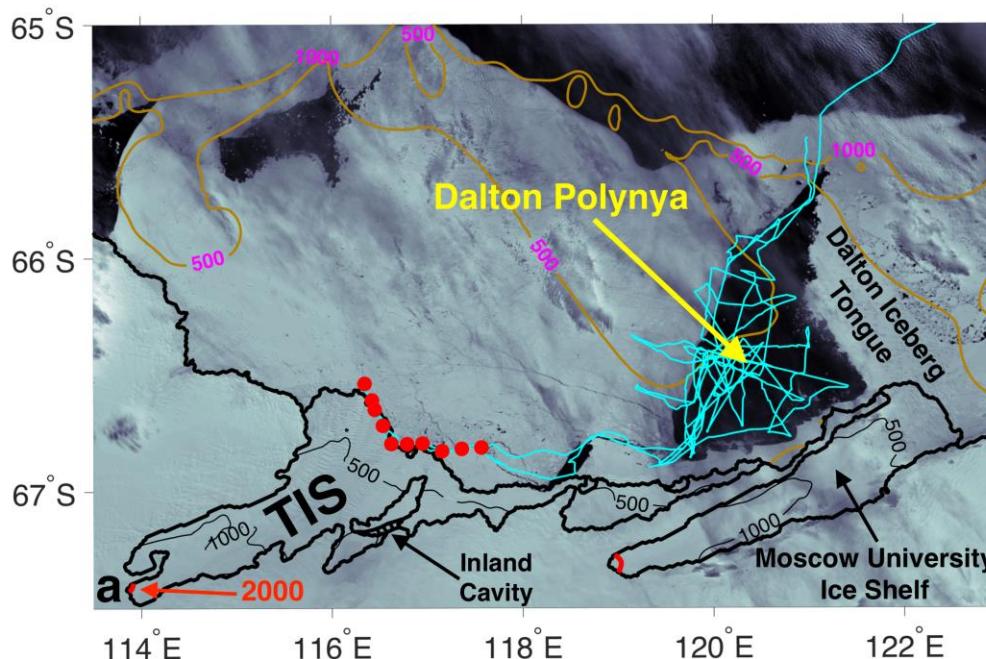
AIS: Amery ice shelf

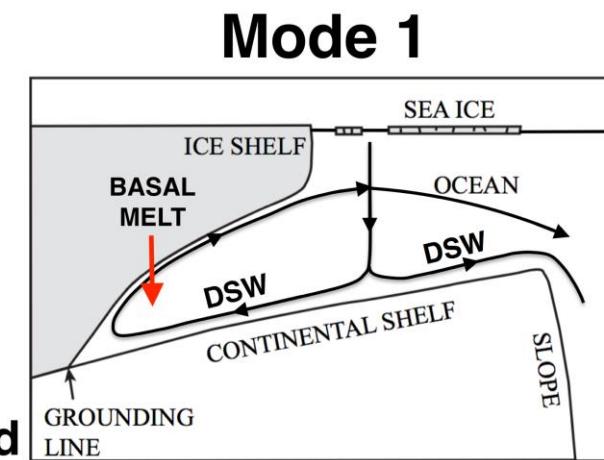
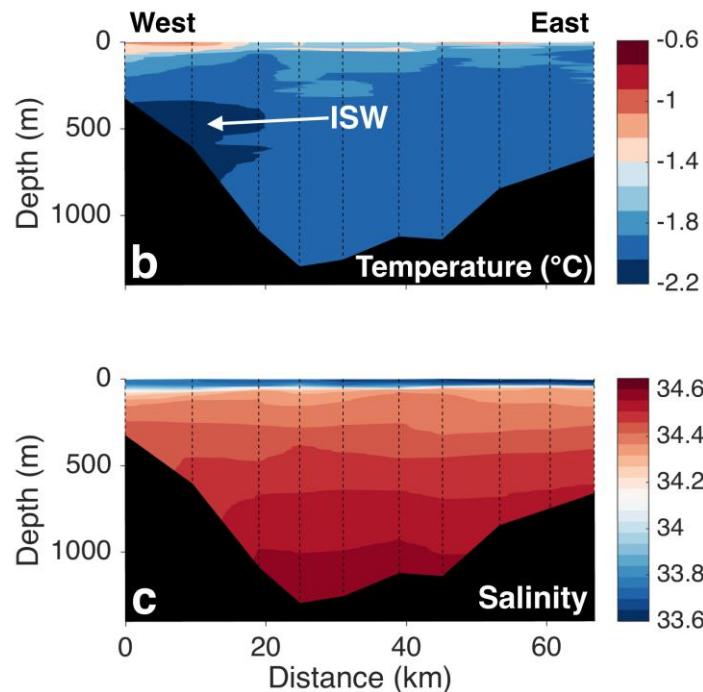
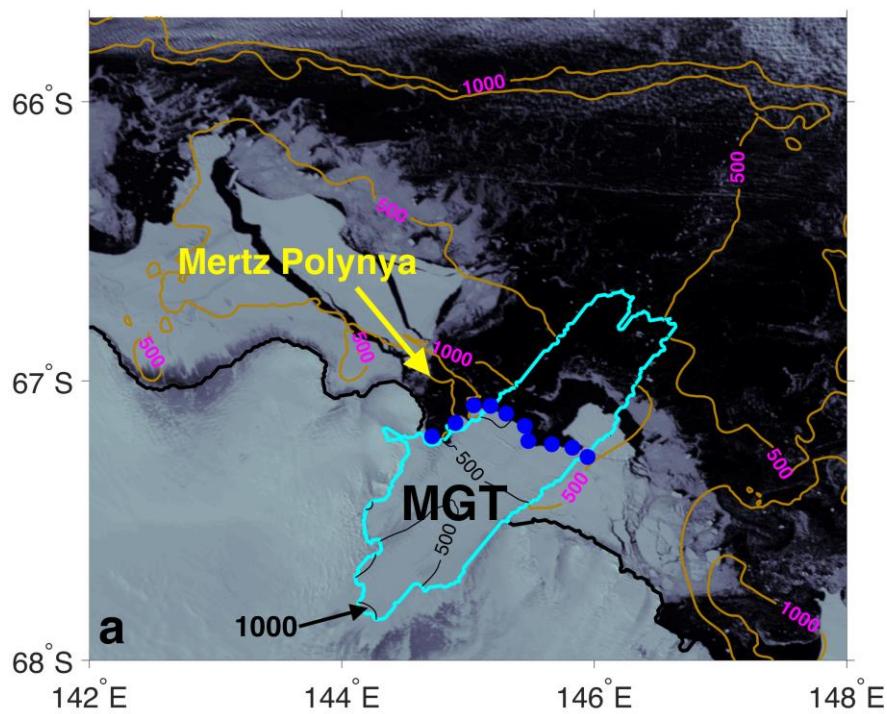
TIS: Totten ice shelf

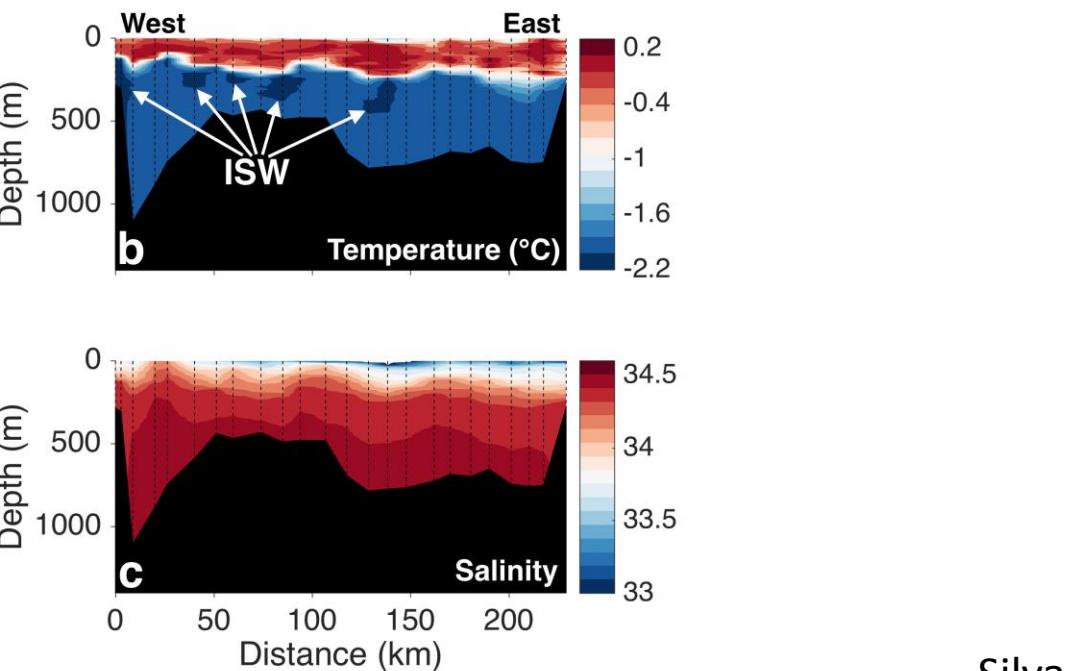
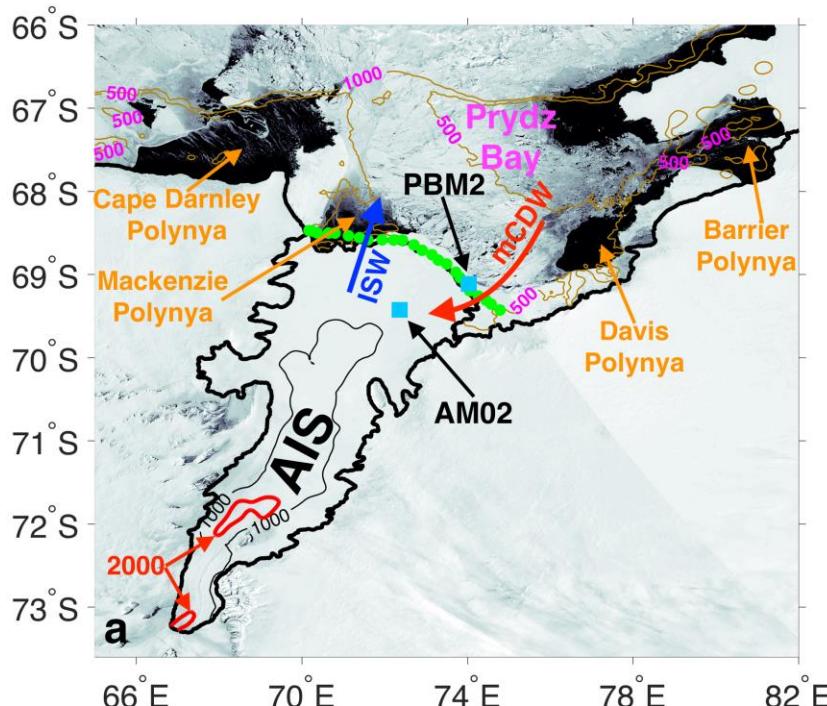
MGT: Mertz Glacial Tongue

Ocean Temperature (°C)

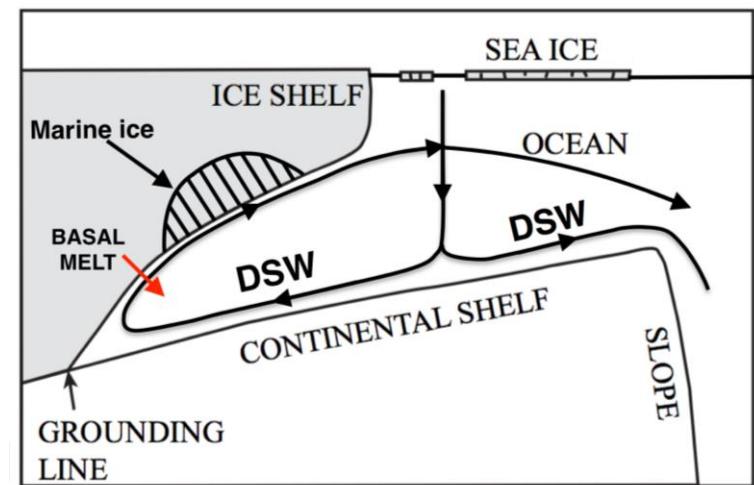




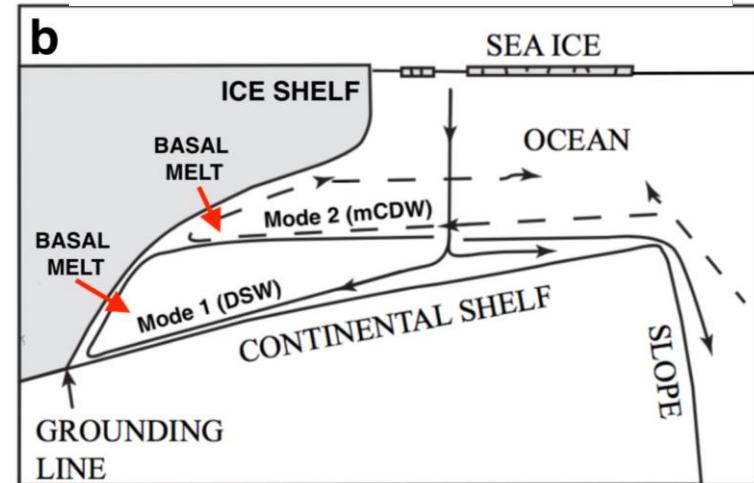




Mode 1

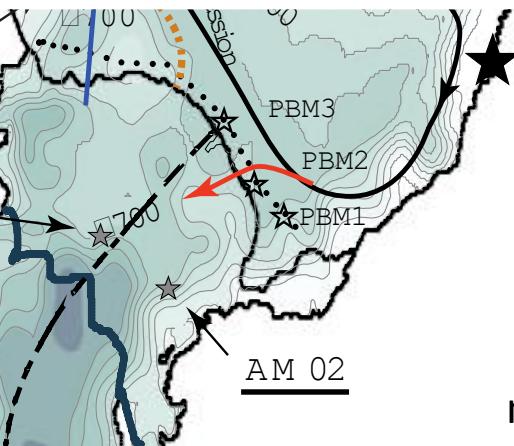


Mode 2

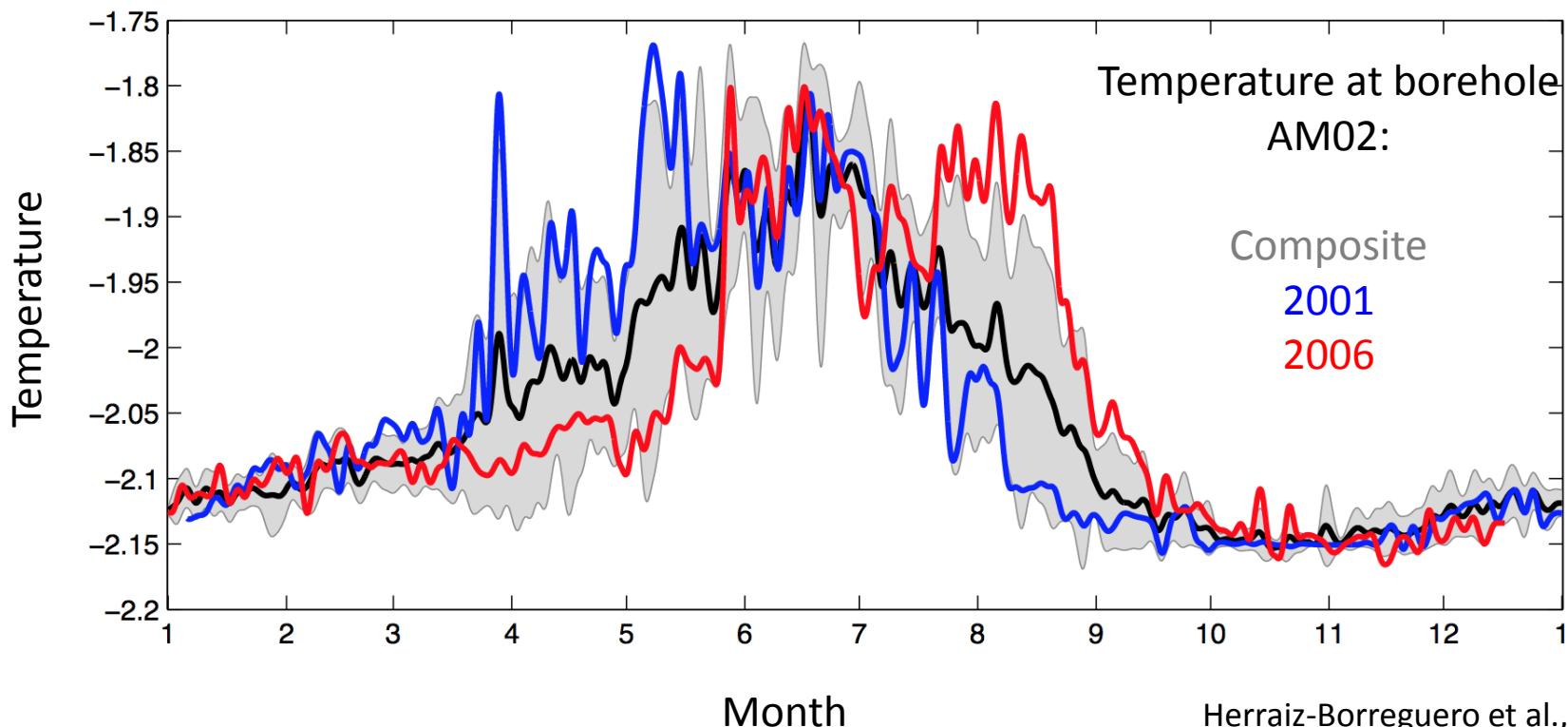


mCDW is responsible for up to 2 ± 0.5 m yr $^{-1}$ during 2001 (23.9 ± 6.52 Gt yr $^{-1}$).

However, heat content flux by mCDW at AM02 shows high intra-annual variability (up to $\pm 40\%$)



mCDW at the ice shelf front



Prydz Bay-Amery ice shelf system: what do we know ?

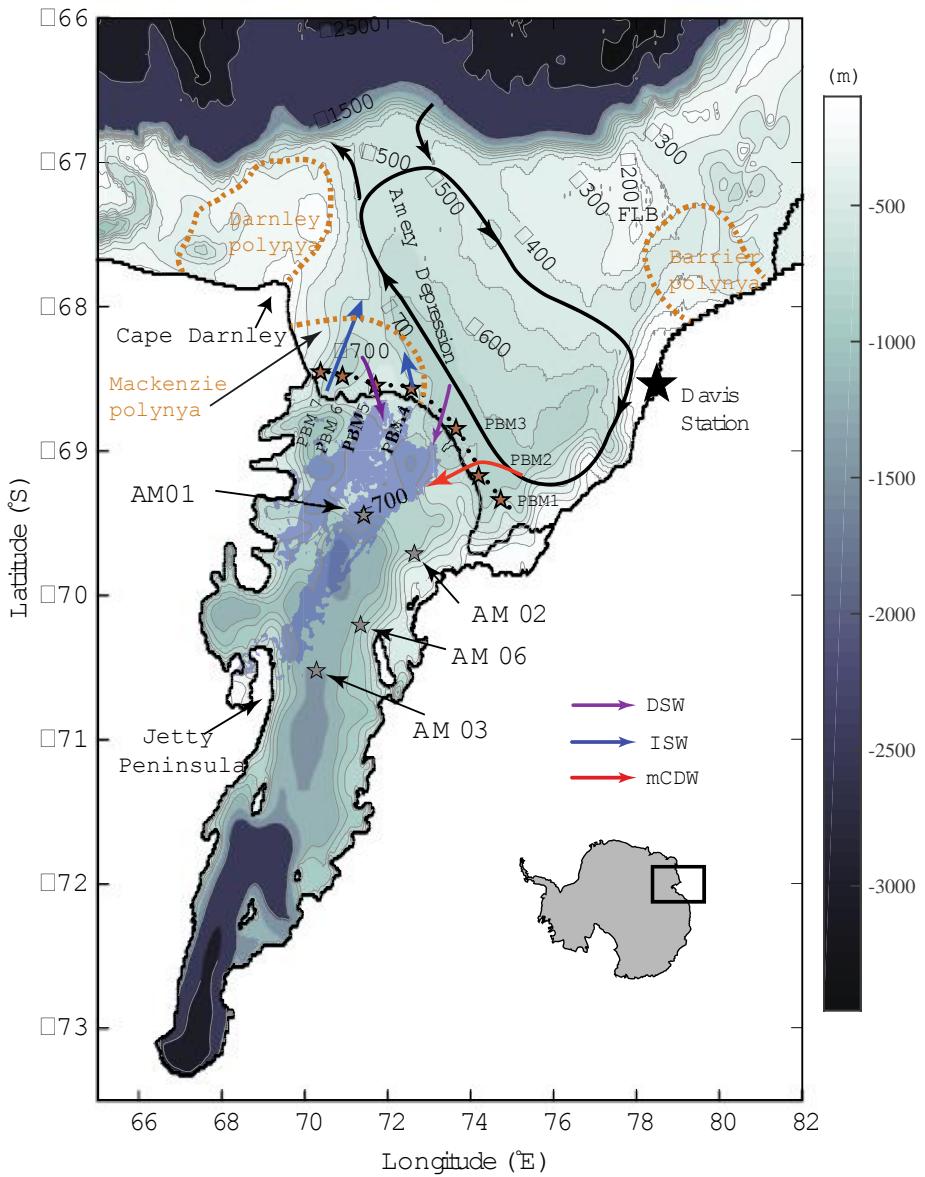


Table 2. Summary of the Most Recent Estimates of Net Basal Melt Beneath the Amery Ice Shelf

Study	Net Basal Melt (m yr^{-1})	Net Basal Mass Loss (Gt yr^{-1})
This study + Herraiz-Borreguero et al. [2015] ^{a,b}	1.0 ± 0.4	57.4 ± 25.3
Depoorter et al. [2013] ^c	0.65 ± 0.35	39 ± 21
Rignot et al. [2013] ^c	0.58 ± 0.4	35.5 ± 22
Galton-Fenzi et al. [2012] ^d	0.74	45.6
Yu et al. [2010] ^c	0.5 ± 0.12	27 ± 7
Wen et al. [2010] ^c	0.84 ± 0.12	46.4 ± 6.9

^aOceanographic study.

^bHas been adjusted to give a net basal melt estimate in m/yr over the whole ice shelf area.

^cGlaciological study.

^dModeling study.

Glacial Melt Water composition: Optimal MultiParameter (OMP) analysis

$$\begin{aligned}
 f_{CDW} &+ f_{AASW} &+ f_{DSW} &+ f_{MGW} &= 1 \\
 f_{CDW} * \theta_{CDW} &+ f_{AASW} * \theta_{AASW} &+ f_{DSW} * \theta_{DSW} &+ f_{MGW} * \theta_{MGW} &= \theta_{obs} \\
 f_{CDW} * S_{CDW} &+ f_{AASW} * S_{AASW} &+ f_{DSW} * S_{DSW} &+ f_{MGW} * S_{MGW} &= S_{obs} \\
 f_{CDW} * {}^{18}\text{O}_{CDW} &+ f_{AASW} * {}^{18}\text{O}_{AASW} &+ f_{DSW} * {}^{18}\text{O}_{DSW} &+ f_{MGW} * {}^{18}\text{O}_{MGW} &= {}^{18}\text{O}_{obs} \\
 f_{CDW} * {}^4\text{He}_{CDW} &+ f_{AASW} * {}^4\text{He}_{AASW} &+ f_{DSW} * {}^4\text{He}_{DSW} &+ f_{MGW} * {}^4\text{He}_{MGW} &= {}^4\text{He}_{obs}
 \end{aligned}
 \quad \left. \right\} E * x = y$$

E: matrix of end-members properties

x: vector of unkowns (**f**)

y: observations

} least square method to resolve our system.

4 water masses used (or end-members):

1. Circumpolar Deep Water (CDW)
2. Antarctic Surface Water (AASW)
3. Dense Shelf Water (DSW)
4. Glacial Melt Water (GMW)

4-5 parameters: Mass conservation, Potential temperature (θ), salinity (S), ${}^{18}\text{O}$ & ${}^4\text{He}$ conc.
contribution of the 4 water masses; f1, f2, f3 and f4

We are not using the data at the top 200 m of the water column

Glacial Melt Water composition:

Cruise 2011

Parameters	Mean fG_{MW} ($\pm 1\text{std}$) per mil	Maximum f_{GMW} Per mil
MC, θ , S, ^{18}O , ^4He	1.22 ± 1.04	4.46
MC, θ , S, ^{18}O	1.88 ± 1.73	6.57
MC, θ , S, ^4He	1.21 ± 1.07	4.91
MC, θ , ^{18}O , ^4He	2.20 ± 1.74	6.69
MC, S, ^{18}O , ^4He	1.46 ± 1.04	4.78



ISW $f_{MGW} = 6.57$ per mil

AABW $f_{MGW} = 1.56$ per mil

GMW fractions, fluxes and melt rates

	Ross Sea		Weddell Sea
	Deep ISW core	Shallow ISW core	Deep ISW (WSW/ISW)
Salinity	4.3 ‰	4.8 ‰	2.9 ‰
Neon	3.9 ‰	4.9 ‰	3.8 – 7 ‰ (${}^4\text{He}$)
$\delta^{18}\text{O}$	3.9 ‰	4.6 ‰	2.8 - 6 ‰

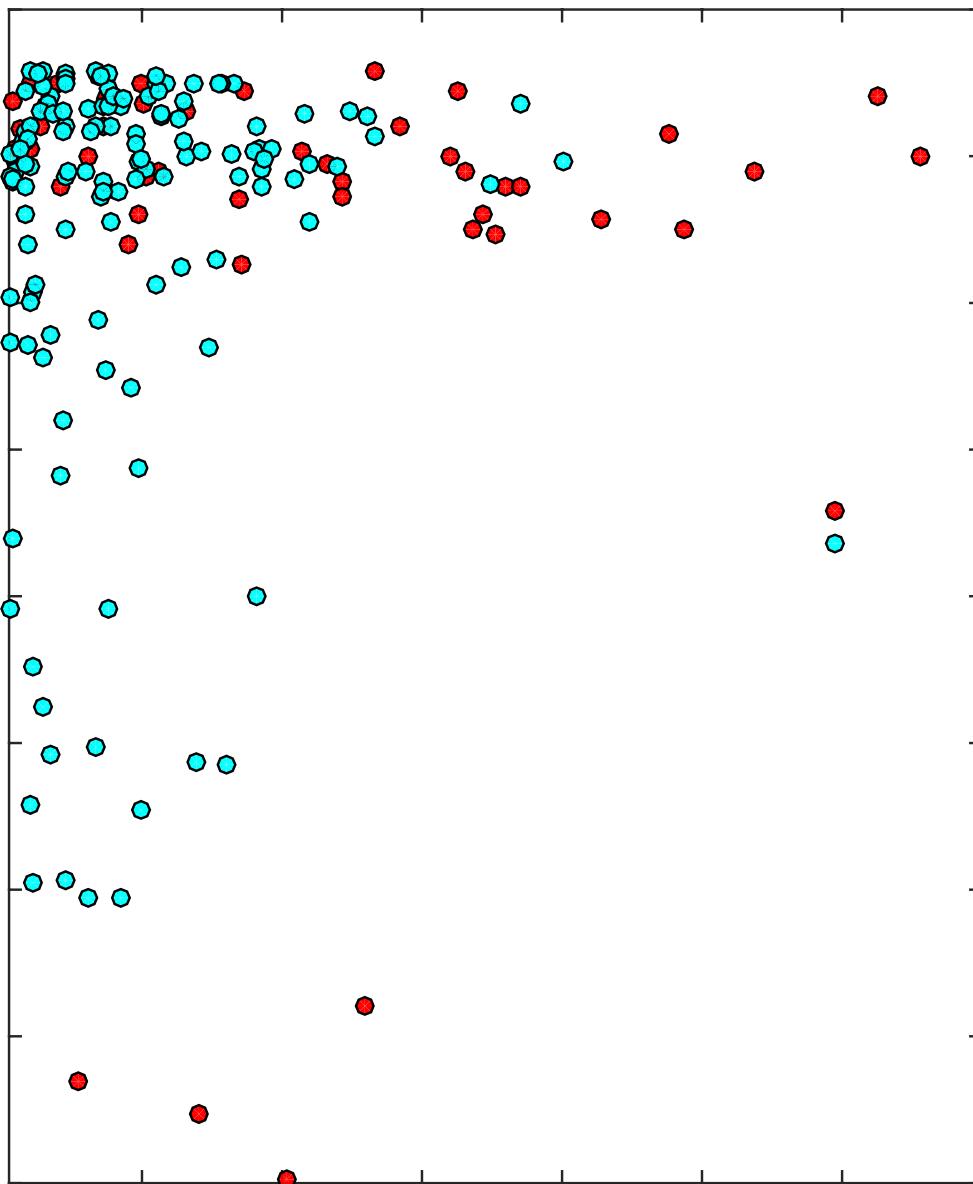
No sign for re-freezing underneath the RIS

Mean residence time: ca. 3.5 years (Bill Smethie)

Deep core: ISW flux: ca. 0.1 Sv

GMW flux: 0.4 mSv

Melt rate for 100 km wide pathway: 0.1 m year⁻¹



x

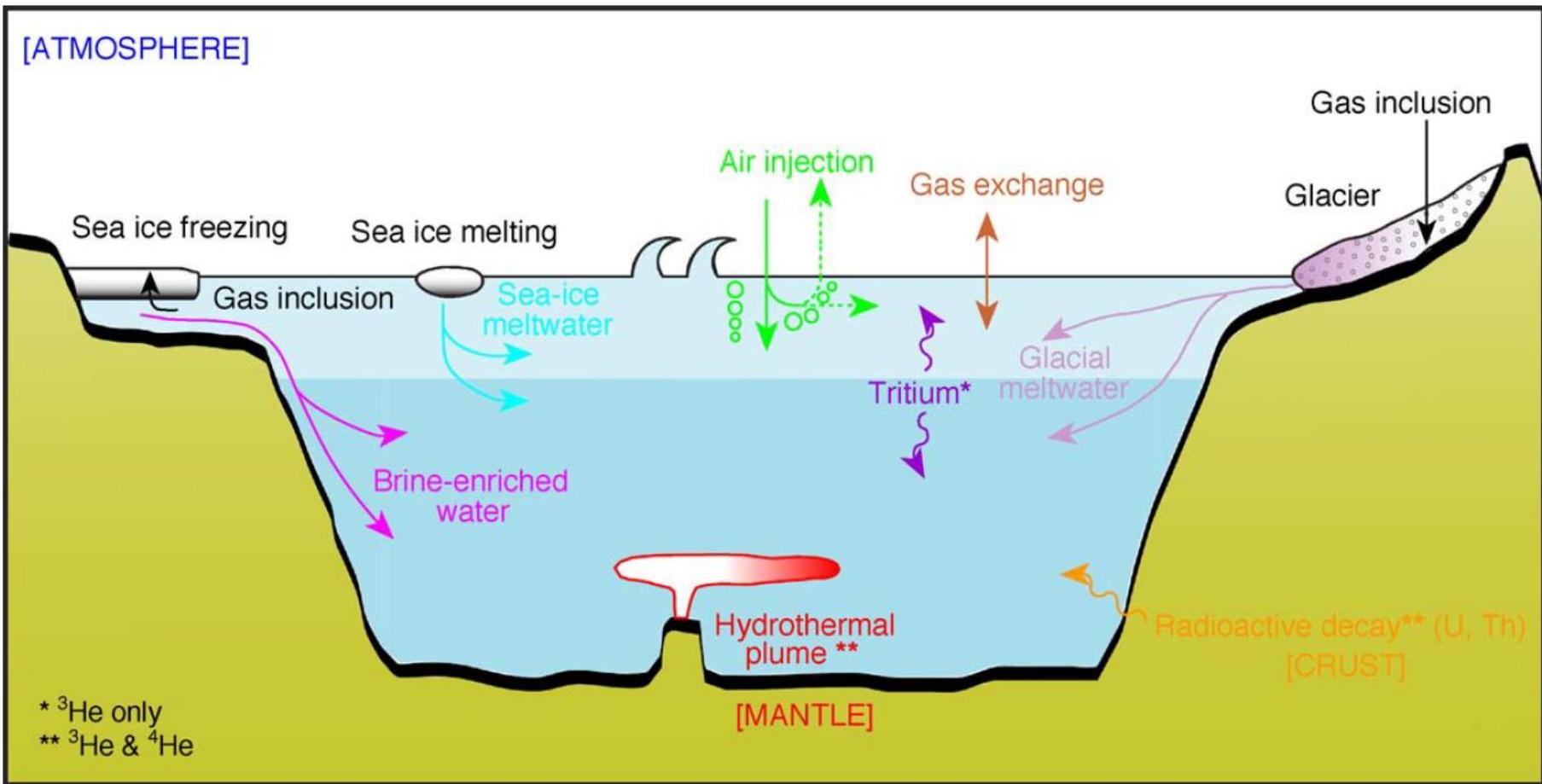


Fig. 1. Mechanisms controlling the distribution of helium and neon in the arctic seas. Helium isotopes have four sources: the atmosphere, mantle, crust and decay of tritium; however, neon has only one, the atmosphere. Helium and neon can be applied to explore the processes which can change the concentration of atmospheric gases, such as air injection, brine injection, sea-ice melting and glacier melting.